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Consiglio Nazionale
delle Ricerche



Istituto di
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Carla Braitenberg



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Mannino



Luigi
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Gabriella
Mangano



Gianmaria
Sannino



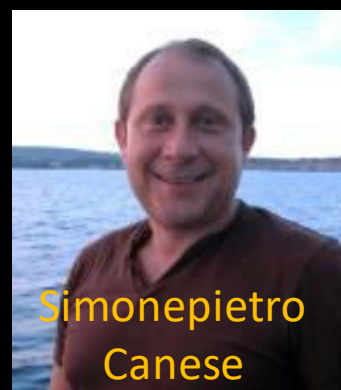
Kurt
Lambeck



Francesco Latino
Chiocci



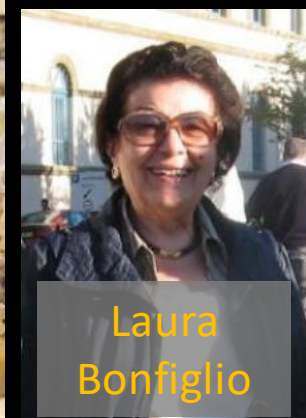
Fabrizio
Antonioli



Simonepietro
Canese



Mariarita
Palombo



Laura
Bonfiglio



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Catalano



Valeria
Lopresti



Maurizio
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Renato Tonielli

Paolo Orrù ed Emanuela Solinas



Eleonora de Sabata



Egidio Trainito



Gianfranco Scicchitano
e
Luigi Ferranti

Surveys and publications with marine archaeologists:

- **Sebastiano Tusa** Sopr. of Trapani
- **Flavio Enei**, Museo di Pirgy
- **Marinella Pasquinucci**, University of Pisa
- **Jonathan Benjamin**, University of Newcastle
- **Heud Galili** University of Haifa, Israel
- **Timothy Gambin**, University of Malta
- **Vladimir Kovacic**, Museum of Porec
- **Silvia Ducci**, Sopr. Archeologica della Toscana
- **Alessandro Porqueddu**, Archaeologist
- **Rubens Doriano**, Sopr. Archeologica Olbia



Irena Radic
University of Zadar



Elena F. Castagnino
University of Catania



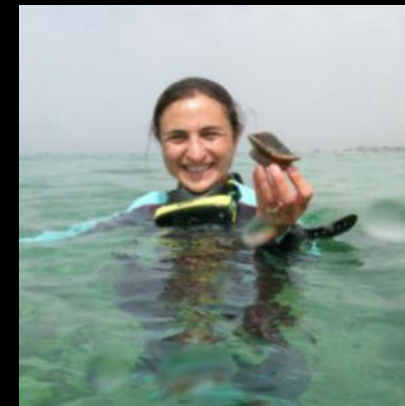
Flavio Enei
Museo di Pirgy



Emanuela Solinas
Museo di Cagliari, Italy



Alessandra Benini
University of
Calabria, Italy



Rita Auriemma
University
of Lecce, Italy

Sebastiano Tusa presso la grotta del Tuono, Marettimo





Presente, perchè

sale il mare? E' una risalita naturale o riscaldamento climatico dovuto all'uomo?

Passato

quali indicatori ci raccontano dove stava il mare, 2000, 20.000, 200.000, 2 milioni di anni fa?

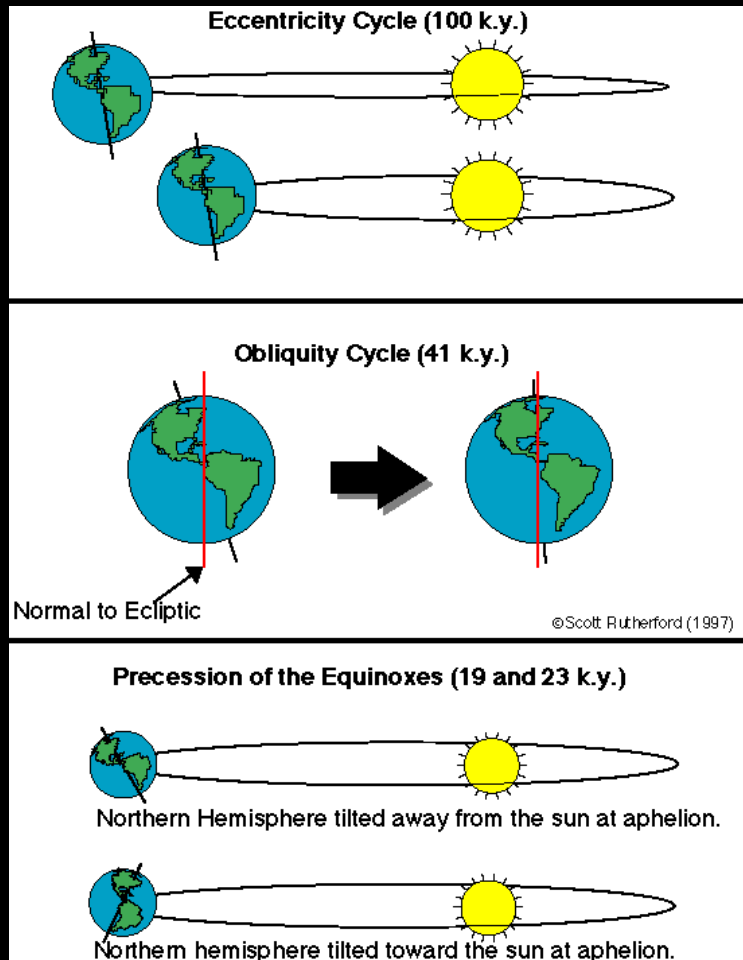
Presente

dati strumentali mareografi satellite;

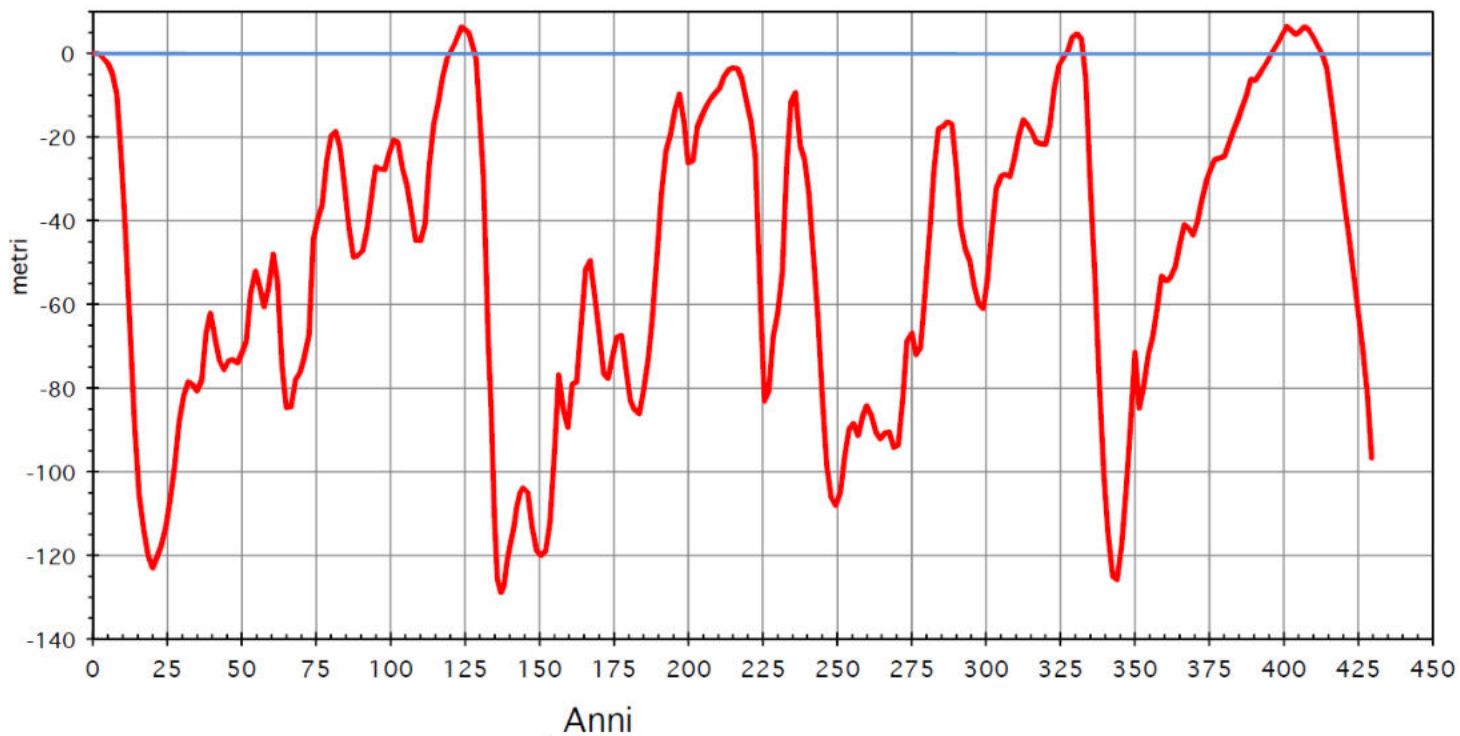
Futuro

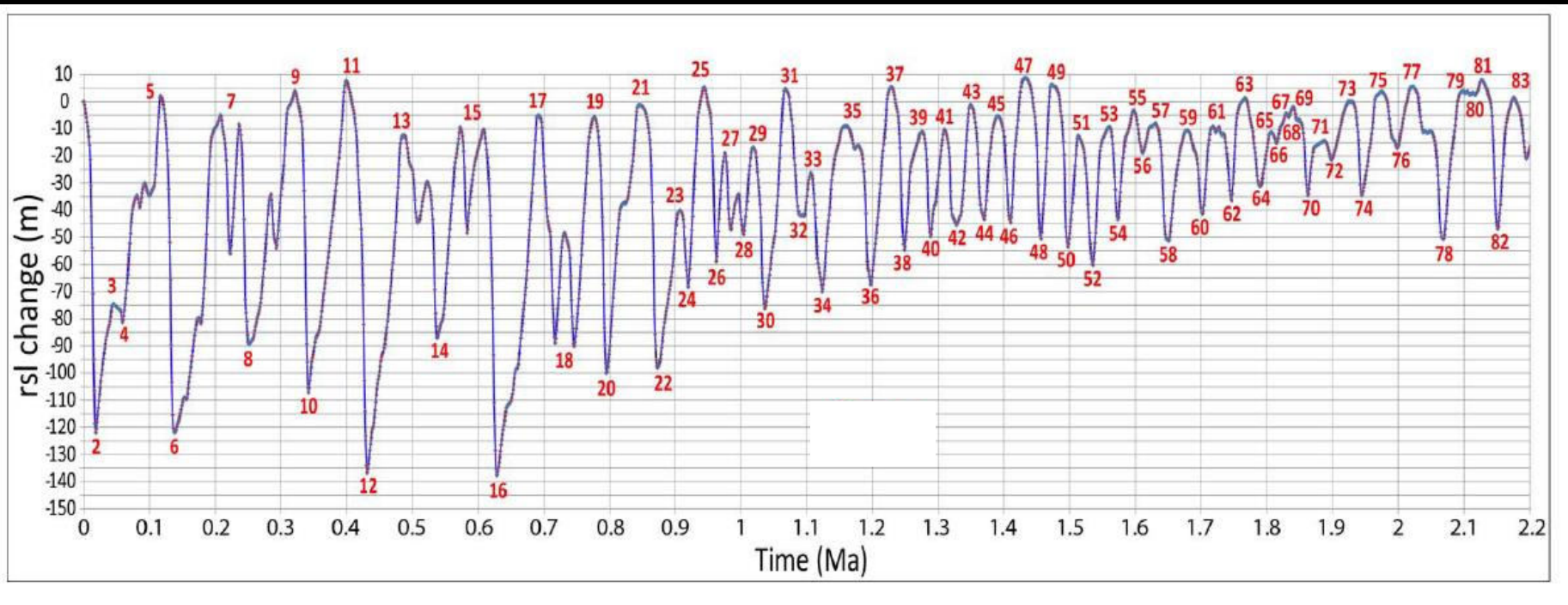
previsioni IPCC fino al 2100-2300

Milankovitch matematico Sloveno che a inizio secolo intuì che le variazioni del clima erano dovute alle variazioni dell'asse terrestre. Tutte le sue teorie sono state confermate.



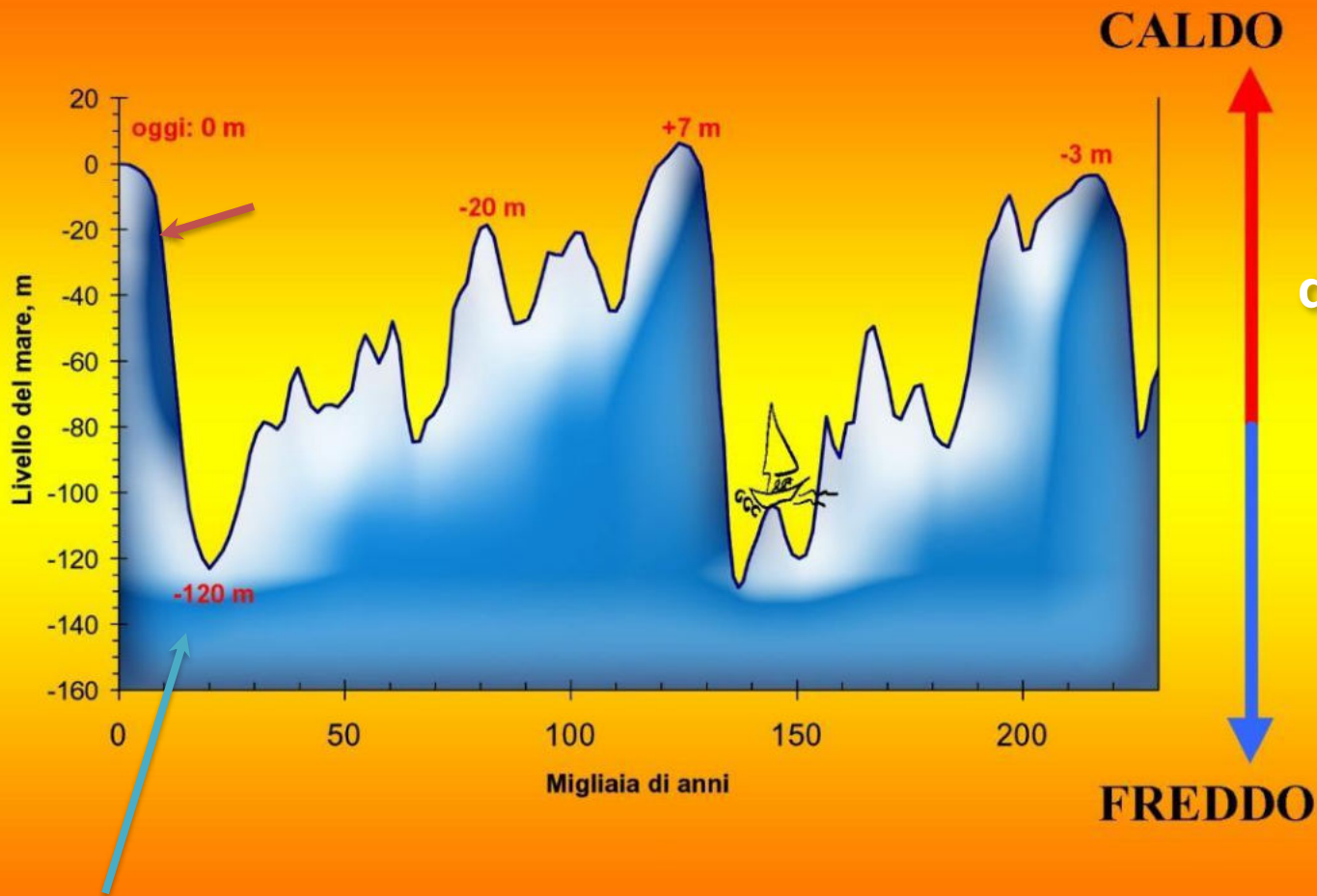
Variazioni del livello del mare da 450 mila anni fa al presente, da SPECMAP, Martinson et al., 1991 et al. 2001





Ocean Drilling Program (ODP), carote sui fondi





Le variazioni eustatiche durante gli ultimi 250.000 anni

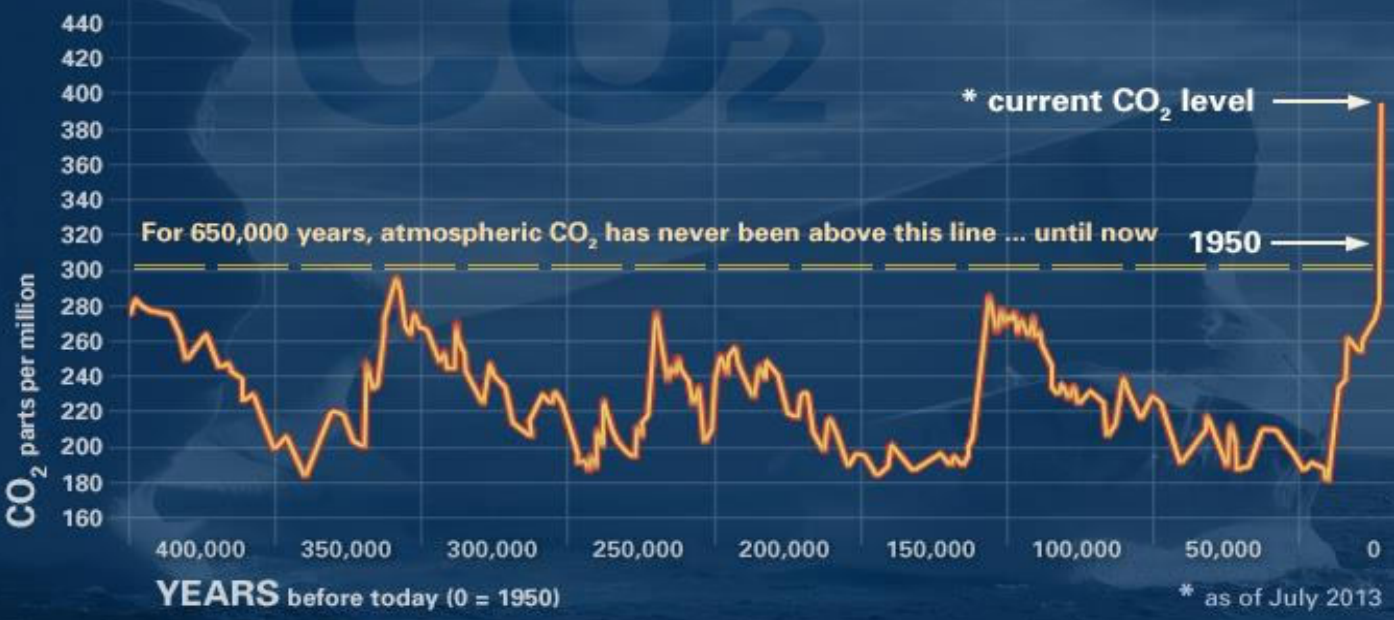
Carote di ghiaccio, Antartide



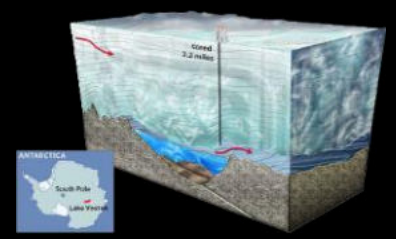
Carote di ghiaccio, Antartide



CO₂



Marzo 2020
415 ppm CO₂



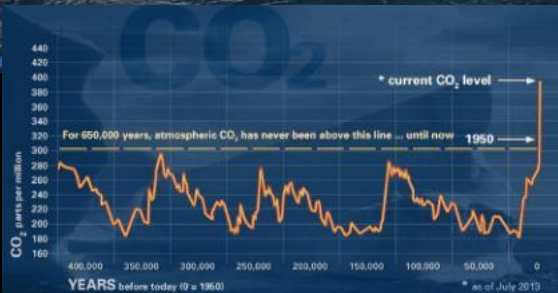
GLOBAL CLIMATE CHANGE

climate.nasa.gov

Sardegna, Golfo di Orosei

Solco di battente fossile, 125.000 anni fa: + 8 m
290 parti per milione di CO²

Solco di battente attuale :
424 parti per milione di CO²





WMO



UNEP

ipcc

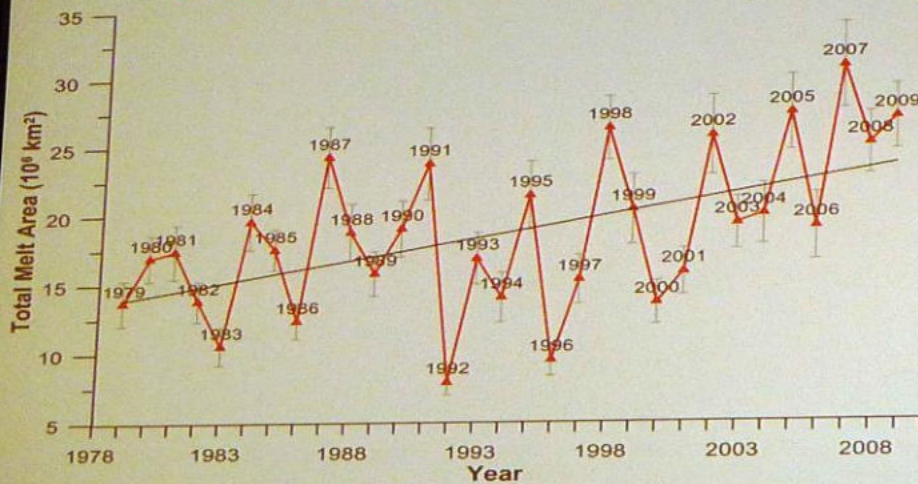
INTERGOVERNMENTAL PANEL ON climate change

Working Group I (WG I) - The Physical Science Basis

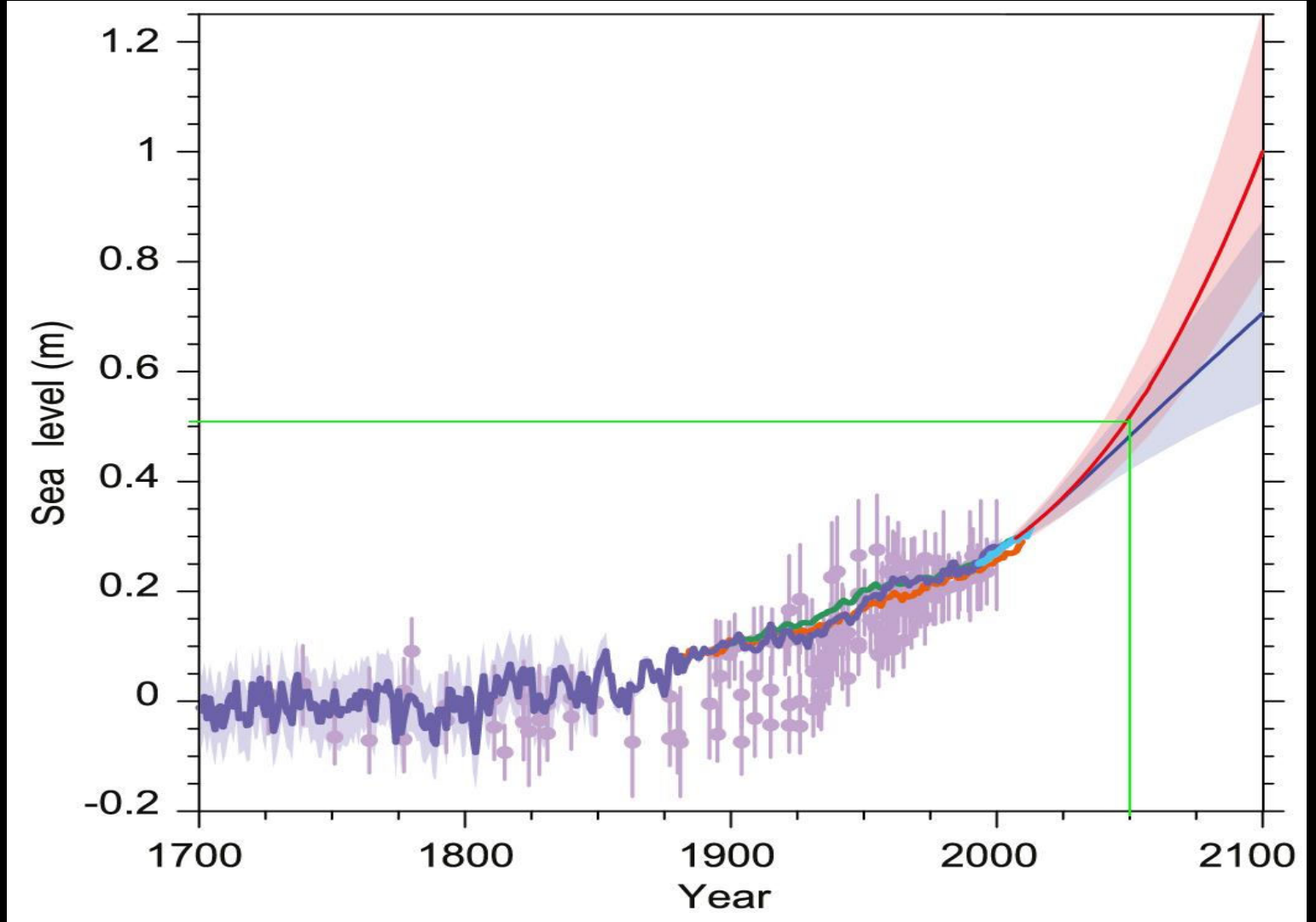
**IPCC Workshop on Sea Level Rise
and Ice Sheet Instabilities
21-24 June 2010
Kuala Lumpur, Malaysia**

Greenland Total Melt Area: 1979-2009

Total Greenland ice sheet melt area increased 65% since 1979 over the 30 year record; on average 2%/year.



The increasing trend in the total area of melting bare ice is at 13% per year



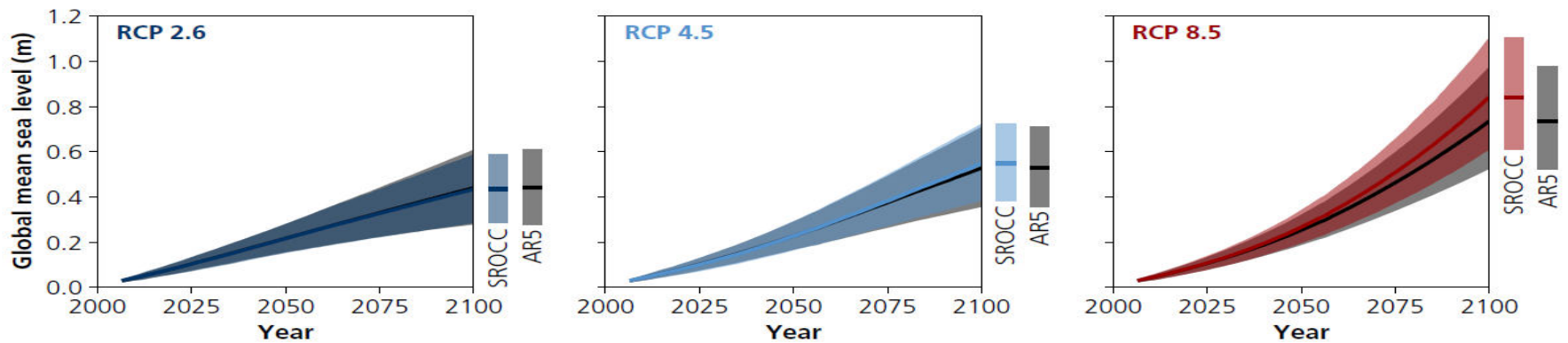


Figure 4.9: Time series of GMSL for RCP2.6, RCP4.5 and RCP8.5 as used in this report and, for reference the AR5 results (Church et al., 2013). Results are based on AR5 results for all components except the Antarctic contribution. Results for the Antarctic contribution in 2081–2100 are provided in Table 4.4. The shaded region should be considered as the *likely range*.

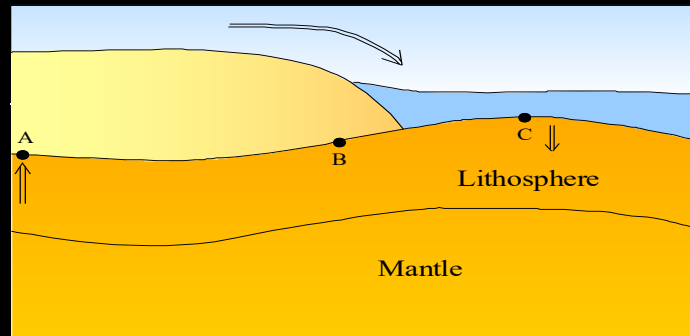
Le variazioni relative del livello del **mare**

costituiscono la sommatoria di:

Eustatismo (Scioglimento ghiacci + dilatazione termica) + **isostasia** + **tettonica**



Scioglimento dei ghiacci e dilatazione termica

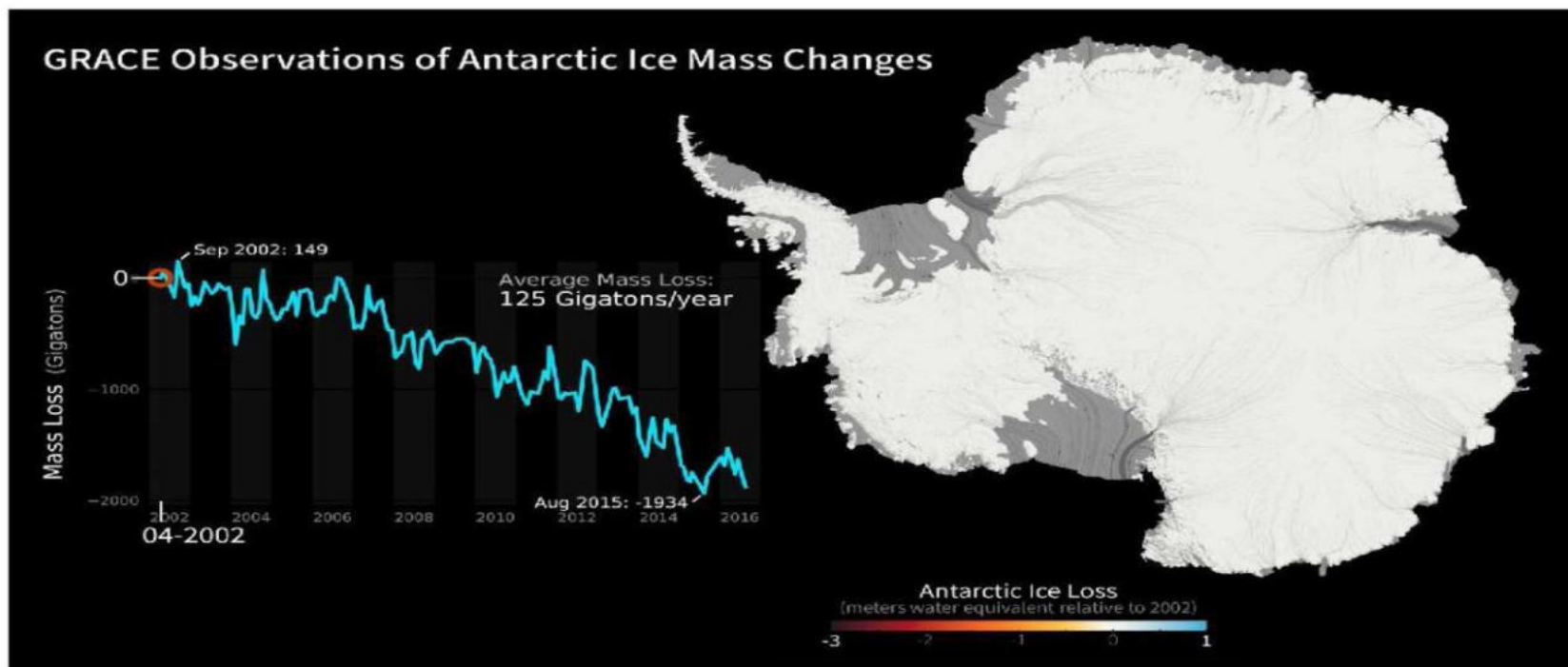


Isostasia



Tettonica e compattazione

Recent changes of the Antarctic ice sheet



Ice Thickness

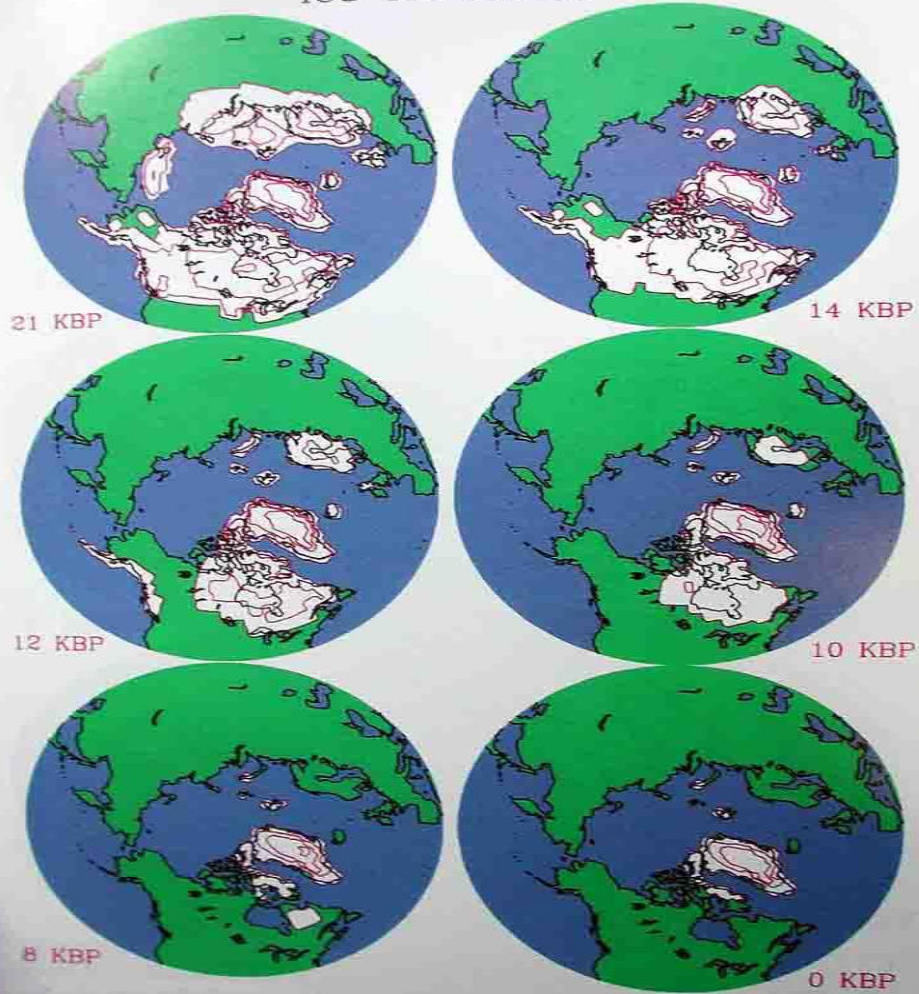
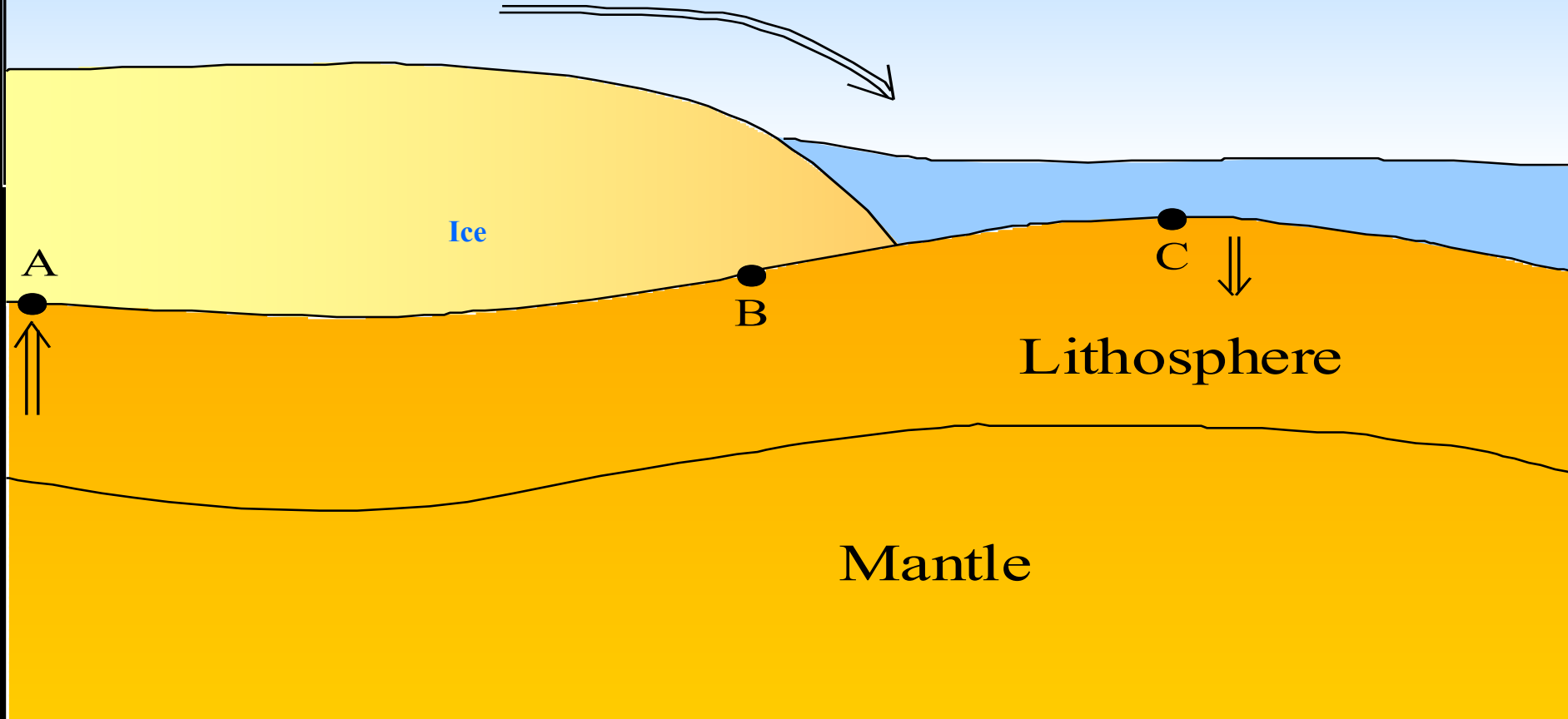


Figure 4.3

Glacio-hydro-isostatic considerations



Ice

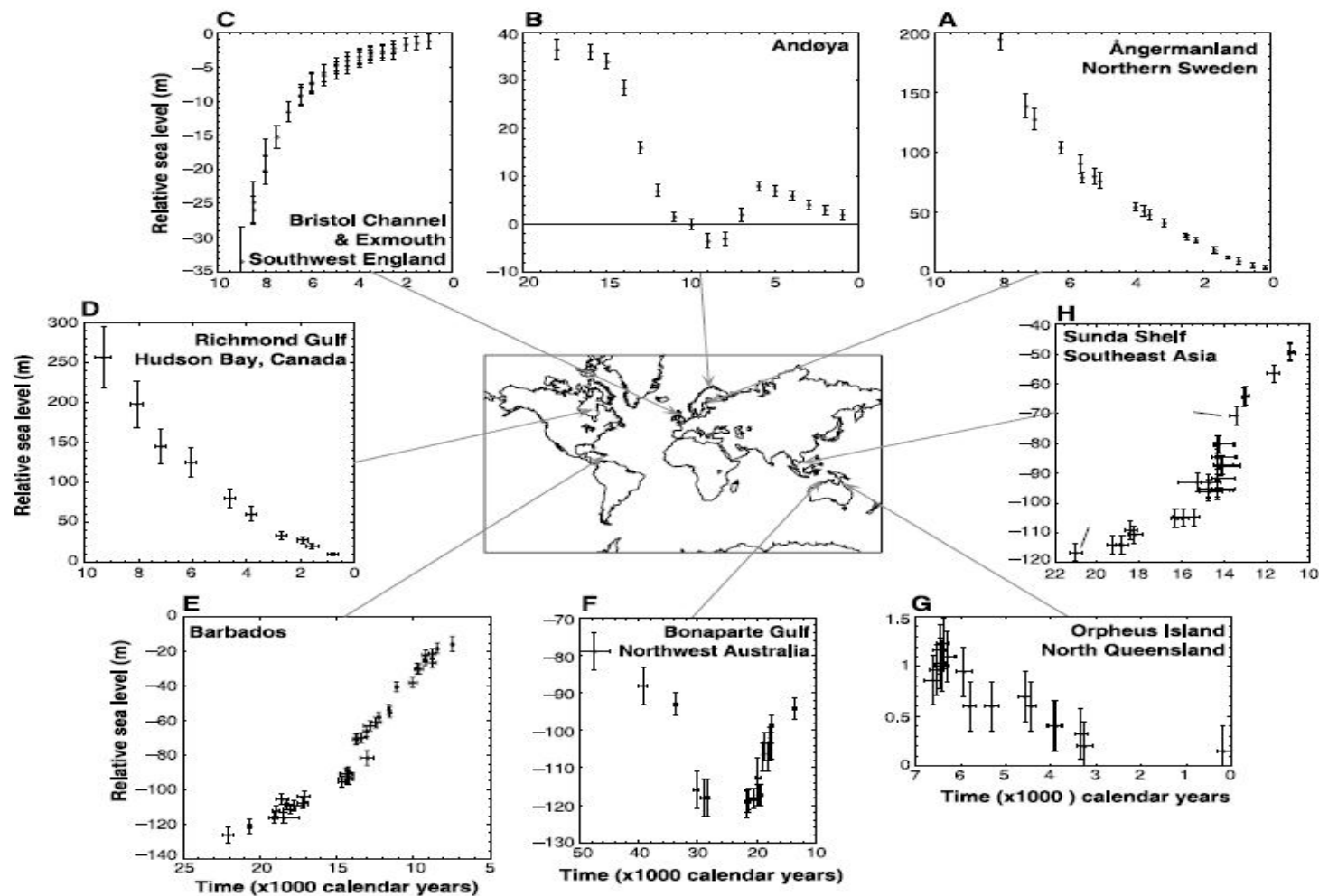
A

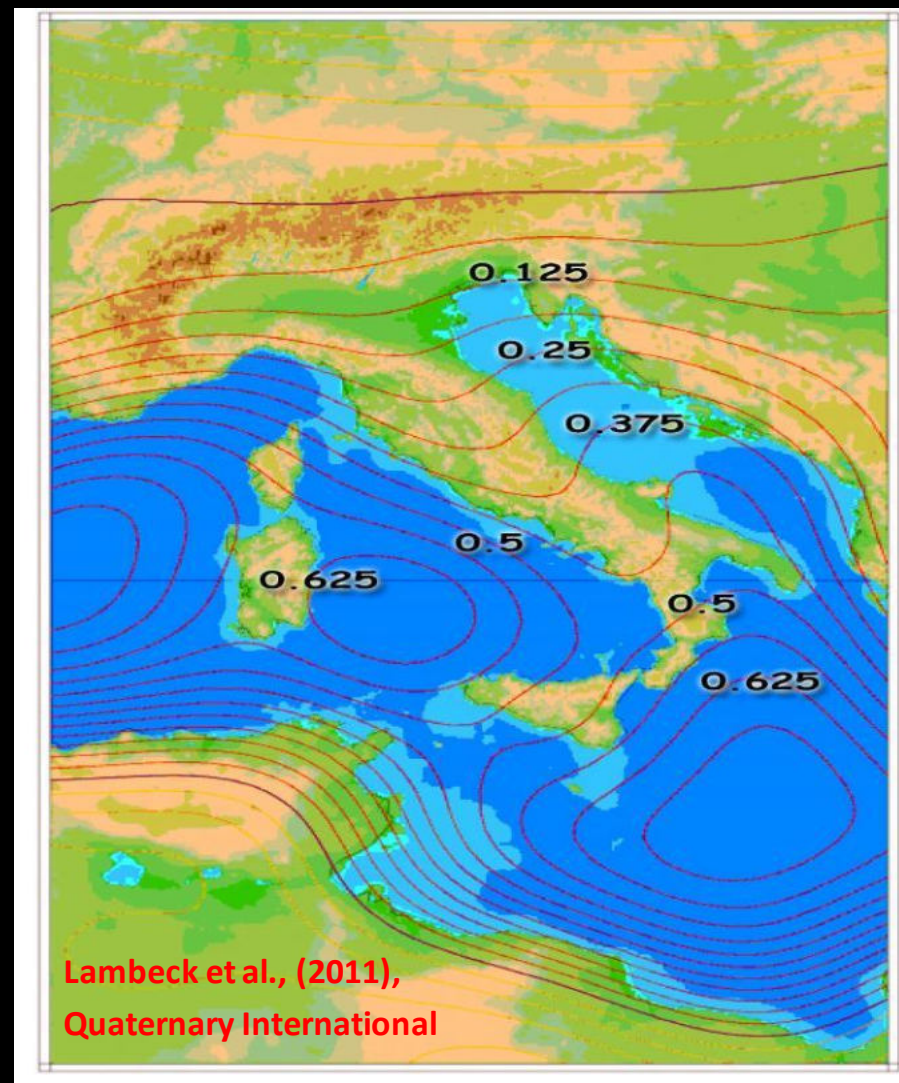
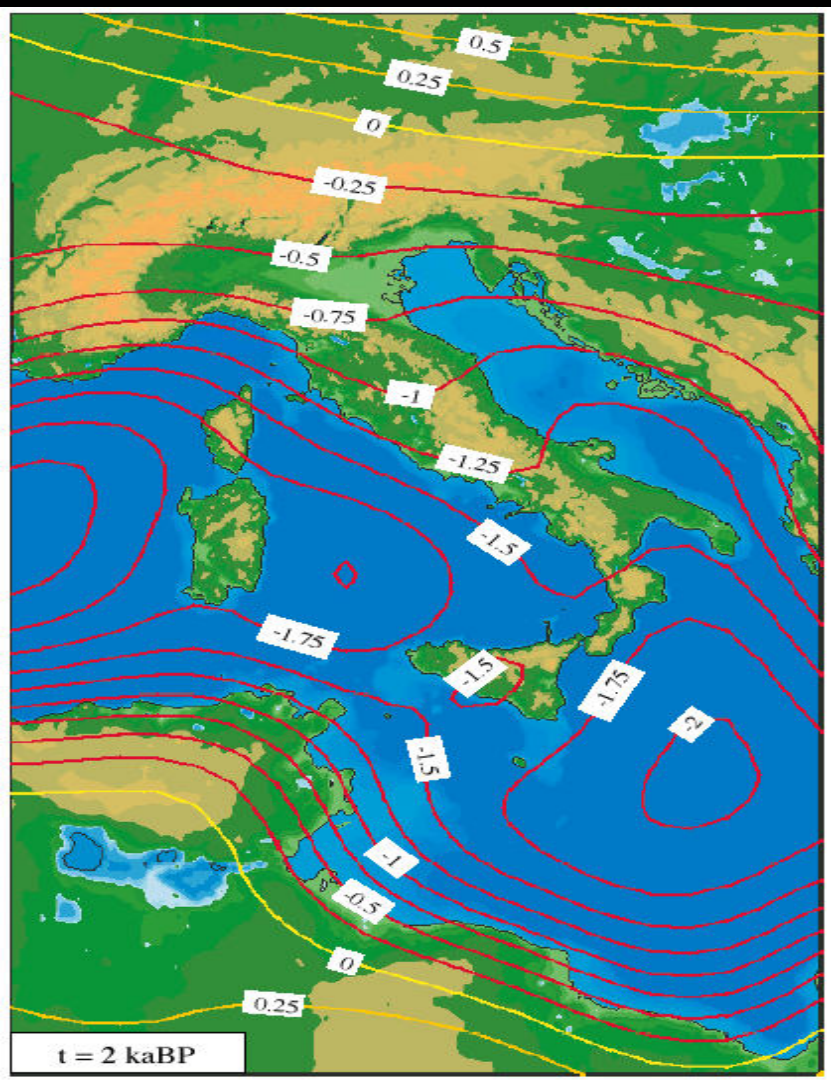
B

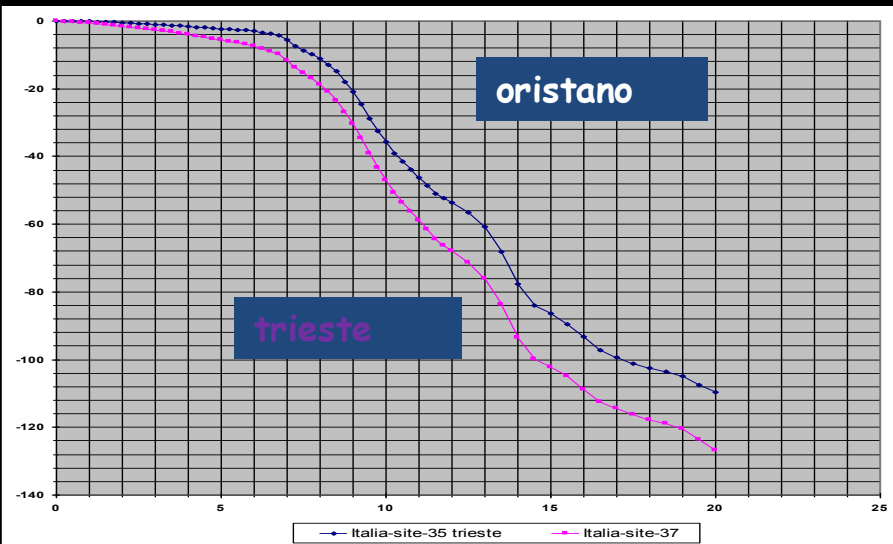
C

Lithosphere

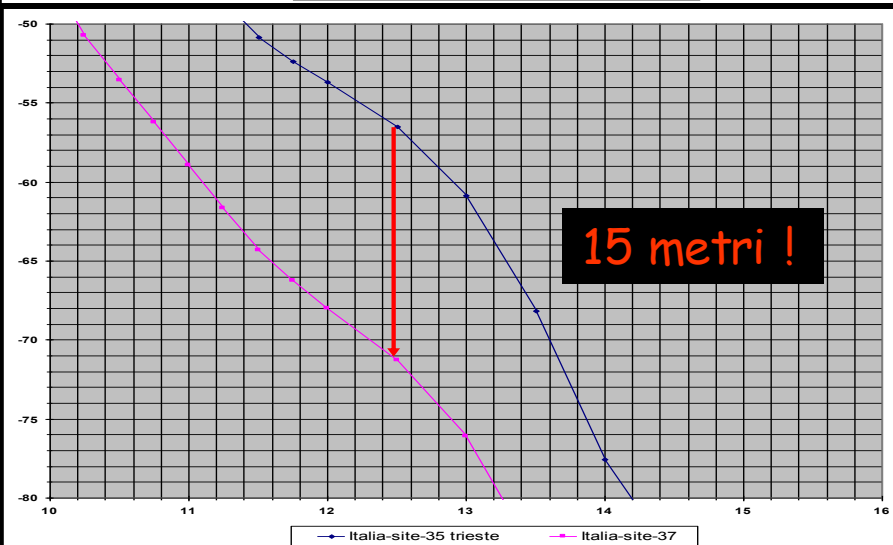
Mantle







Curve di risalita di livello del mare predette dal modello di Lambeck et al., 2011. E possibile osservare, le enormi differenze dovute solamente all'isostasia (GIA) tra l'area costiera di Trieste (minimi valori di riaggiustamento isostatico e Oristano (massimi valori di isostasia per l'Italia



Indicatori s.l.

Geomorfologici

Terrazzi solchi

Sedimentologici

Speleotemi...

Biologici

Fossili (lagunari) Vermetidi

Archeologici

Piscine, moli...

...

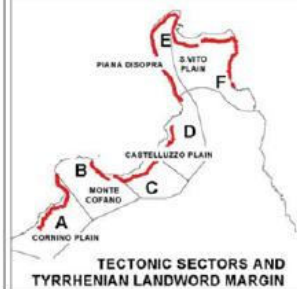
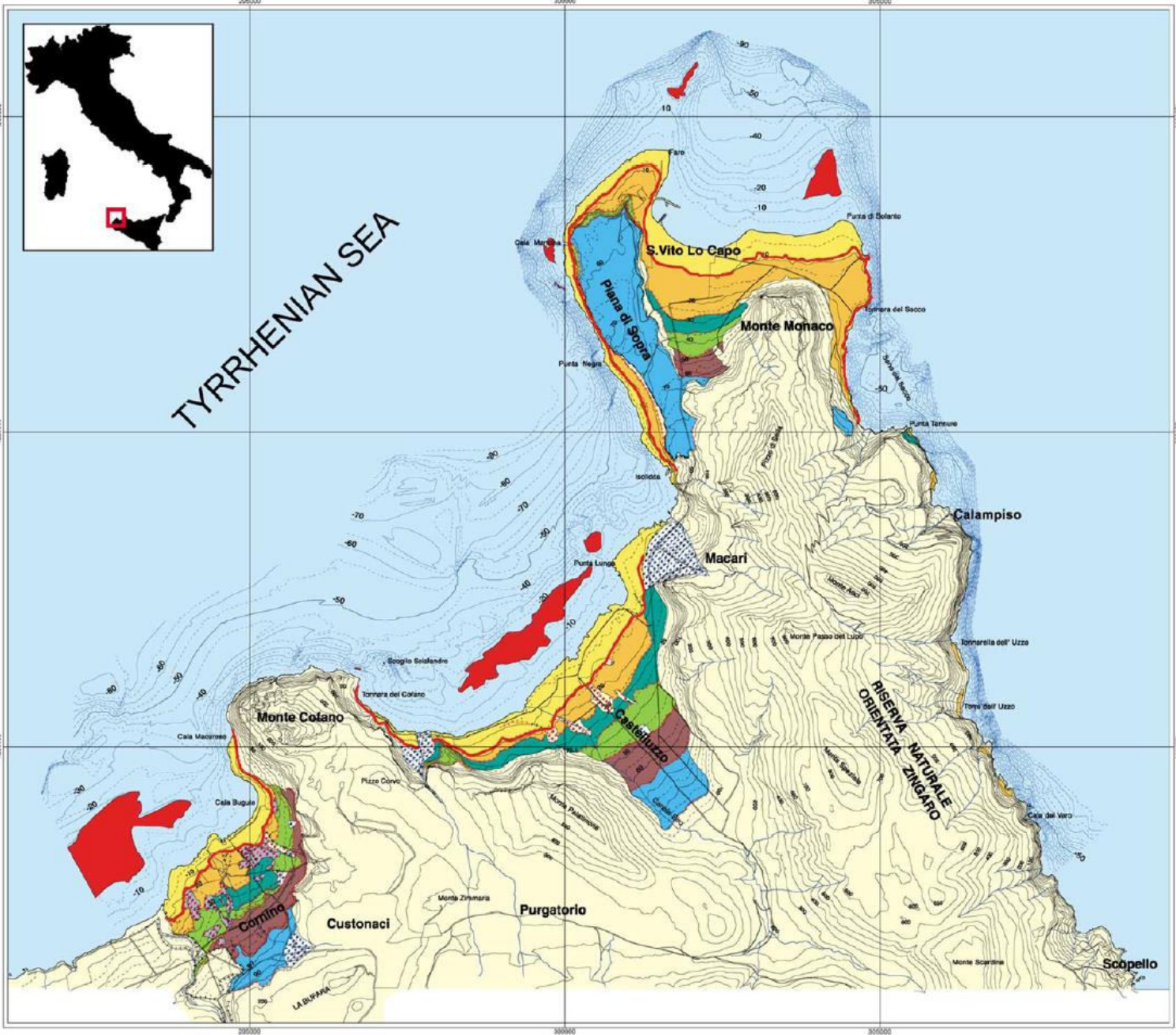
**CAPO SAN VITO PROMONTORY
(SICILY N-W - ITALY)**

COASTAL MORPHOLOGICAL MAP

SCALE 1 : 60,000

Antonelli F., Cremona G., Imbordino F., Puglisi C.,
Romagnoli C., Silenzi S., Valpreda E., Verrubbi V.

TYRRHENIAN SEA



**TECTONIC SECTORS AND
TYRRHENIAN LANDWARD MARGIN**

LEGEND

MARINE TERRACES

- VII order
- VI order (Eutyrrhenian)
- V order
- IV order
- III order
- II order
- I order

LAND-WARD MARGINS

- Eutyrrhenian
- Others

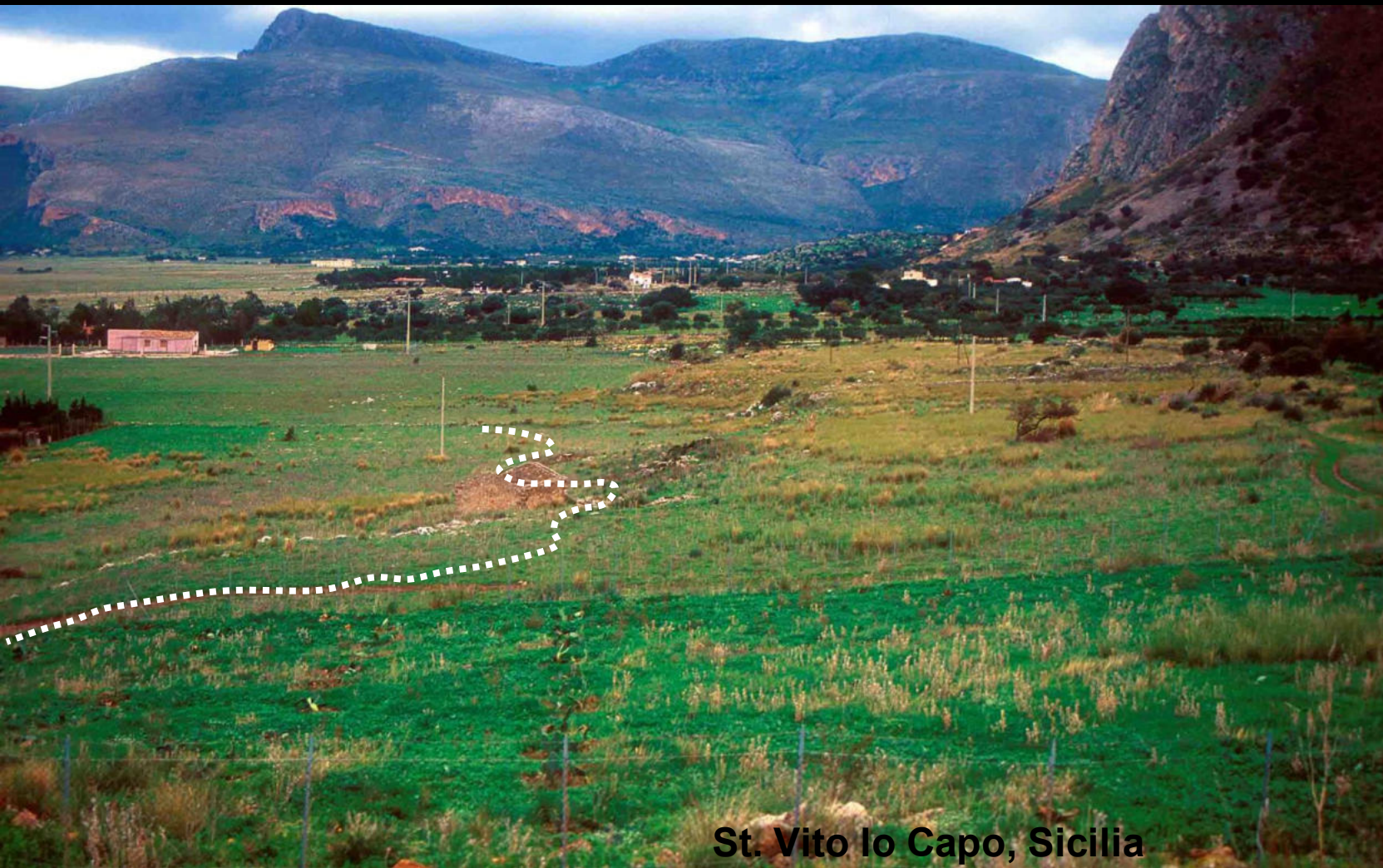
OTHER MORPHOLOGIES AND DEPOSITS

- Alluvial fan
- Fluvial bed surface
- Coastal Dunes and Eolianites
- Fluvial slope

FIG. 2

The coordinates are referred to European Datum (E.D. 55) and the projection is in UTM system.

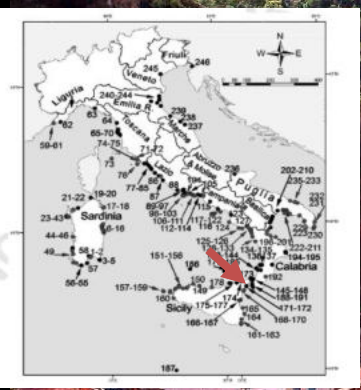
**Margine interno di un terrazzo marino risalente a circa
125.000 anni fa**



St. Vito lo Capo, Sicilia



MIS 5.5 inner margin 127 m



Article

MIS 5.5 highstand, and future sea level flooding at 2100 and 2300 in tectonically stable areas of central Mediterranean sea: Sardinia and the Pontina Plain (southern Latium), Italy

Giacomo Deiana^{1,2*}, Fabrizio Antonioli³, Lorenzo Moretti⁴, Paolo E. Orrù^{1,2,3}, Giovanni Randazzo⁵, Valeria Lo Presti⁶



Figure 13. a, Google earth image of the Pontina and Fondi Plain with the carbonatic promontory of Terracina, Sperlonga on which are carved the fossil tidal notches (b and c) aged MIS 5.5 the yellow arrow indicate samples number and altitude (see Table 3) Pia-nura Pontina e di Fondi. d, the coastal area of northern Pontina Plain (on background he limestone Circeo Promontory); e, the coastal area of northern Pontina Plain (on background he limestone Terracina promontory. f, the coastal area of Fondi Plain (on background the limestone Terracina Promontory; g same point but a southern view, with Sperlonga promontory.



Figure 14. a: *Cerastoderma edulis* sampled in a section outcropping in the channel of this figure in e. b *Tapes decussatus* from the outcrop on the reclamation drainage c channel; c and e the channel d the Mussolini channel during the excavation in 30s; f the outcrop *Nassa mutabilis* *Tapes decussatus*, *Cerastoderma edulis*; f a fossiliferous level very rich of lagoon fauna. See also Table 3 site 7.1.

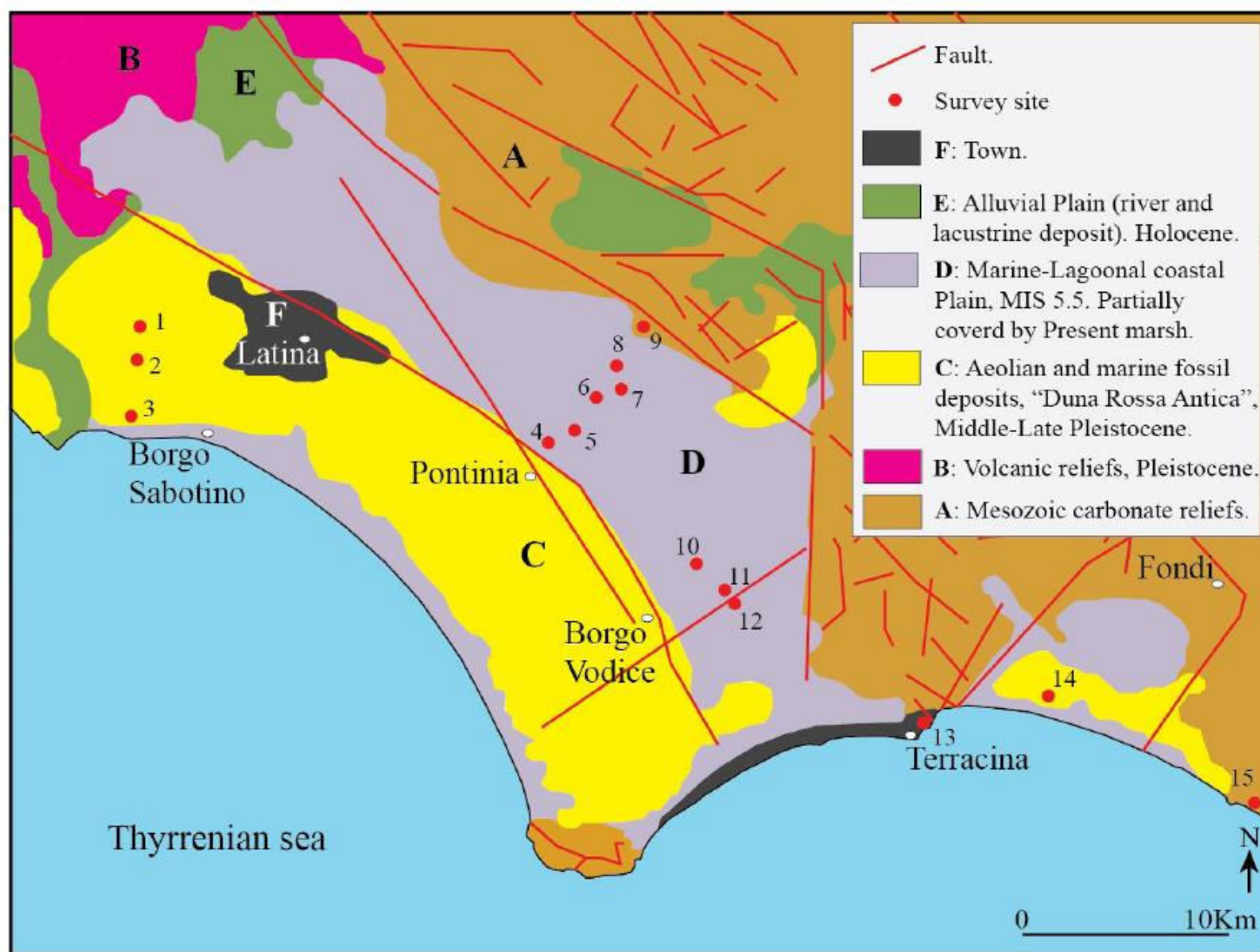


Figure 12. - Main geological outcroppings of the Pontina Plain, this map is a compilation of: i) Italian Geological Survey sheet numbers 170, 158 and 159; ii) Map of the soil [51], iii) sinkhole map of Regione Lazio http://www.regione.lazio.it/binary/rl_main/tbl_documenti/AMB_PBL_Carta_Sinkholes_Lazio_2011.pdf. The red dots refer to the sites described in Table 3.

2	Canale Mussolini	41.4483 12.8107	5.1 ± 0.1	Fossil beach containing <i>Persistrombus latus</i>	Senegalese Fauna Aminoacid	[37,40,43]
3	Nuclear power plant Borgo Sabotino	41.4231 12.8053	-4.3 +5 ± 0.5	Fossil beach containing <i>Persistrombus latus</i>	Senegalese Fauna	[43]
4	Pontinia 1	41.4129 13.0449	+5.3 ± 0.5	Lagoonal facies with <i>Cerastoderma s.p.</i>	Geomorphological correlation	[17,46]
5	Pontinia 2	41.4172 13.0600	+4.4 ± 0.5	Lagoonal facies with <i>Cerastoderma</i>	Geomorphological correlation	[17,46]
6	Pontinia 3	41.4323 13.0721	+2.3 ± 0.5	Lagoonal facies with <i>Cerastoderma s.p.</i>	Geomorphological correlation	[17,46]
7	Pontinia 4	41.4355 13.0864	+0.8 ± 0.5	Lagoonal facies with <i>Cerastoderma s.p.</i>	Geomorphological correlation	[17,46]
7.1	Check in field	41.434771 13.062667	-1± 0.5	Cerastoderma e Tapes, travertino con incrostazioni	Geomorphological correlation	This paper
7.2	Check in field	41.3757510 13.1281060	-2± 0.5	Lagoonal facies with <i>Cerastoderma s.p.</i>	Geomorphological correlation	This paper
7.3	Check in field	41.363875 13.140631	-3± 0.5	Lagoonal facies with <i>Cerastoderma edulis</i> , <i>Tapes decussatus</i> , <i>Nassa</i> <i>mutabilis</i>	Geomorphological correlation	This paper
8	Pontinia 5	41.4424 13.0751	-0.5 ± 0.5	Lagoonal facies with <i>Cerastoderma</i>	Geomorphological correlation	[17,46]
9	Mezzaluna core	41 27 47 13 06 01	-14.30 -11.41 ± 0.5	Venus and <i>Cerastoderma</i>	Pollen Analysis, U\Th and aminoacid	[45]
10	Borgo Vodige 1	41.3571 13.1317	1 ± 0.5	Lagoonal facies with <i>Cerastoderma</i>	Aminoacid	[17,46]
11	Borgo Vodige 2	41.3497 13.1293	-0.6 ± 0.5	Lagoonal facies with <i>Cerastoderma</i>	Aminoacid	[17,46]
12	Borgo Vodige 3	41.350 13.117	-1.80 ± 0.5	Lagoonal facies with <i>Cerastoderma</i>	Aminoacid	[17,46]
13	Terracina	41.288 13.260	7.96 ± 0.1	Tidal notch	Geomorphological correlation at 5 km from aged MIS 5.5 deposit	[7]
14	Fondi APT4	41.0065 13.331	-6 \ -24	Marsh with <i>Cerastoderma</i> Aminozone E	Aminoacid	[44]







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Tidal notches in Mediterranean Sea: a comprehensive analysis



Fabrizio Antonioli ^a, Valeria Lo Presti ^{b, a, *}, Alessio Rovere ^{c, d}, Luigi Ferranti ^e, Marco Anzidei ^f, Stefano Furlani ^g, Giuseppe Mastronuzzi ^h, Paolo E. Orru ⁱ, Giovanni Scicchitano ^j, Gianmaria Sannino ^a, Cecilia R. Spampinato ^k, Rossella Pagliarulo ^l, Giacomo Deiana ^l, Eleonora de Sabata ^m, Paolo Sansò ⁿ, Matteo Vacchi ^o, Antonio Vecchio ^f

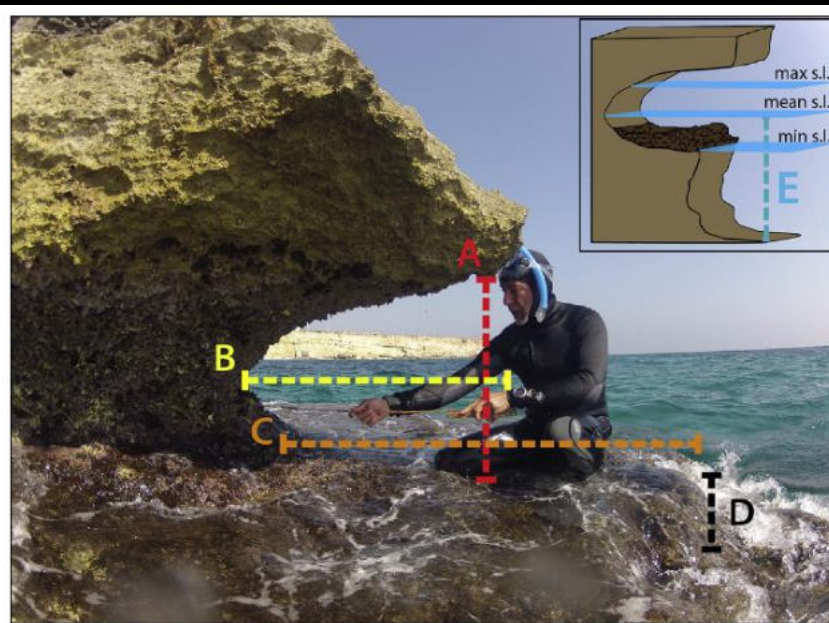
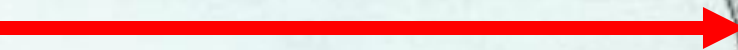


Fig. 3. Morphometric measures: A) Average notch width, B) Notch depth, C) Bottom depth (reef when present), D) reef and step (if present) thickness, E) Depth of cliff toe at mean sea level.

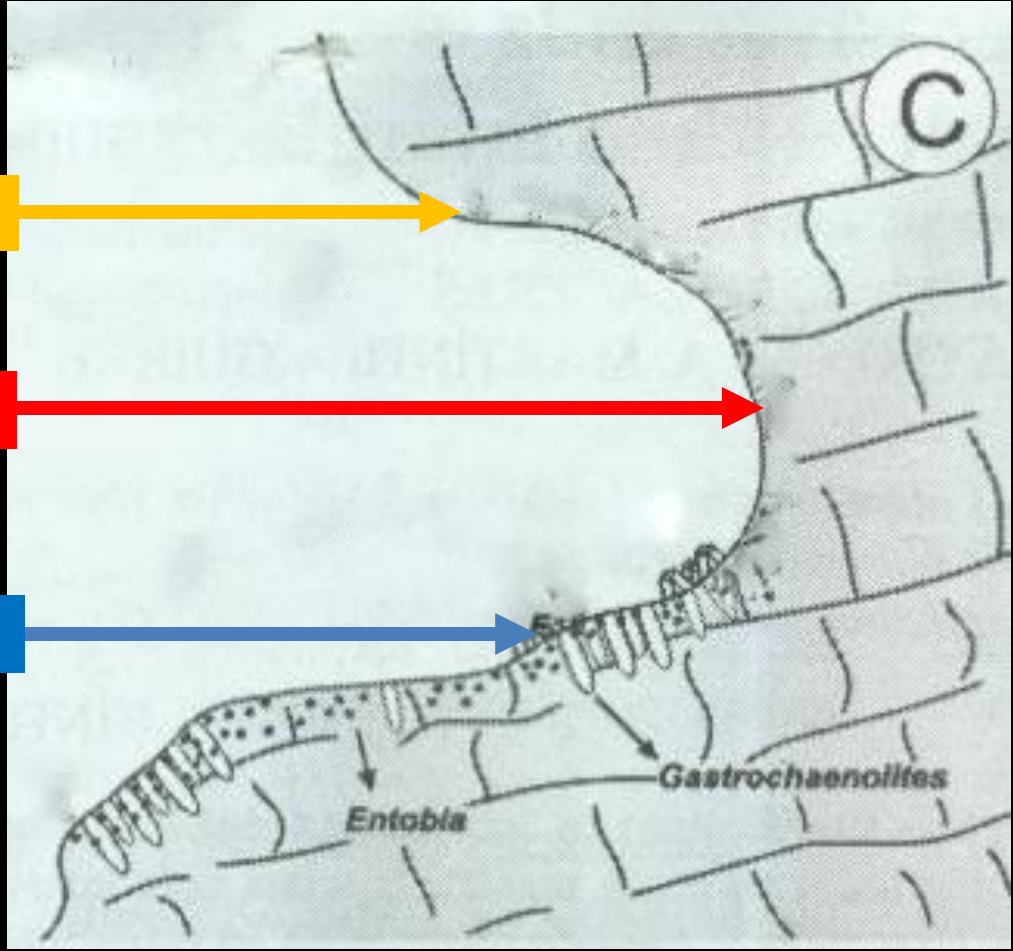
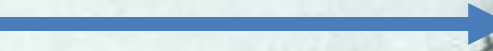
high tide



Mean sl



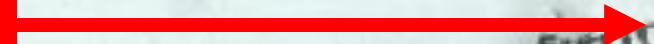
Low tide



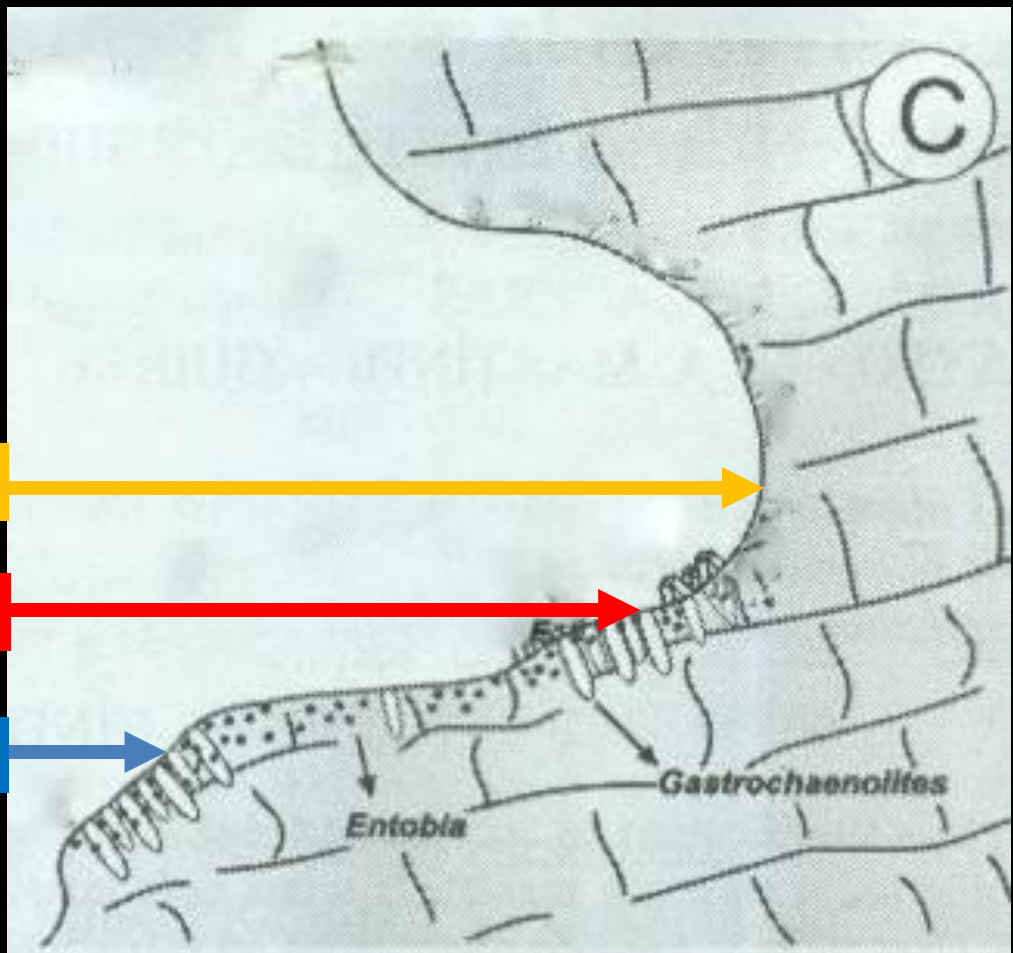
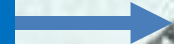
high tide



Mean sl



Low tide



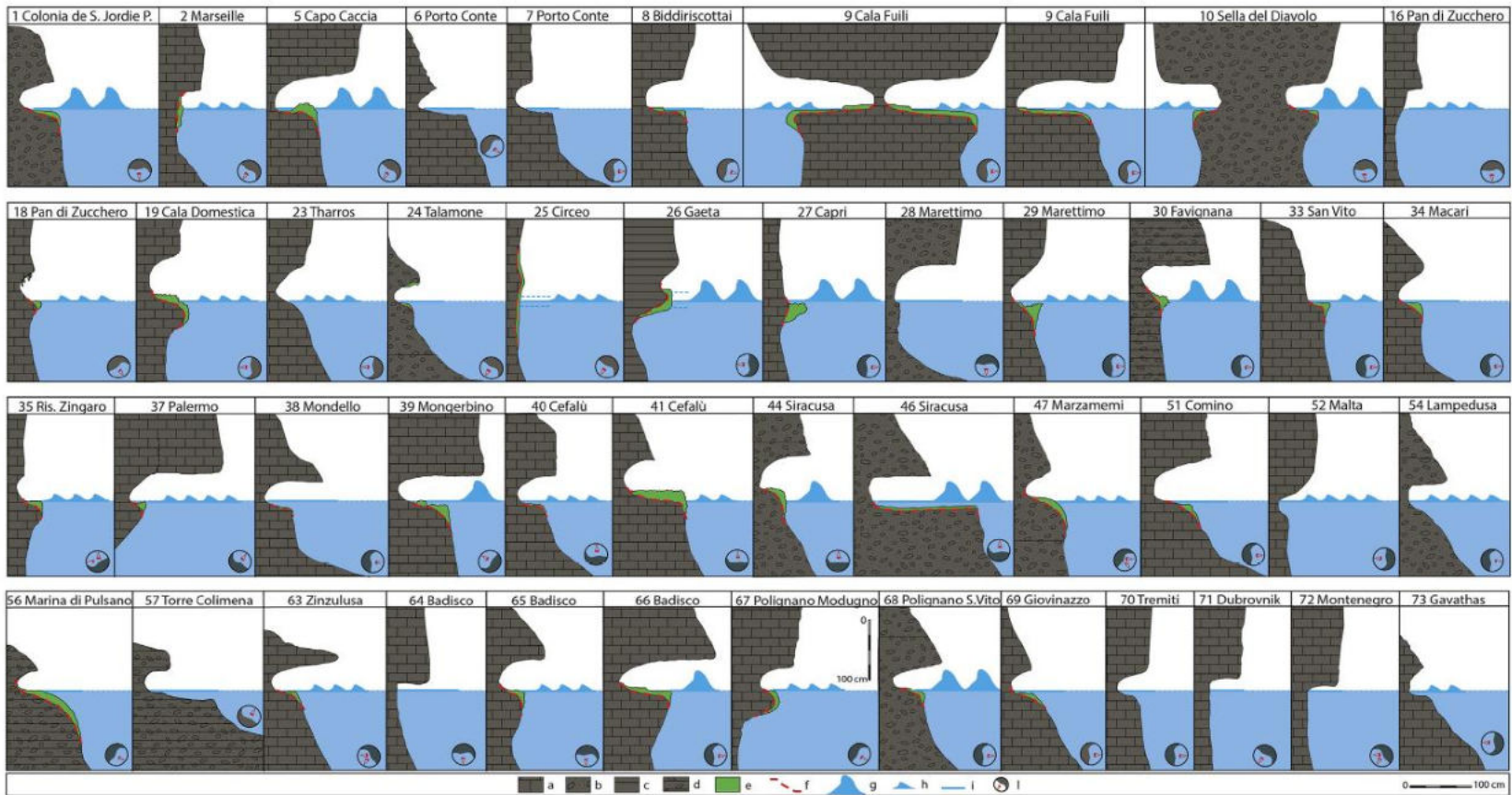


Fig. 6. Representative sections of most significant tidal notches studied. a) limestone, b) sandstone and very erosive limestone, c) stratified limestone, d) stratified sandstone and very erosive limestone, e) reef, f) supposed limit between rock and reef. Fetch and kind of sea energy: g) very exposed, h) exposed, i) sheltered. l) geographical exposure (see Table 2).

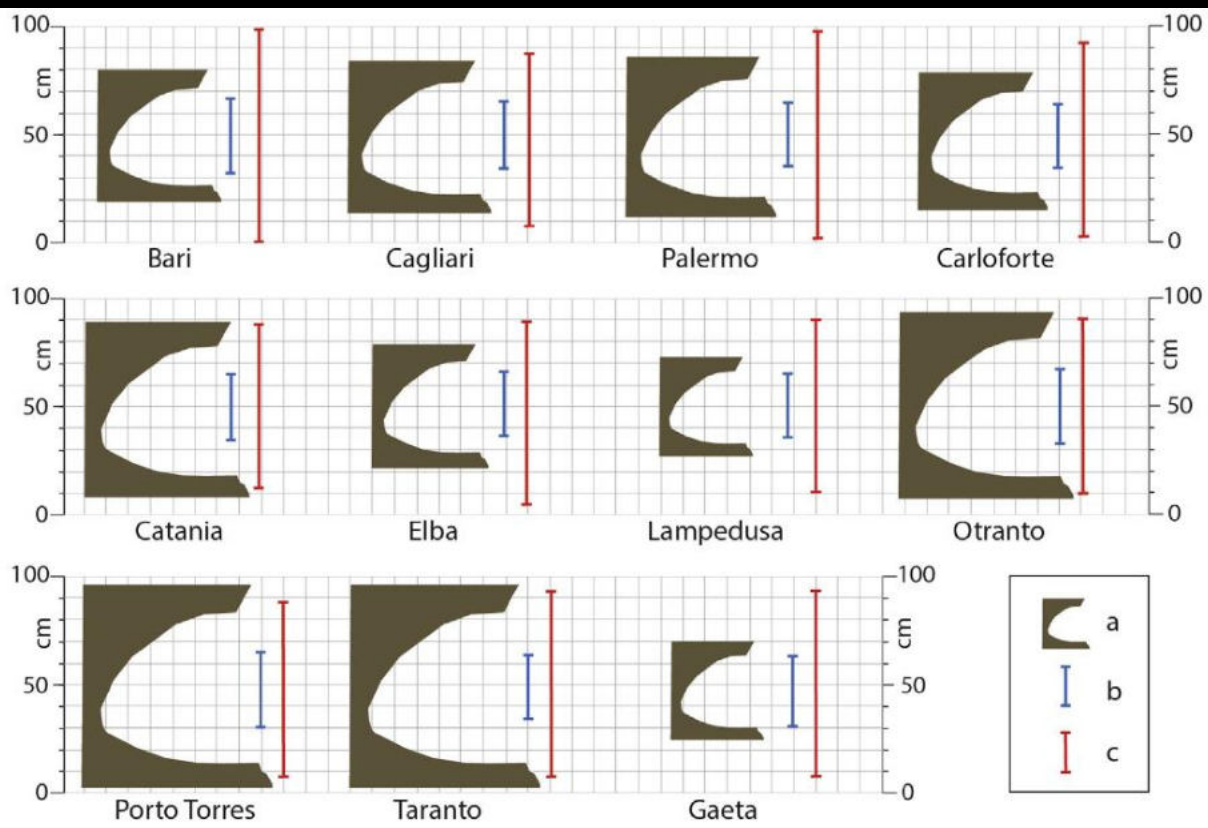


Fig. 7. Relationship between notch width (a), mean tide values (b) and extreme (max-min) tide values (c) in locations where notches have been measured near a tide station.

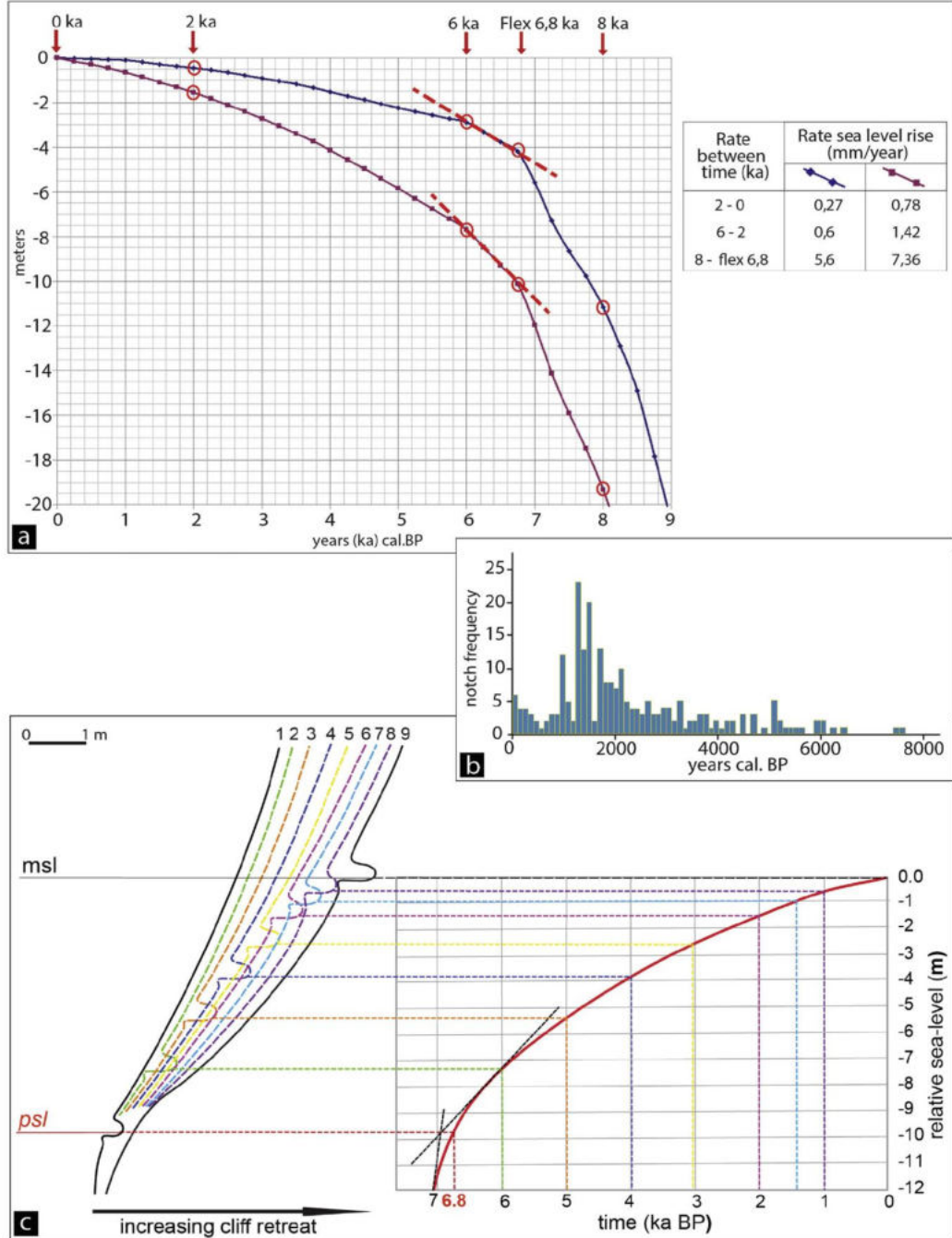


Fig. 10. **a)** Sea level rise rates from 8 ka cal BP to the Present using the predicted sea level rise curves (Lambeck model, 2011) with maximum (Cagliari, purple line) and minimum (Trieste, blue line) isostatic subsidence values. The inflection at 6.8 ka marks the change of rise. **b)** Frequency of notches formation from 8 ka to 0 ka (modified from Boulton and Stewart, 2015). **c)** Model of tidal notches formation from 6.8 ka to the present. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)

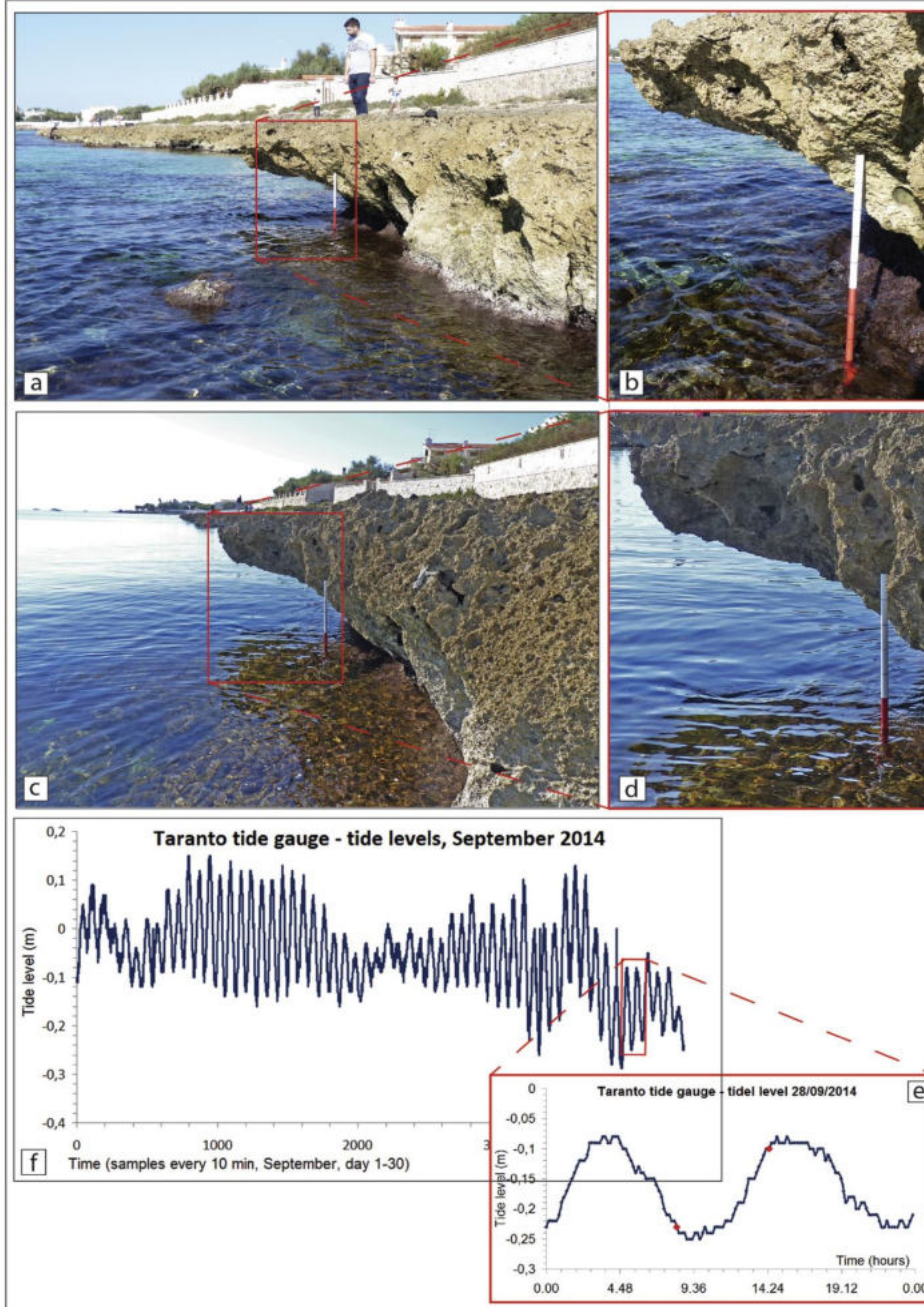


Fig. 12. Tidal range measured on the Taranto – San Vito notch (site 68 Table 1 and 51. Measured on 09/28/2014 at 10.00 am (a,b), and 16.30 pm (c,d). Observations are in agreement with the instrumental data collected at the nearest tide gauge located at Taranto. Plots show the daily tide during the observations (e) while in (f) are the tides for one month cycle of September. The red arrow indicate the time of the notch measures (a,b,c,d). (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)









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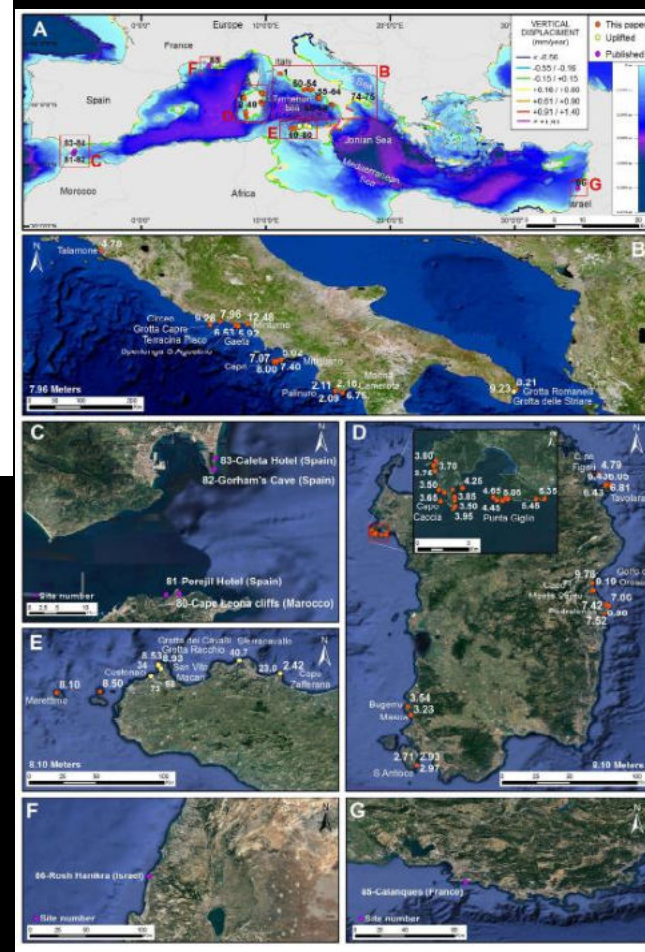
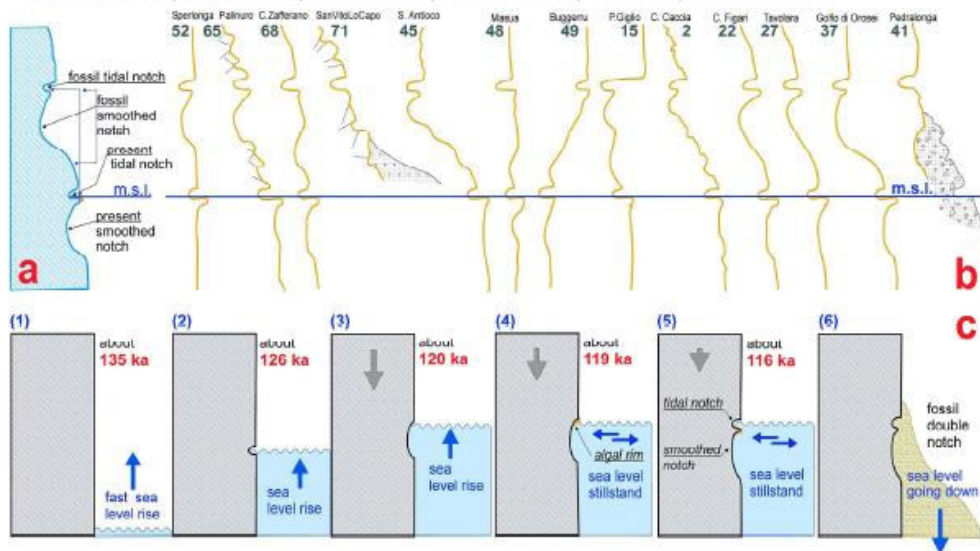
Earth-Science Reviews

journal homepage: www.elsevier.com/locate/earscirev



Morphometry and elevation of the last interglacial tidal notches in tectonically stable coasts of the Mediterranean Sea

Antonoli F.^a, Ferranti L.^b, Stocchi P.^c, Deiana G.^d, Lo Presti V.^a, Furlani S.^e, Marino C.^b, Orru P.^d, Scicchitano G.^f, Trainito E.^g, Anzidei M.^h, Bonamini M.ⁱ, Sansò P.^j, Mastronuzzi G.^k



A. F. et al.

Earth-Science Reviews 185 (2018) 600–623

Fig. 2. a) MIS 5.5 and PTN notches morphological sketch. b) Tidal and smoothed notch as is possible to observe today in some sections we studied. c) Evolution of the tidal and smoothed notch during MIS 5.5, the final coverage of fans or aeolianites deposits preserves the notches from the dissolution.

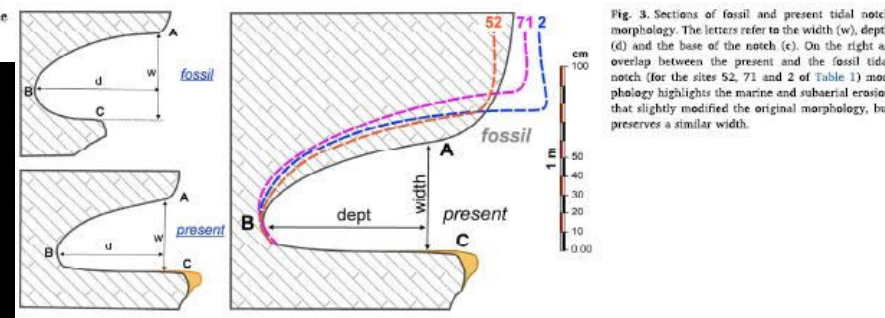


Fig. 3. Sections of fossil and present tidal notch morphology. The letters refer to the width (w), depth (d) and the base of the notch (c). On the right an overlap between the present and the fossil tidal notch (for the sites 52, 71 and 2 of Table 1) morphology highlights the marine and subaerial erosion that slightly modified the original morphology, but preserves a similar width.

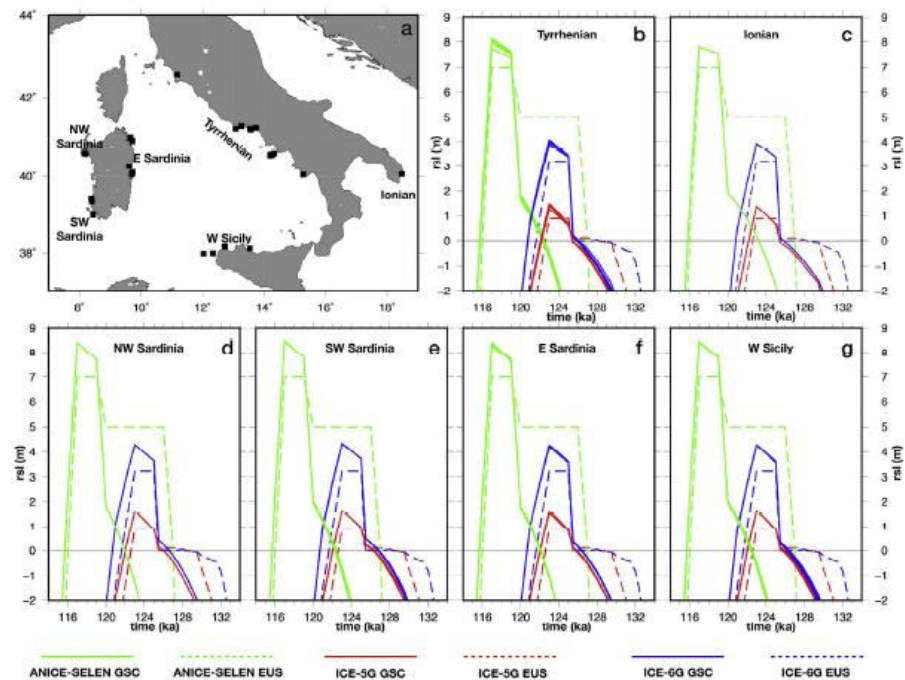
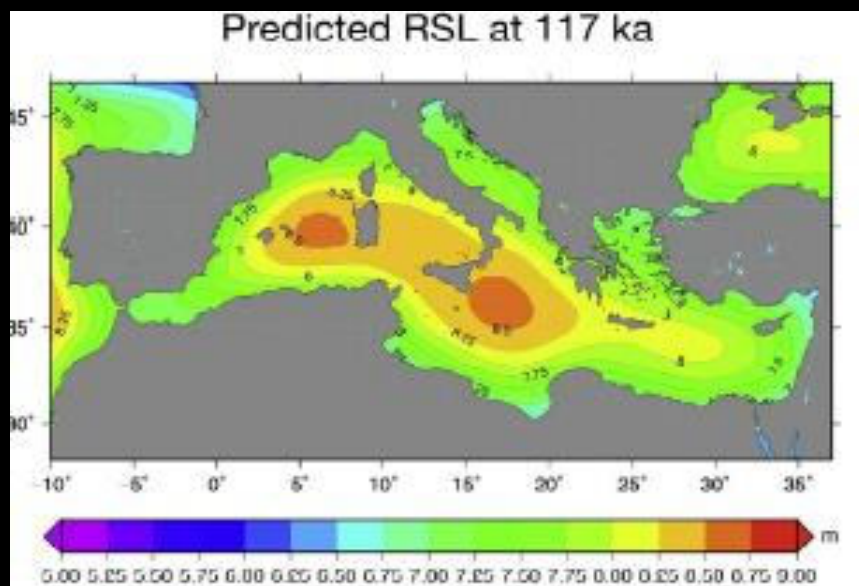


Fig. 6. Predicted MIS 5.5 RSL curves at sites along the Italian coastlines according to ICE-5G (red curves), ICE-6G (blue curves) and ANICE-SELEN (green curves) ice-sheet models. The dashed curves represent the eustatic trend, while the solid curves represent the GIA-induced RSL changes. The RSL curves are computed at each investigated site and are plotted cumulatively for different sub-regions. (For interpretation of the references to colour in this figure legend, the reader is referred to the web version of this article.)





Fori di *Litophaga* fossili a Marettimo, si osservano fino a 7.30 metri. Ad 8 metri è presente un solco di battente fossile







San Vito lo Capo, Sicily, scogliere a Vermetidi (*Dendropoma petreum*), un gasteropode coloniale che vive a livello del mare, e risulta utilissimo quando trovato fossile. Qui sotto la sezione di un frammento di reef, i Vermetidi oggi fossili, campionati e datati nella parte inferiore, hanno fornito una data di circa 600 anni a -30 centimetri, Antonioli et al., 1991, *Marine Geology*. Alcune note aree della Sicilia vengono naturalmente ripascite per la distruzione da parte di tempeste, di questi gasteropodi coloniali che forniscono preziosissimo materiale sabbioso sulla costa.





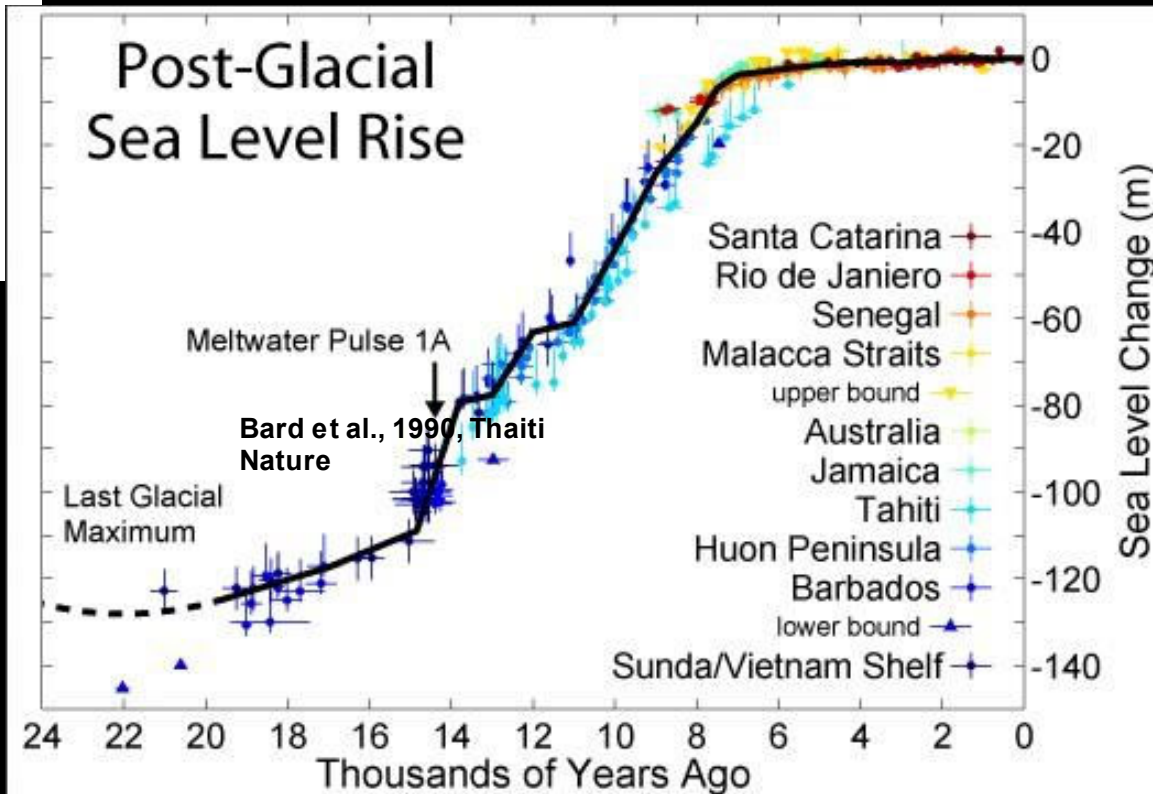
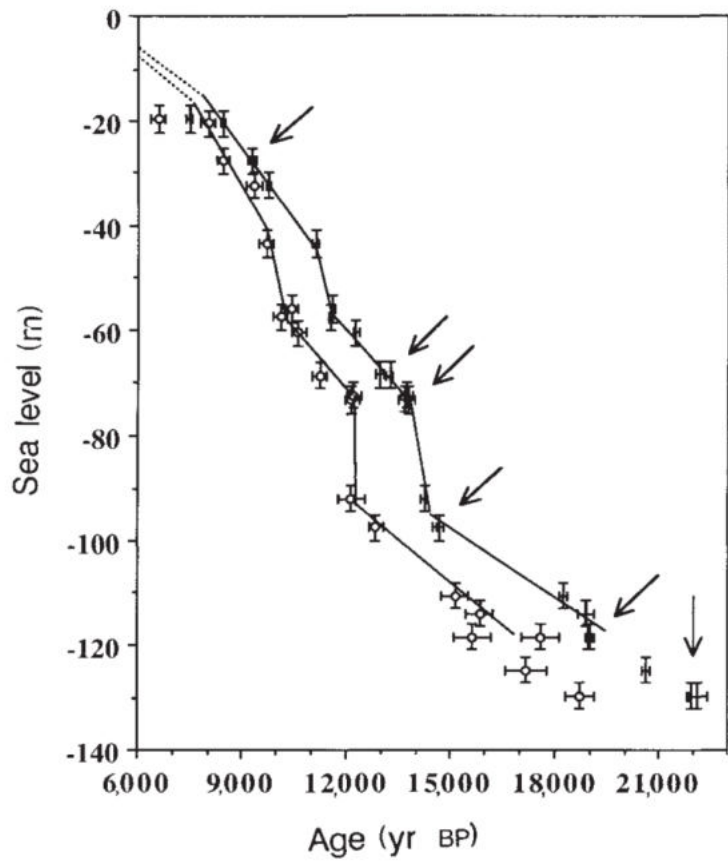
Data SIO, NOAA, U.S. Navy, NGA, GEBCO

Google earth

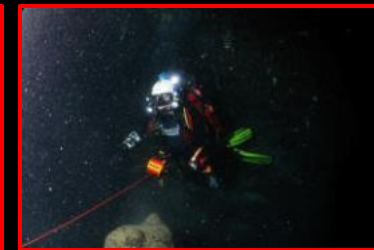








Thanks to: Giorgio Caramanna,



Marco Oliverio,



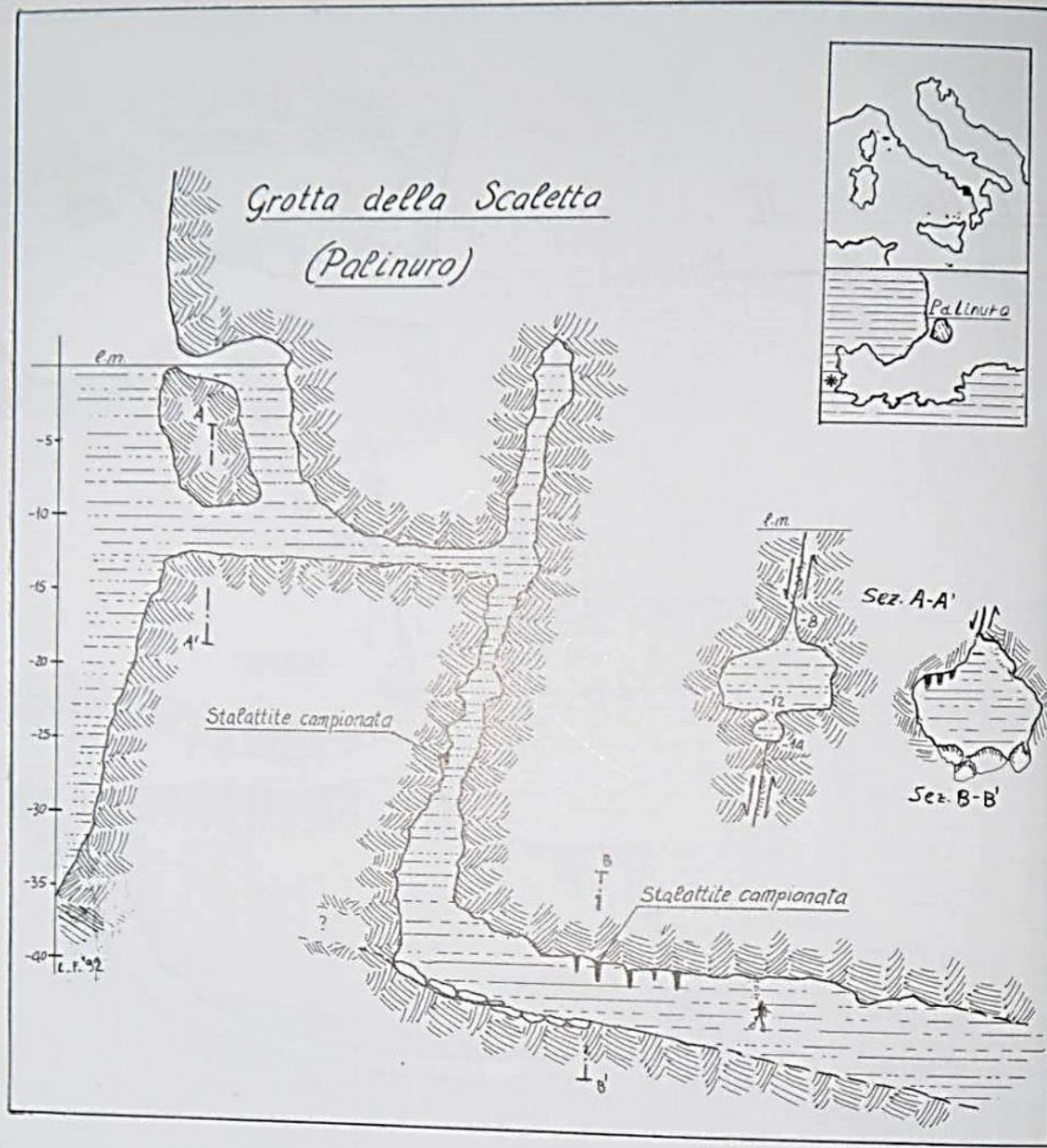
Marco with the speleo sampled at -48 m



Gianfra Scicchitano



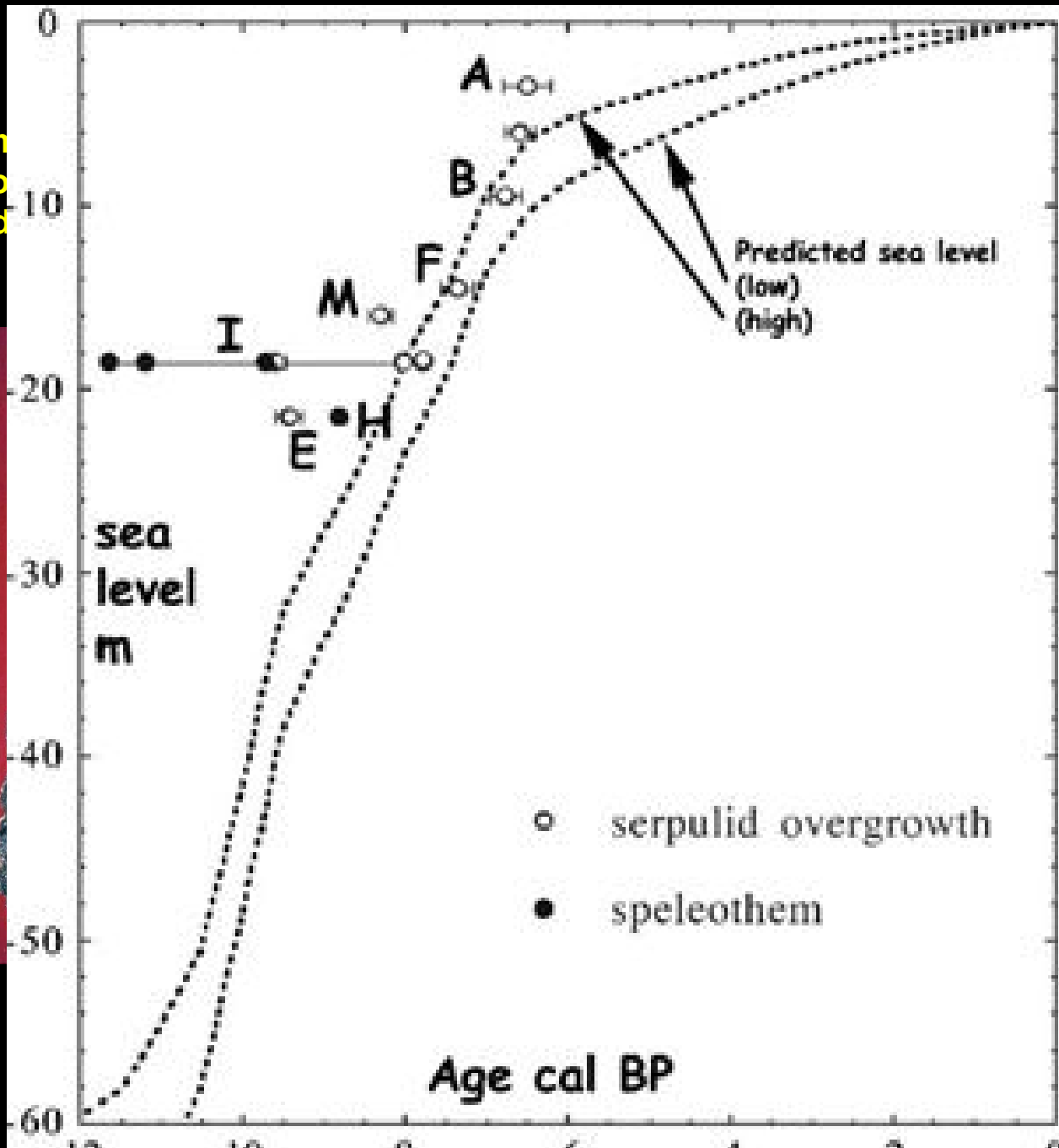




ne longitudinale della Grotta della Scaletta, Capo Palinuro.
itudinal section of Grotta della Scaletta at Cape Palinuro.

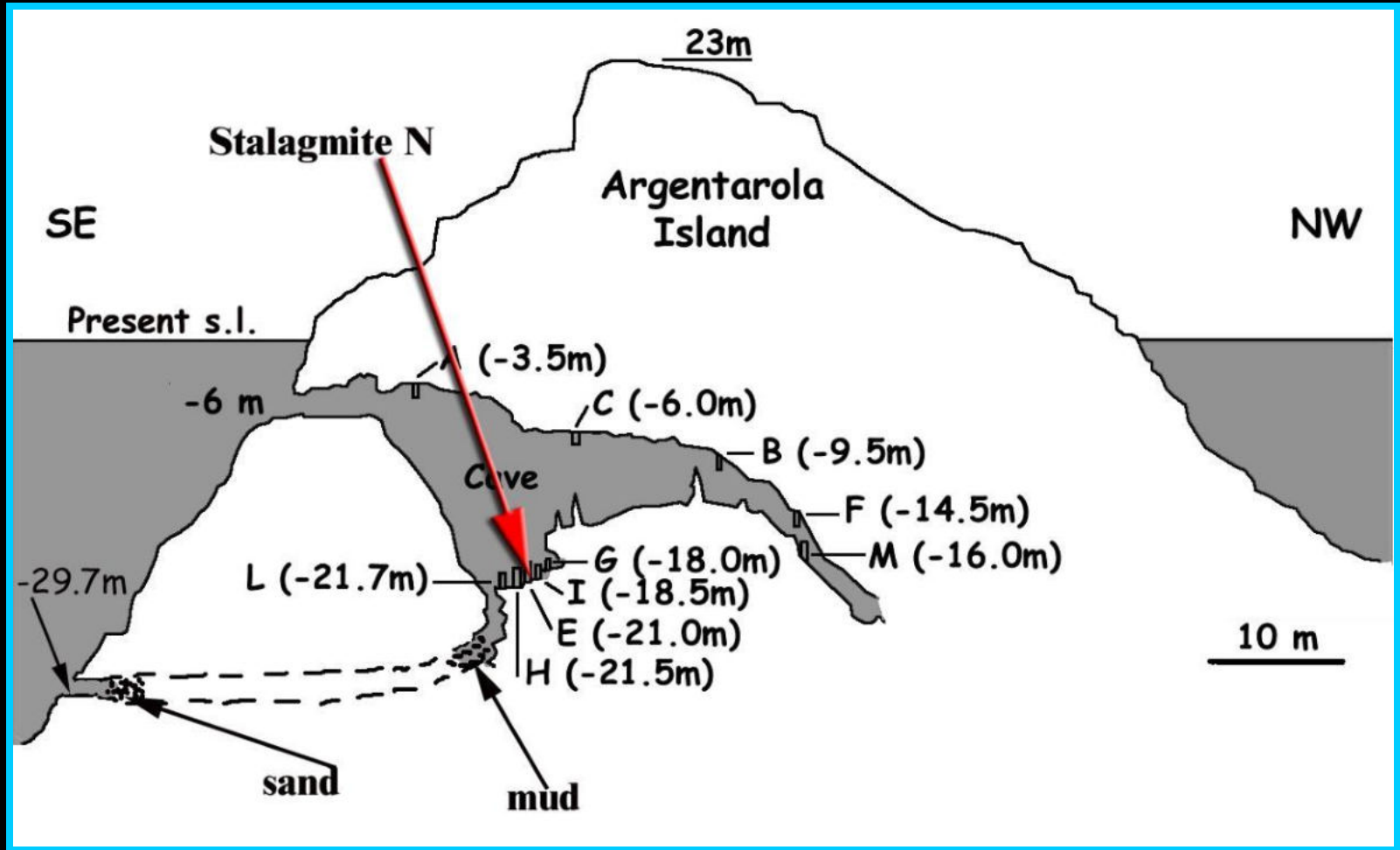


La stalattite cava
e forata da Lithothamnion
la data precisa di



ta da Serpuldi
ga, ha fornito
rio 1989

Cross Section of Argentarola Cave





MIS 5

145.2±1.1

•
•
164.9±1.9

•
168.7±1.4

MIS 6

•
171.9±1.3

•
179.2±1.3

•
189.7±1.5

MIS 7.1

•
201.6±1.8

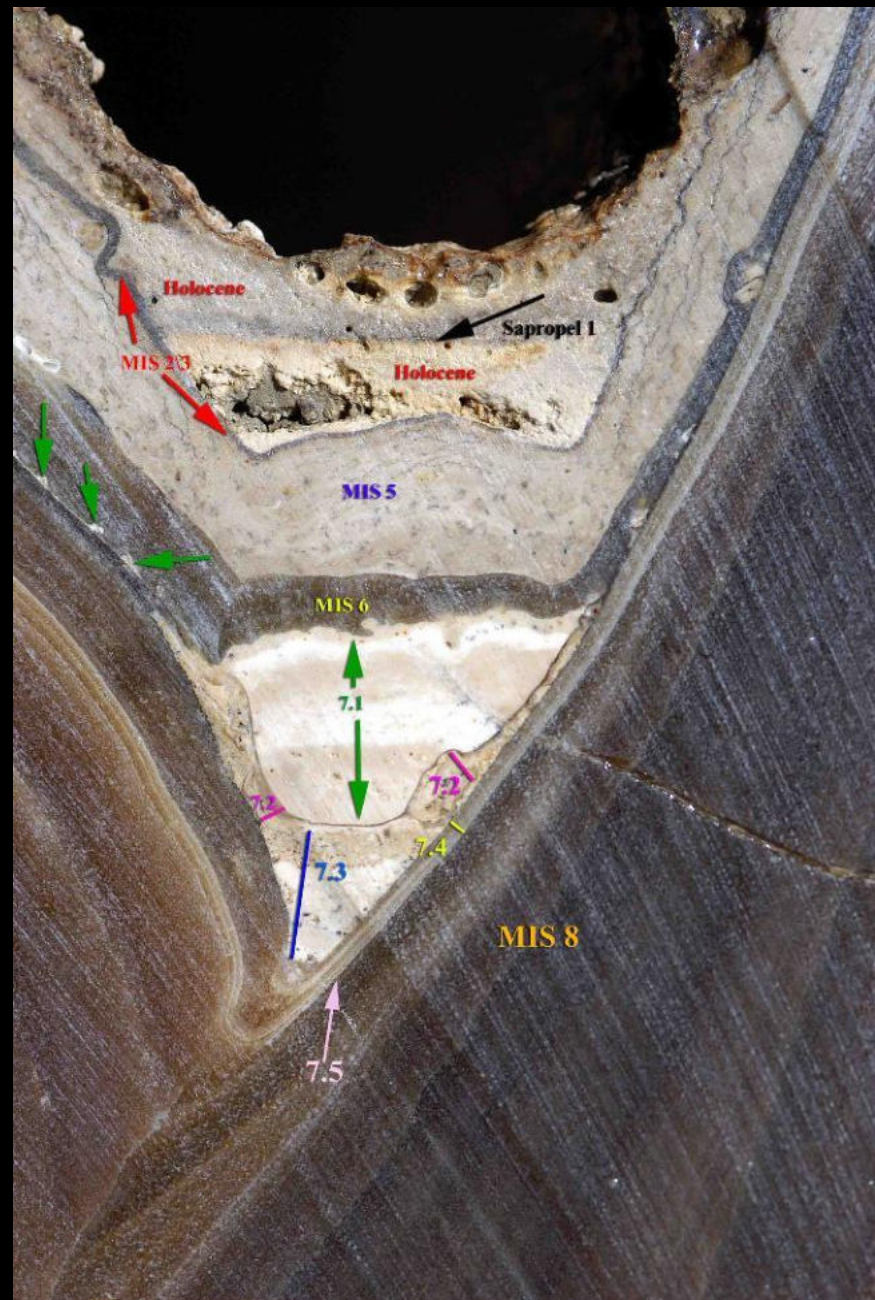
•
201.5±1.7

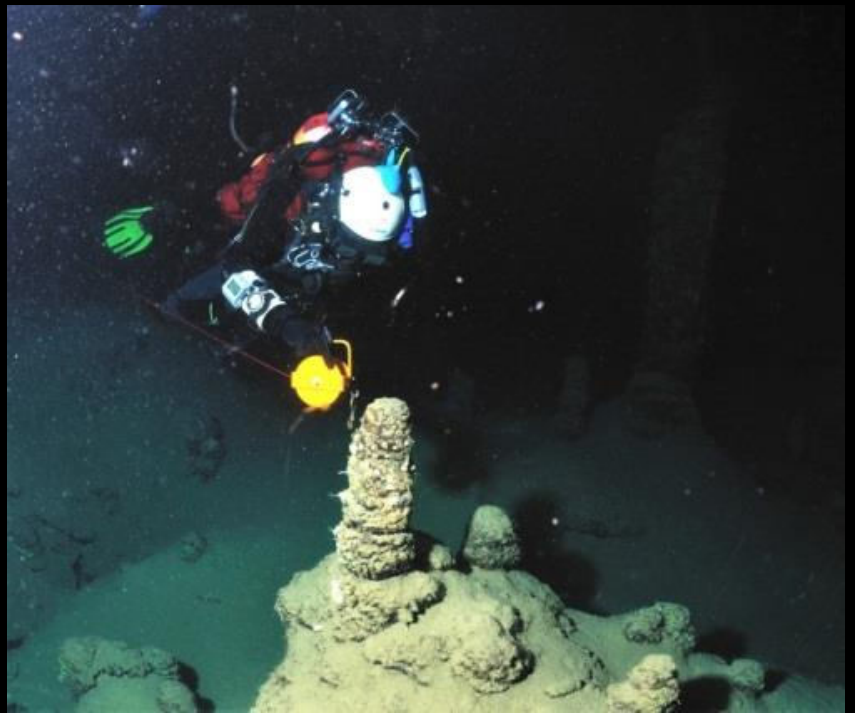
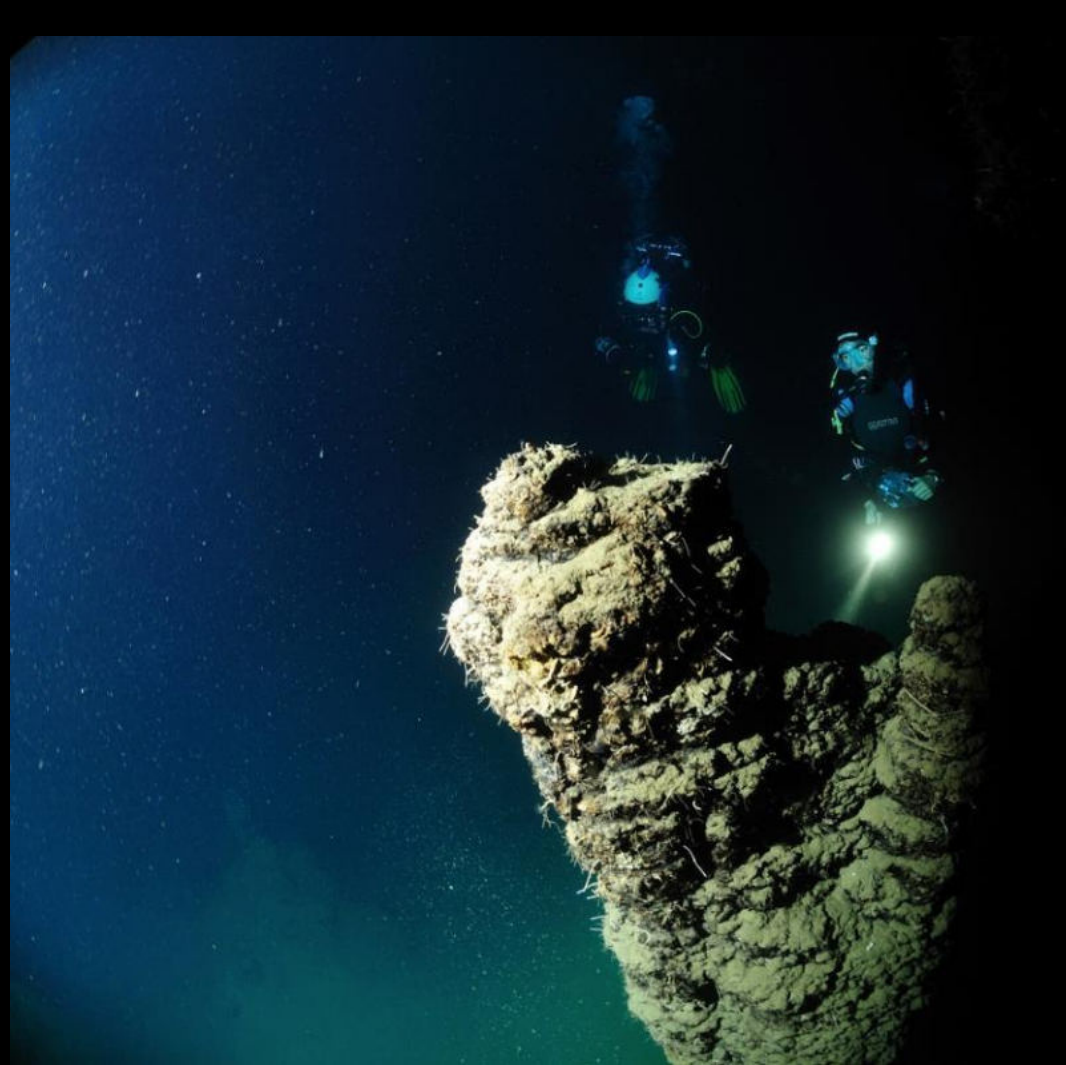
MIS 7.2

•
206.0±1.9

5 cm







Phasing and amplitude of sea-level and climate change during the penultimate interglacial

Andrea Dutton^{1*}, Edouard Bard², Fabrizio Antonioli³, Tezer M. Esat^{1†}, Kurt Lambeck¹ and Malcolm T. McCulloch¹

1 Earth's present climate has evolved through oscillations between short-lived interglacial and extended glacial periods for the past million years. Direct markers of sea-level variations within these cycles that are absolutely dated are rare for periods older than the last interglacial; hence, our knowledge of sea-level change driven by the waxing and waning of continental ice sheets before that time is largely based on proxy records from deep-sea cores¹⁻³. Here we present precise U-Th ages from a collection of submerged speleothems^{4,5} from Italy, which record three sea-level highstands during the penultimate interglacial period, Marine Isotope Stage 7, from 245,000 to 190,000 years ago. We find that in the first and third highstands maximum sea levels of about -18 m (relative to modern sea level) were reached several thousand years before maximum northern hemisphere insolation, whereas in the second, Marine Isotope Stage 7.3, the highstand is essentially synchronous with the insolation maximum. Sea level during Stage 7.3 also peaked at about -18 m, even though the concurrent insolation forcing was the strongest of the three highstands. We attribute the different phasing and amplitude of this highstand to the extensive continental glaciation that preceded Marine Isotope Stage 7.3, and conclude that the response time of the cryosphere is an important component of the climate system.

Reconstructions of previous interglacial periods can help elucidate the phasing and causal mechanisms linking temperature, sea level and greenhouse gas concentrations that are critical to our understanding of climate dynamics during the present interglacial and into the future. Although the penultimate interglacial offers an important test for theories about the timing, duration, magnitude and driving mechanisms of sea-level highstands that have evolved out of studies of the last interglacial and the Holocene, fewer data are available for this period of time. Hence, the additional data presented herein provide an important test for several competing reconstructions of sea level during Marine Isotope Stage (MIS) 7 (refs 2,3,6-8) (Fig. 1).

A common approach to reconstruct past sea level is to measure the age and elevation of geologic archives that formed at a known position relative to the sea surface, such as corals⁶ that grow in shallow water, or speleothems^{9,10} that grow in caves and serve as an upper limit to sea level elevation. The strength of this method lies in the degree of confidence afforded by an accurate and precise chronology, such as provided by the U-Th dating technique, and in the ability to precisely measure the elevation relative to present sea level. We have applied this technique to submerged speleothems

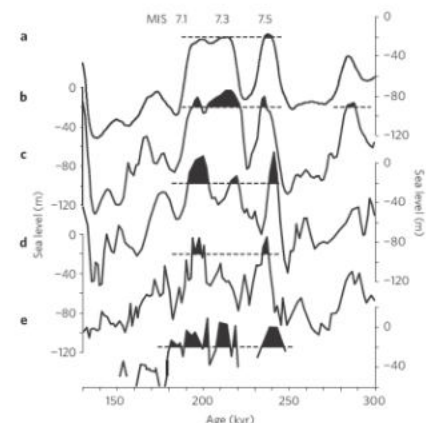
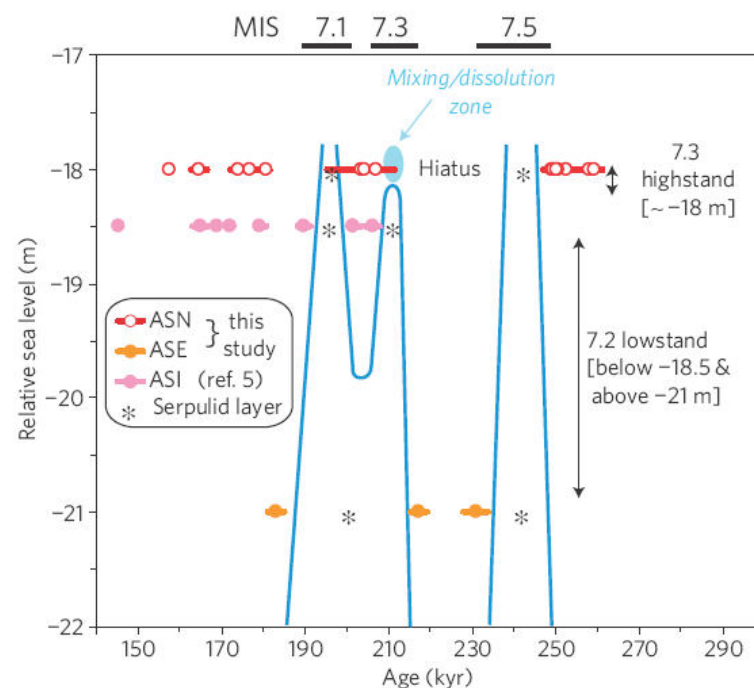


Figure 1 | Sea-level curves derived using five different methods.

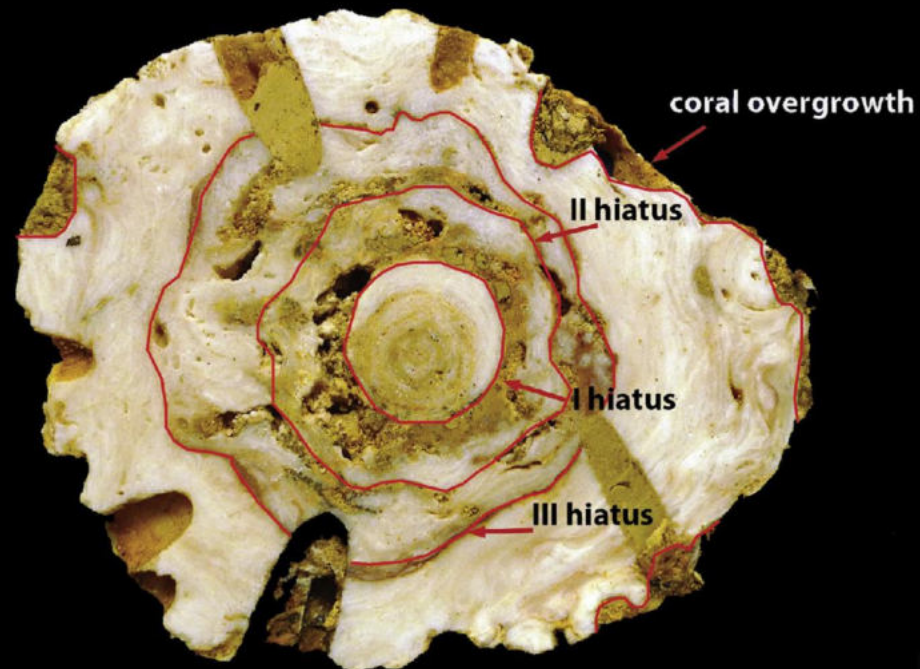
a-e. Reconstructions on the basis of model⁶ driven by benthic oxygen isotope ($\delta^{18}\text{O}$) stack (**a**), benthic $\delta^{18}\text{O}$ data and ice-volume model⁶ (**b**), seawater $\delta^{18}\text{O}$ calculated from paired Mg/Ca and $\delta^{18}\text{O}$ measurements on planktonic foraminifera² (**c**), reconstruction of Red Sea seawater $\delta^{18}\text{O}$ (ref. 3) (**d**) and open-system U-Th ages of corals⁷ (**e**). Shading indicates periods of time for which sea level rises above -20 m (dashed lines), the approximate depth of Argentarola Cave speleothems, Note differences in the number of MIS 7 highstands predicted to exceed -20 m and the difference in elevation predicted for MIS 7.3 in particular.

recovered from Argentarola Cave, Italy, that preserve alternating layers of spelean calcite that grew when the caves were above sea level and biogenic calcite secreted by serpulid worms that colonized the speleothems during seawater submergence⁴. One of the advantages of the speleothem archive relative to corals is that dense spelean calcite is less susceptible to alteration, which allows for reconstructions farther back in time. Furthermore, by dating the timing of speleothem growth, the entire duration of time for a given sea-level highstand can be bracketed. In contrast, it is much more challenging to unambiguously determine the entire duration



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la sezione di una stalattite campionata a Custonaci (Tp, Sicilia). Lo speleotema è ricoperto da coralli e contiene 3 latus, corrispondenti a 4 trasgressioni marine nella grotta che ora si trova a 97 metri di quota. La datazione dei coralli è di 1.1 Milioni di anni. Le 4 trasgressioni corrispondono agli stadi isotopici compresi tra 25,41 fino a 1.5 milioni di anni. Si tratta della più antico speleotema





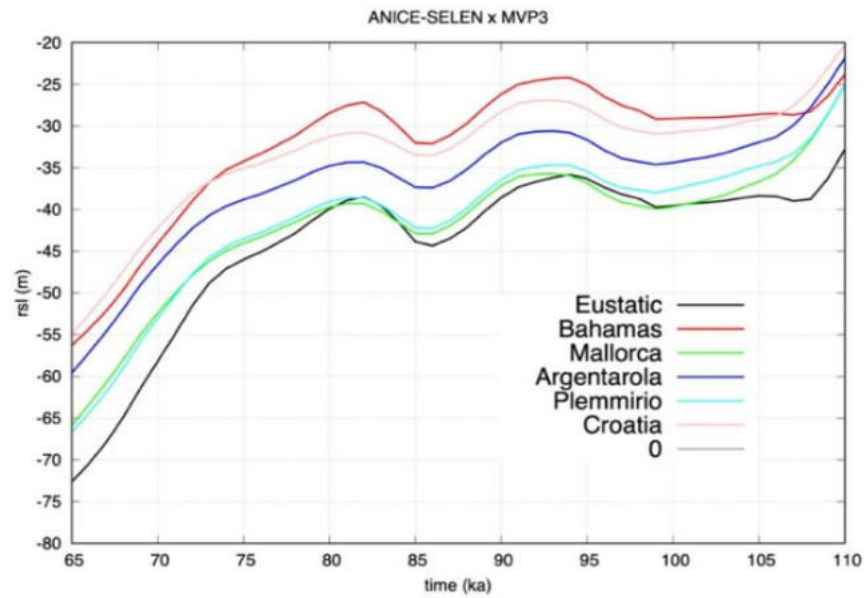


Figure 11. Predicted MIS 5.1–5.3 RSL curves for the Bahamas and at the relevant Mediterranean sites for ANICE-SELEN and the MVP 3 mantle viscosity profile. The black curve shows the eustatic.

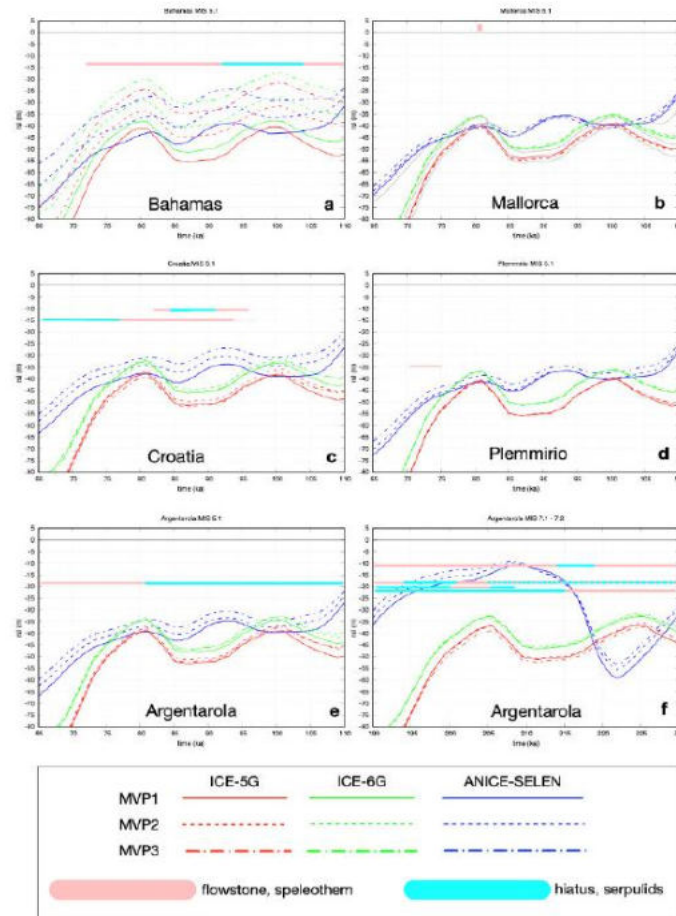


Figure 8. Predicted RSL curves for ICE-5G (red curves), ICE-6G (green curves), and ANICE-SELEN (blue curves) ice sheet models, in combination with mantle viscosity profiles (MVP) 1–3 (solid, dashed, and dotted, respectively) at each site and with respect to the measured elevations. (a) predicted RSL curves at MIS 5.1 in the Bahamas and with respect to the elevation of the flowstone. (b) predicted RSL curves at MIS 5.1 in Mallorca. (c) predicted RSL curves at MIS 5.1 in Croatia and elevation of the speleothems. (d) predicted RSL curves at MIS 5.1 at Plemmirio and elevation of the speleothem. (e) predicted RSL curves at MIS 5.1 at Argentarola and elevation of the speleothems. (f) predicted RSL curves at MIS 7.1–7.2 at Argentarola and elevation of the speleothems.

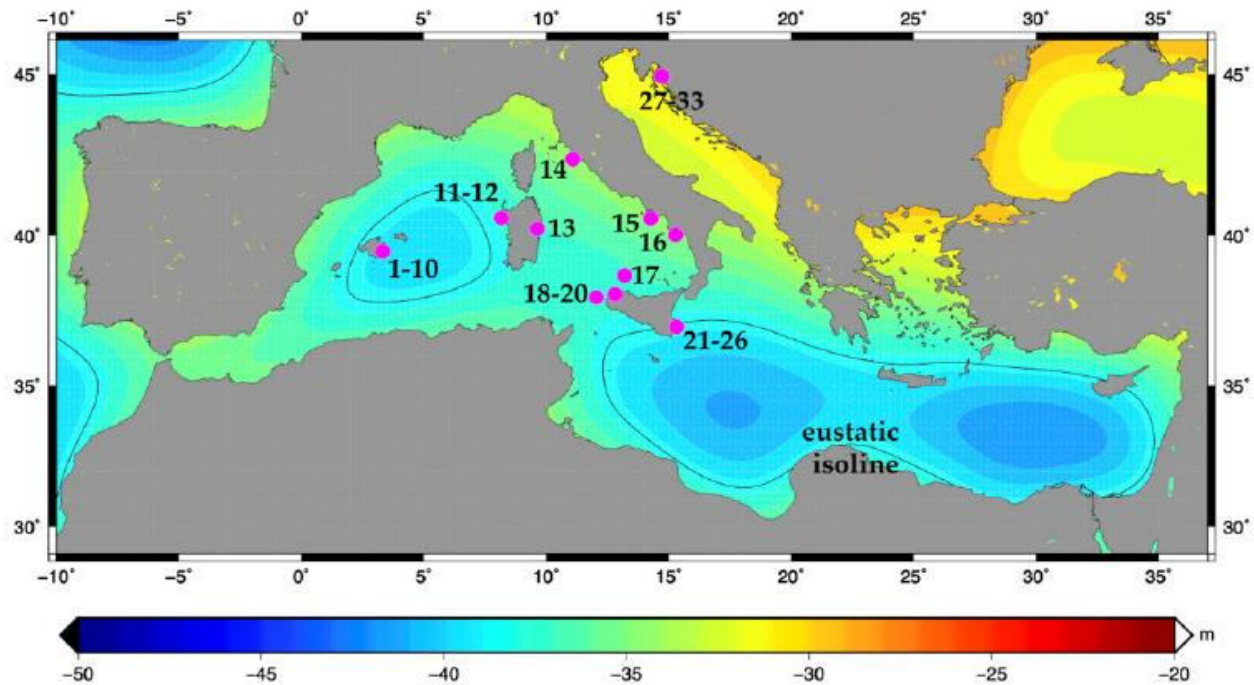


Figure 12. Predicted RSL elevation in the Mediterranean Sea during MIS 5.1 according to ANICE-SELEN and the MVP 3 mantle viscosity profile. The pink dots indicate the location of the sites. The black isoline corresponds to the eustatic value (-38.5 m).

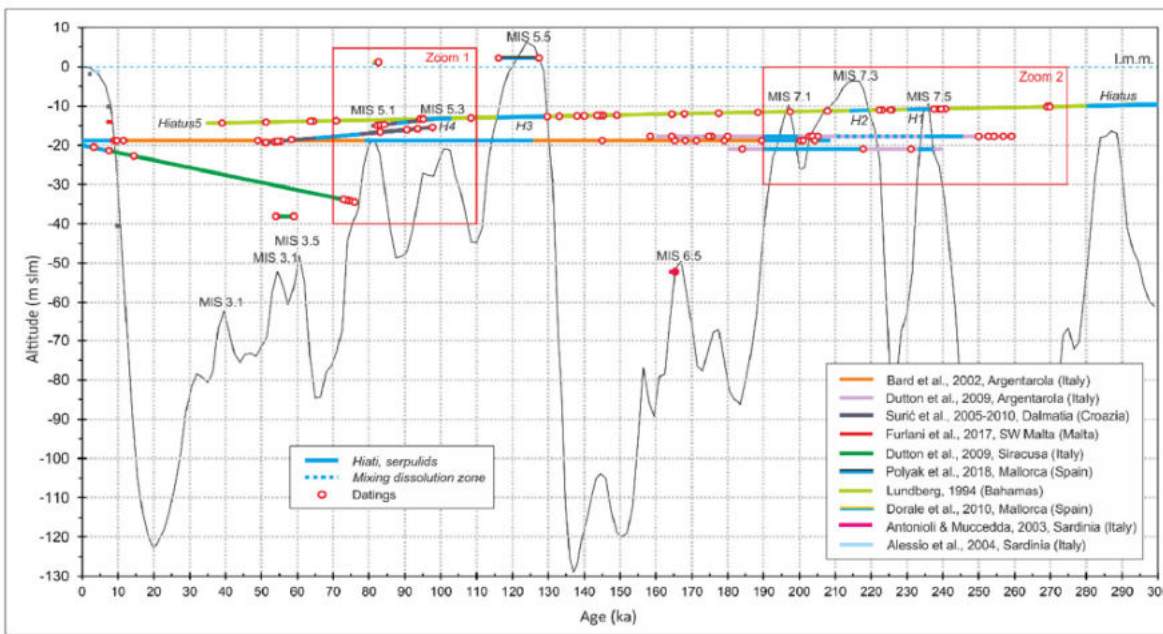


Figure 13. Elevations and radiometric ages (¹⁴C and U-Th) of the Mediterranean speleothems and marine overgrowths discussed in the present study and a comparison with the DWBAH flowstone Table A1. Argentarola (Italy), -18.5 m [2] and -18 and -21.7 m [22]; Plemmirio (Italy), -23 m [64]; Grotta di Nettuno (Italy), - 3 m [81]; U vode Pit (Krk Island, Croatia), Stalagmite K- 4 (-14.5 m) and K-18 (-18.8 m) [3]; POS from Mallorca (Spain), + 1.5 m [4,5]; Malta [65]; DWBAH Flowstone (Bahamas), -15 m [11]. Black line: global sea level curve reconstructed by [88].

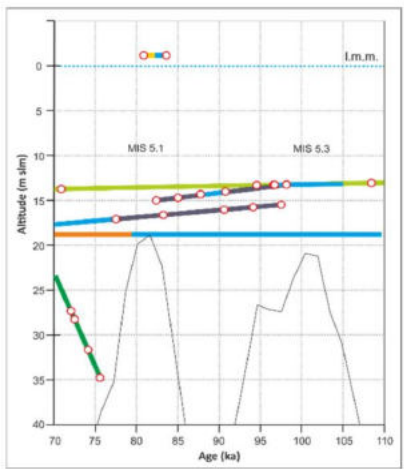


Figure 14. Zoom 1 of Figure 13. Light green line: DWBAH flowstone [11]; orange line: Stalagmite ASI (-18.5) from Argentarola cave, -18.5 m [2]; black line: Speleothems K14 and K18; yellow and blue line: Mallorca POS; green line: Plemmirio [64]; fine black line: global sea level curve reconstructed by [88].

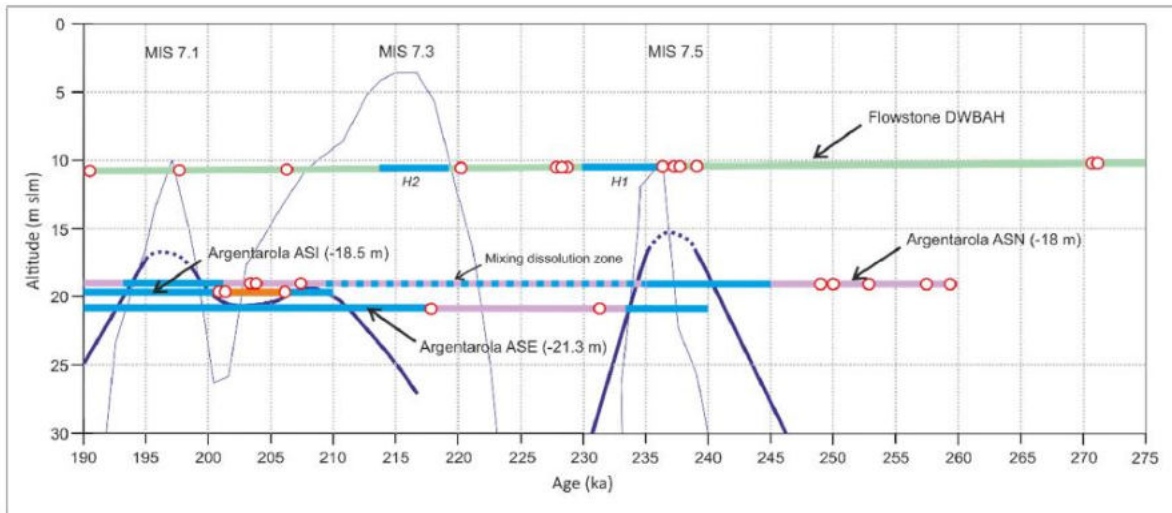


Figure 15. Zoom 2 of Figure 15. Green line: DWBAH flowstone [11]; orange line: Stalagmite ASN (-18 m) from Argentarola cave [22], and stalagmite ASI (-18.5) from Argentarola cave [2]; black line: global sea level curve reconstructed by [88]; dark blue line: sea level drawn with observed data.

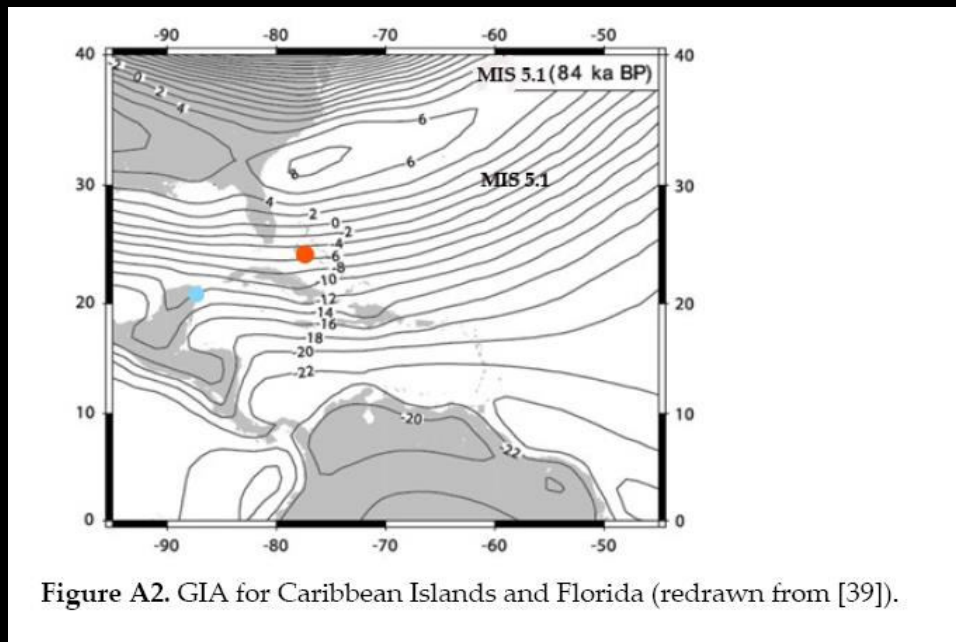


Figure A2. GIA for Caribbean Islands and Florida (redrawn from [39]).

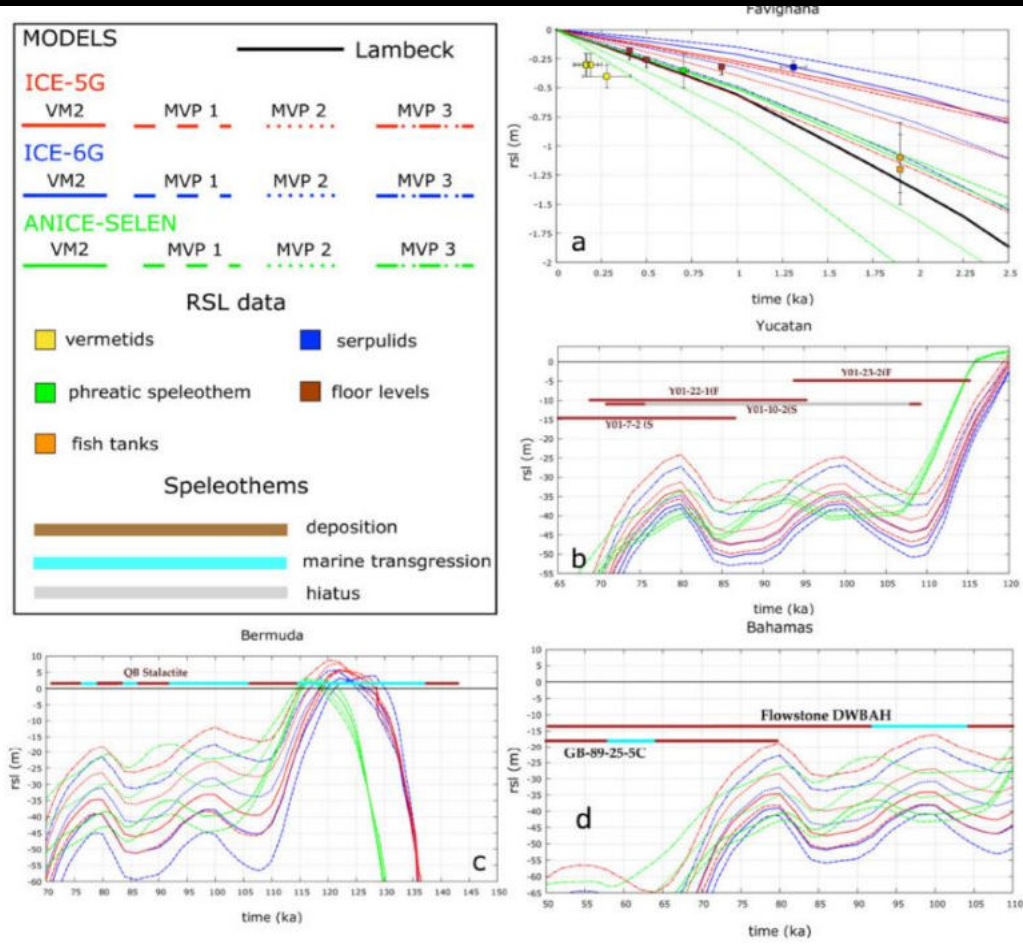


Figure 10. Predicted RSL curves for the ICE-5G (red curves), ICE-6G (green curves), and ANICE-SELEN (blue curves) ice-sheet models in combination with MVP 1–3 mantle viscosity profiles (solid, dashed and dotted lines, respectively) at each site and with respect to the measured elevations. (a) Favignana. (b) Yucatán. (c) Bermuda. (d) The Bahamas.

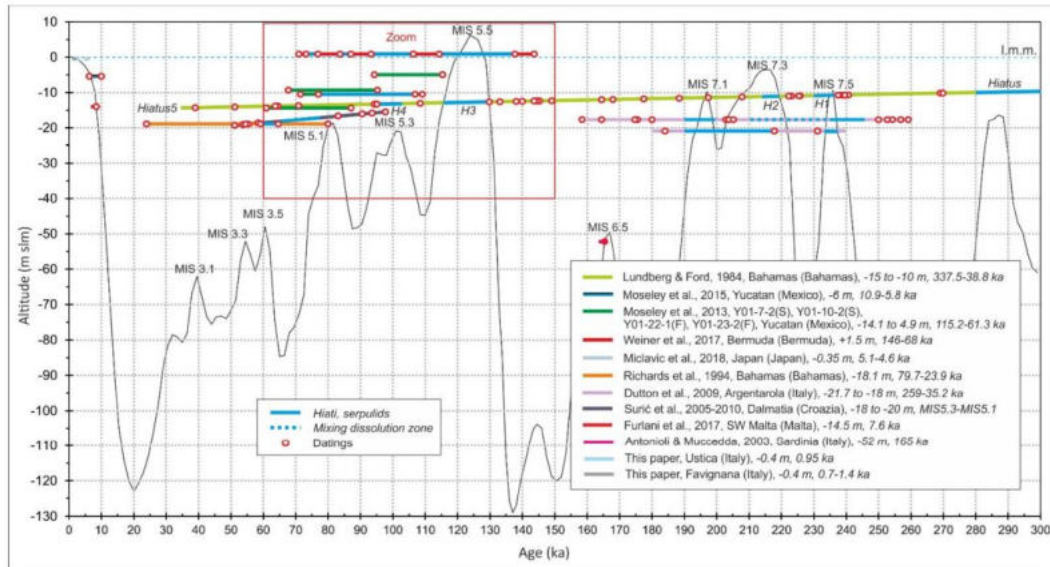


Figure 13. Elevations and radiometric ages (14C and U/Th) of the Caribbean, Japan, and some Mediterranean speleothems discussed in the present study and comparison with the: global sea level curve reconstructed by [79] Black line.

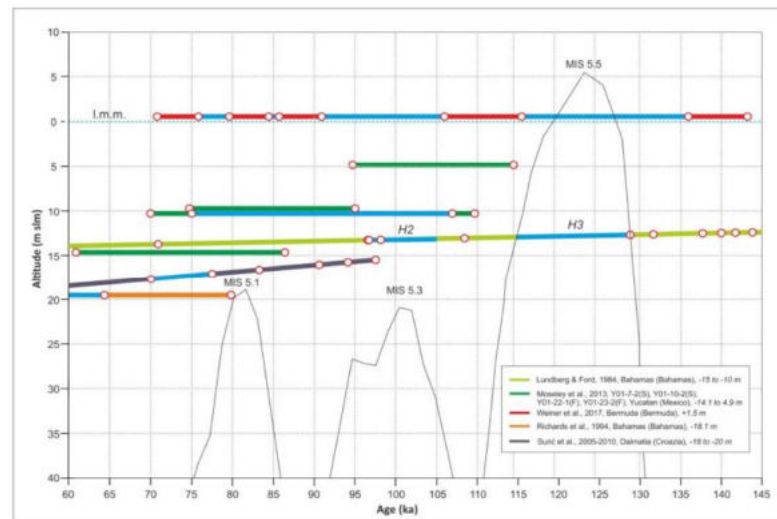
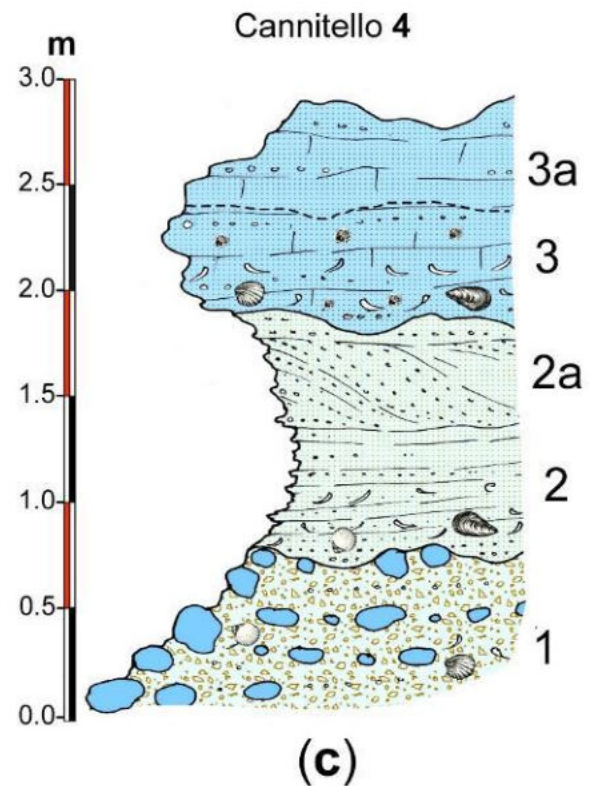
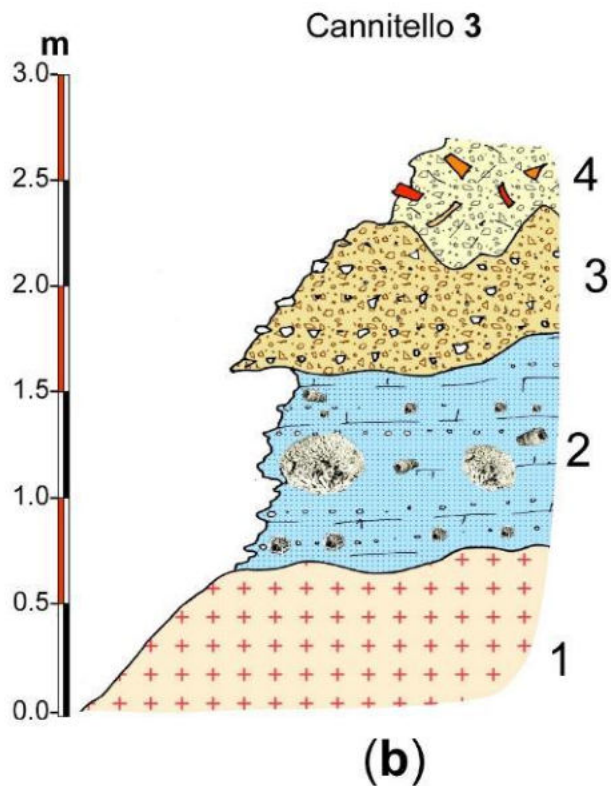
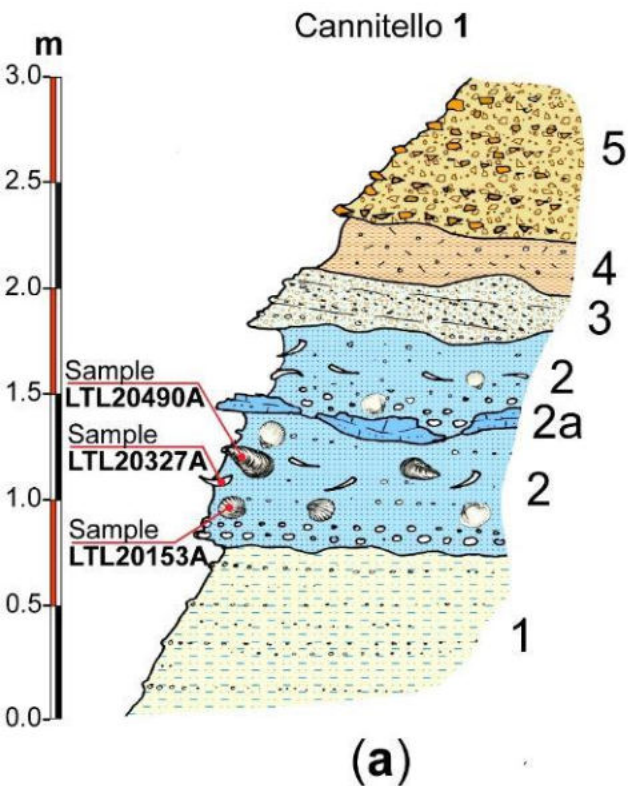


Figure 14. Zoomed in view of part of Figure 13. Fine black line is the global sea-level curve reconstructed by [79].




Article

New Evidence of MIS 3 Relative Sea Level Changes from the Messina Strait, Calabria (Italy)

Fabrizio
Paolo O
Paolo St

Given the overall scarceness of MIS 3 marine outcrops that are explorable in coastal areas subject to important uplift, we consider the southern Calabrian site relevant for the assessment of past sea levels during this still poorly known interstadial.

11000 5,6,7 

The GIA results suggest that the $\delta^{18}\text{O}$ -based ice sheet models appear to significantly overestimate the ice sheet volumes during the MIS 3.1 and 3.3. Our data are in agreement with Gowan et al. [101], raising, by 40 metres, the eustatic contribution to sea level during interstadials MIS 3.1, 5.1, and 5.3 with respect to the current global sea level curve scenarios. Further, our reconstruction agrees well with the records proposing MIS 3 sea levels at depths between -18 and -40 m.

Our numerical results and observational data confirmed that the MIS 3 RSL changes at Cannitello are governed by glacio-eustasy, whereas GIA plays a secondary role. While the $\delta^{18}\text{O}$ dependent ice sheet models result in RSL curves that are always significantly lower than the observations, PaleoMIST 1.0 is the only model capable of returning a MIS 3.1 elevation that is in agreement with the observations. Indeed, we observe that there is a discrepancy (of at least 30–40 m) between the eustatic altitude of the MIS 3 of all global curves and those suggested by observations. Therefore, our results confirm previous evaluations by Pico et al. [63] and Gowan et al. [101] and support the contention that a reduction of global ice sheet volumes across the MIS 3, and specifically at the MIS 3.1 and 3.3, is needed.

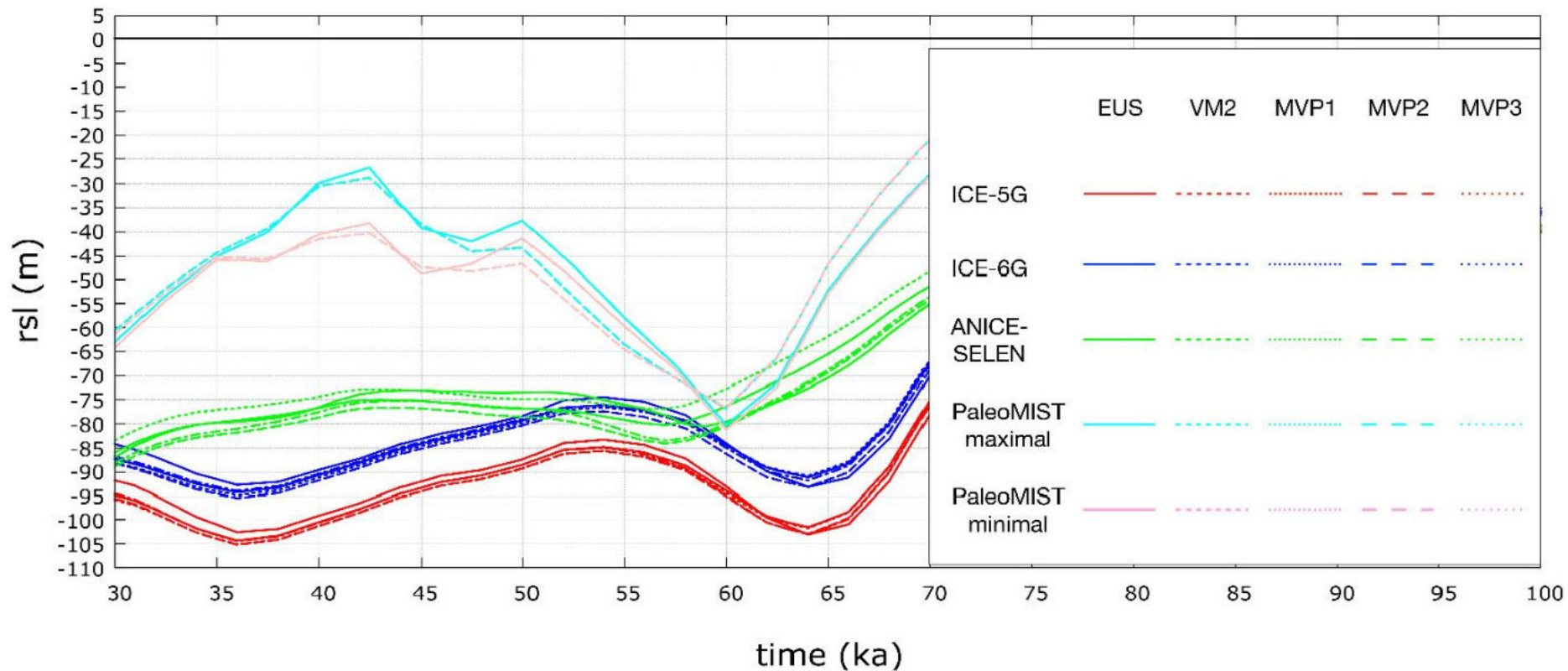


Figure 7. Predicted RSL curves for four ice sheet models: ICE-5G (red), ICE-6G (blue), ANICE-SELEN (cyan), and PaleoMIST 1.0 maximal and minimal (cyan and pink, respectively). The solid curves represent the eustatic. The dashed and dashed-dotted curves represent the GIA-modulated RSL predictions.

Time (ka)	ICE-5G (m)	ICE-6G (m)	ANICE-SELEN (m)	PaleoMIST maximal (m)	PaleoMIST minimal (m)
30	-95	-88	-65	-85	-88
35	-102	-92	-45	-80	-92
40	-98	-88	-30	-75	-88
45	-92	-82	-40	-75	-82
50	-88	-78	-45	-75	-78
55	-85	-75	-65	-75	-75
60	-95	-85	-80	-80	-85
65	-102	-92	-55	-85	-92
70	-95	-85	-30	-80	-85



Sicily, St Vito Lo Capo -56 m. Is this an archaeological marker for sea level change?

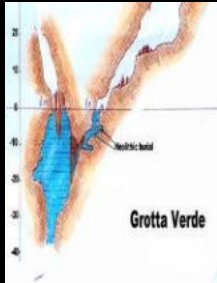
Sequenza temporale delle evidenze archeologiche di variazioni del livello del mare negli ultimi 22 mila anni

-11 m

-2.5 m

-0.3 m

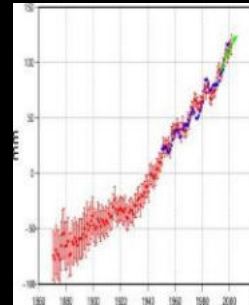
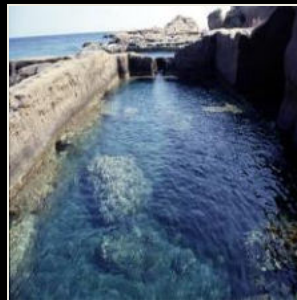
-130 m



-13 m

-1.35 m

3 mm/yr



~22 mila anni
Grotta Cosquer
(Francia) graffiti

~7 mila
Pozzi (Israele)
Grotta Verde
(Sardegna)

~2.4 mila
Capo Malfatana
Molo frangiflutti
Epoca Punica

~2 mila
Piscine
(Lutiane)

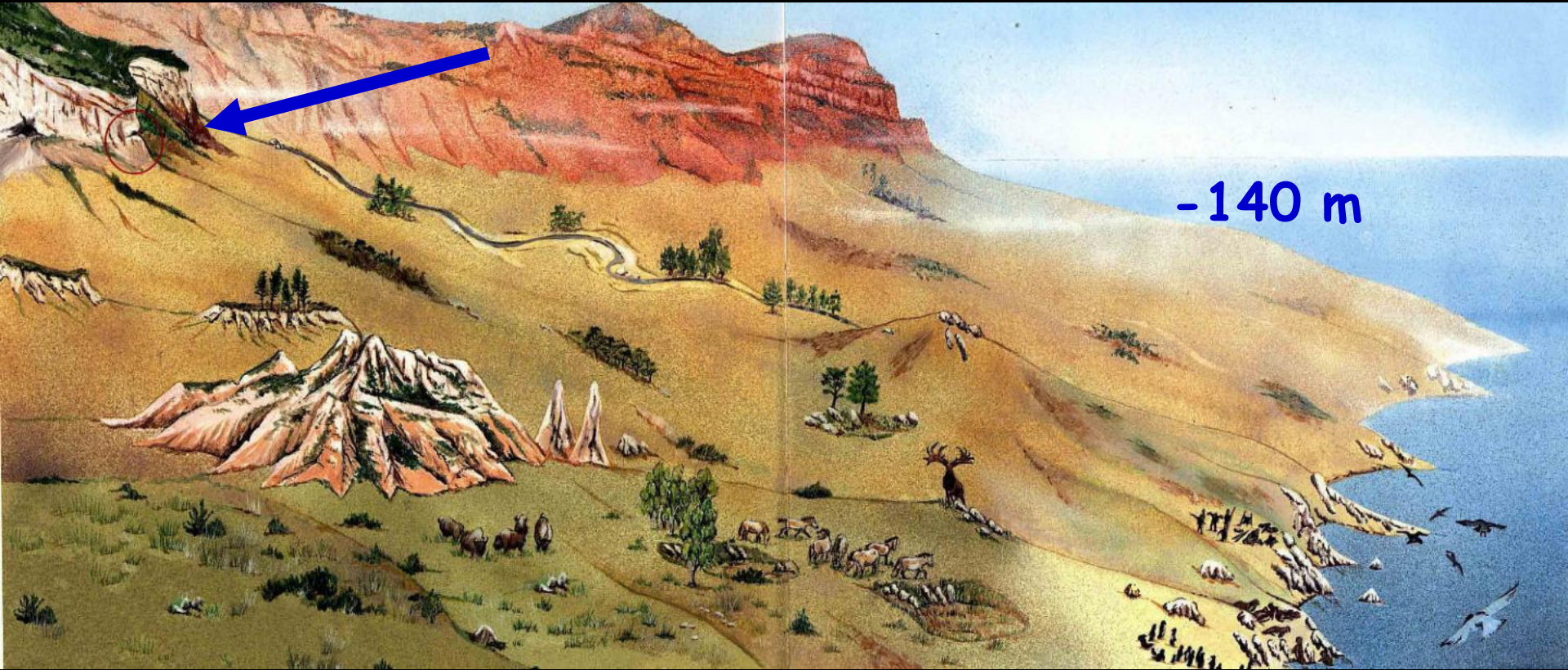
Mille
Millstones

Oggi
Mareografi



Cosquer cave - France

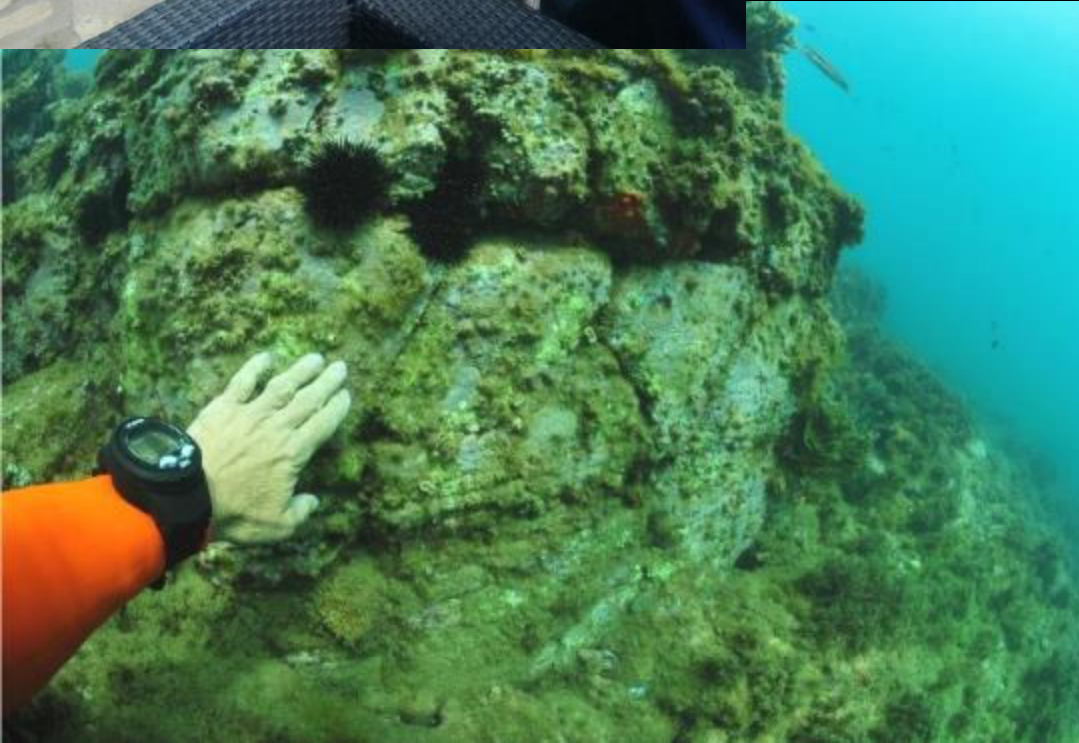
Today at -37 m

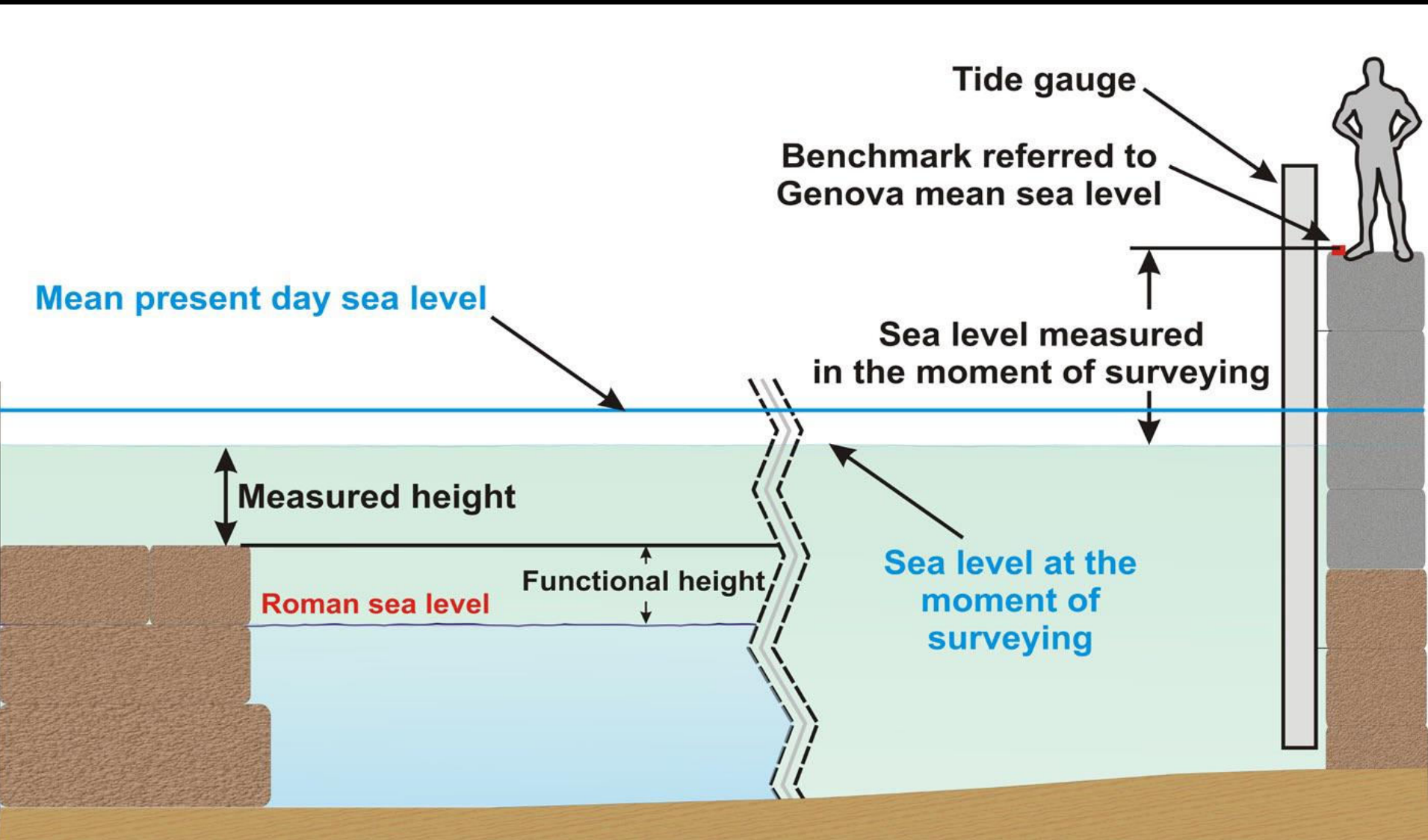


-140 m

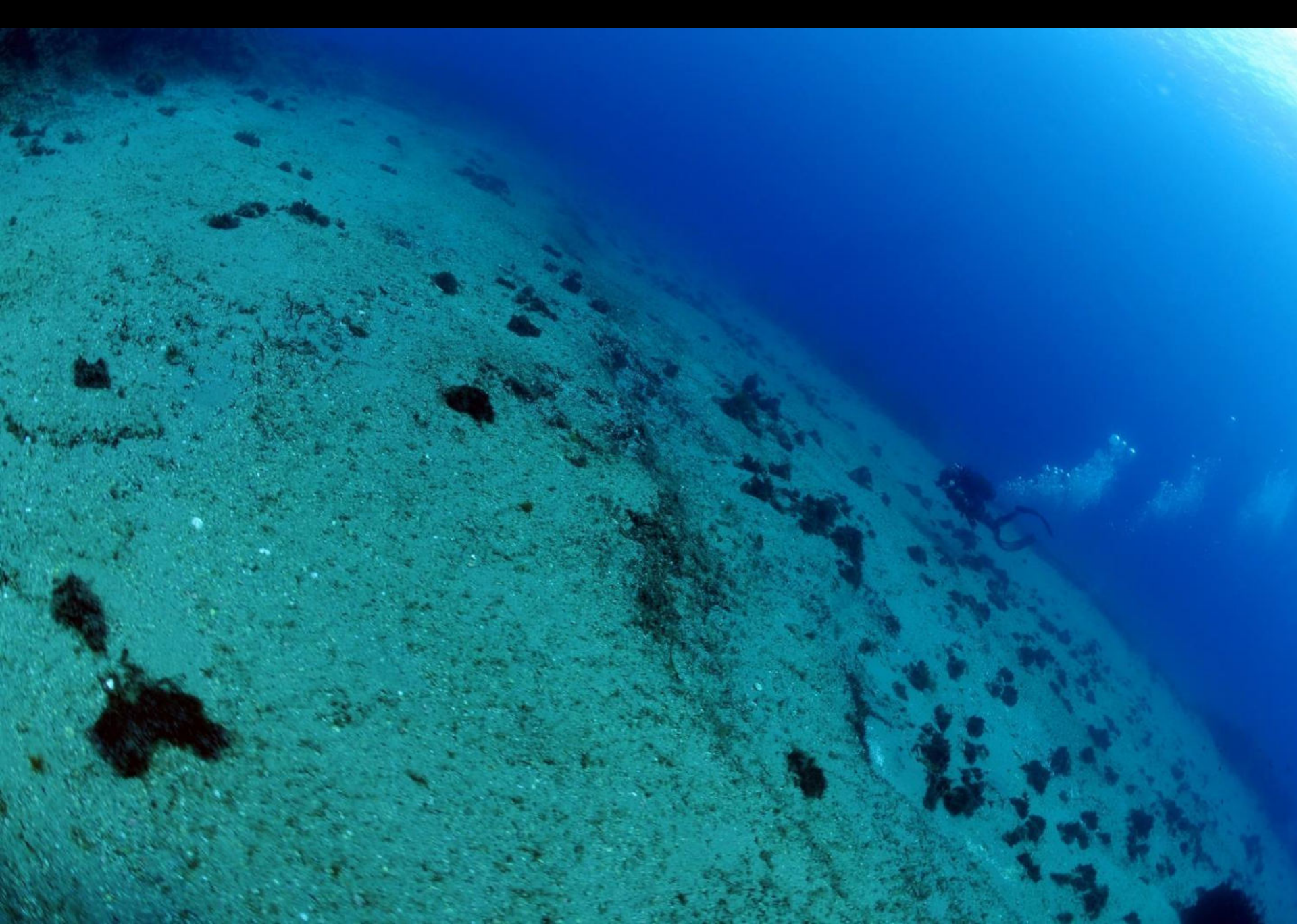
22 ka cal BP (dated from wall paintings, Clottes et al., 1992)

Sott'acqua puo essere facile prendere abbagli, quella che sembra una strada o un muro o un prosciutto, sono invece formazioni geologiche naturali.

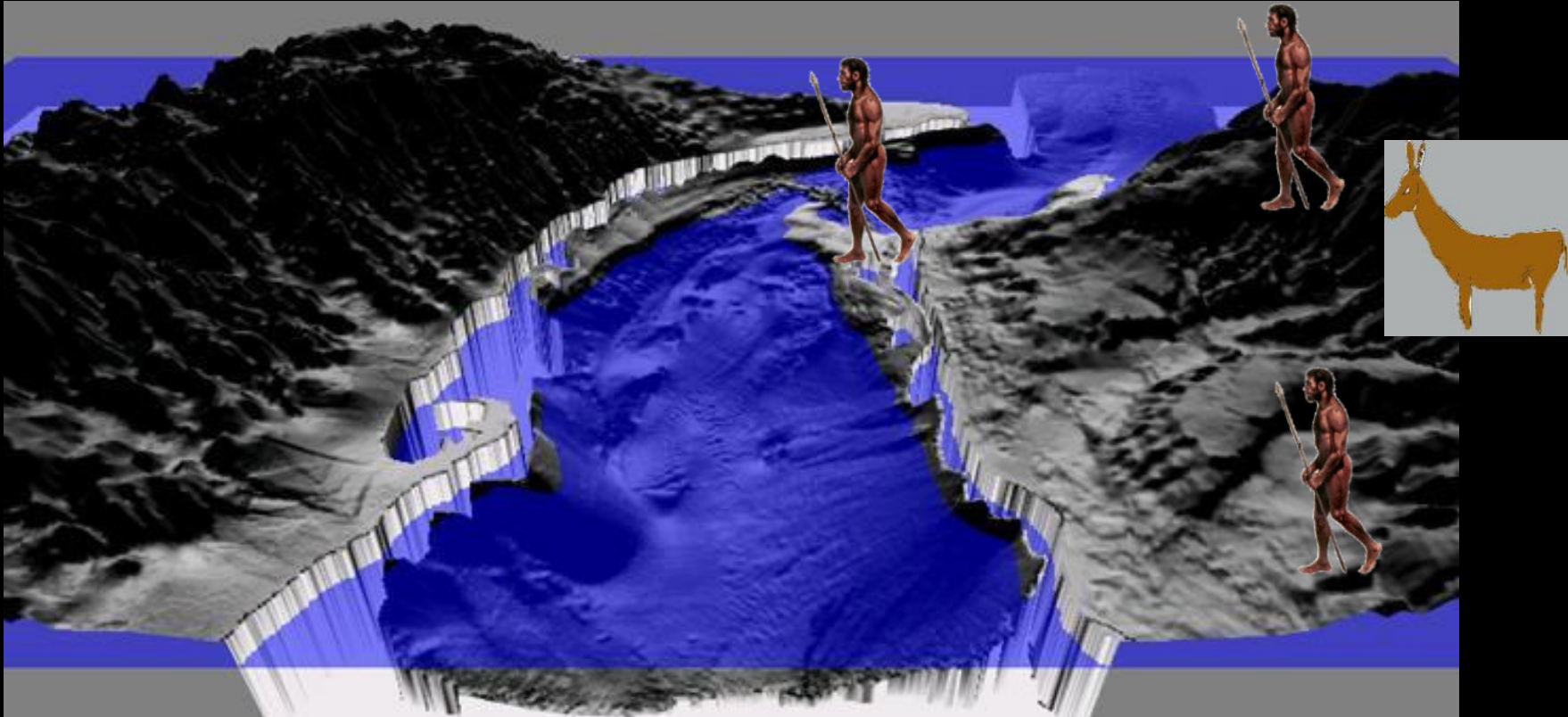






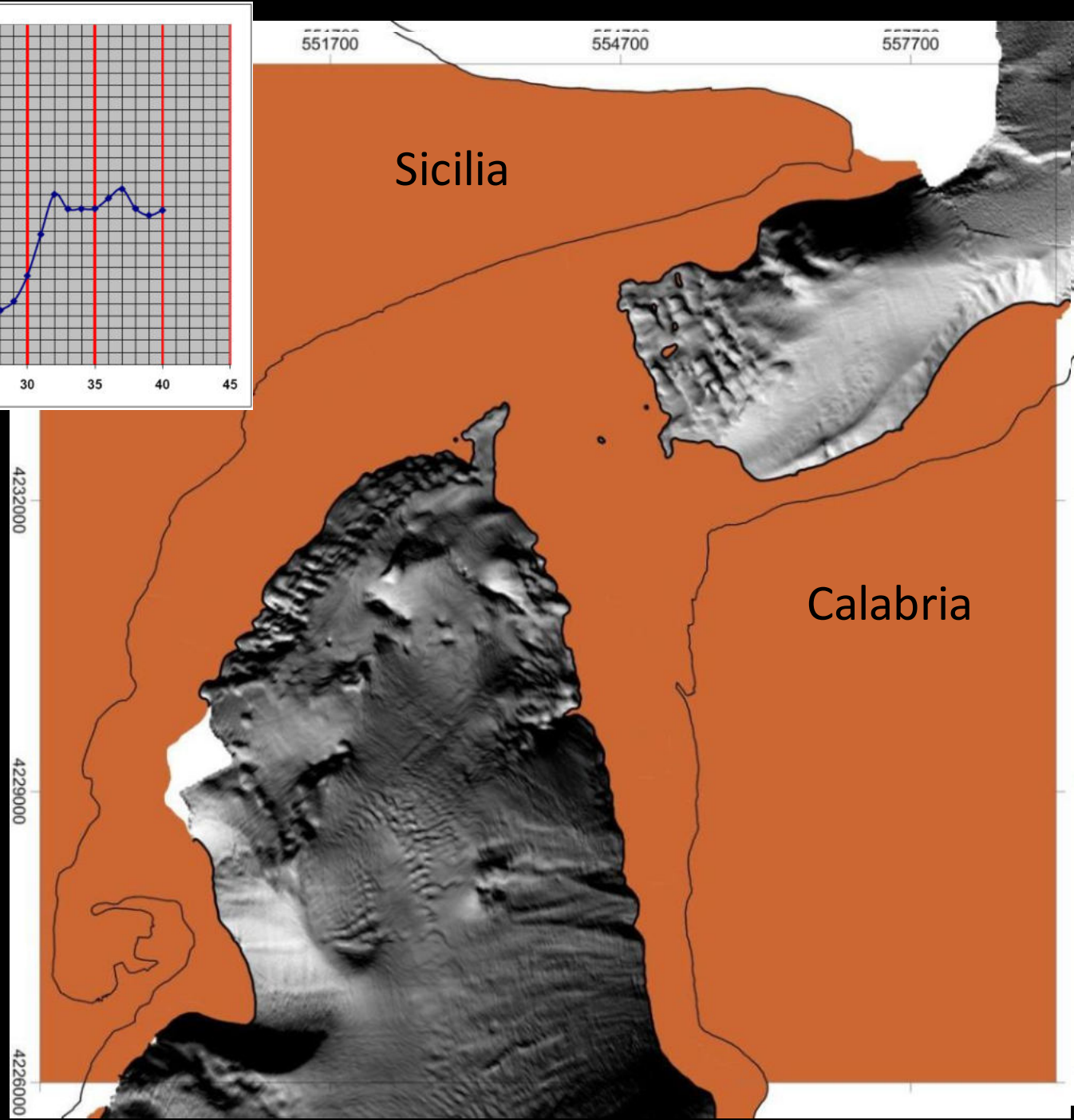
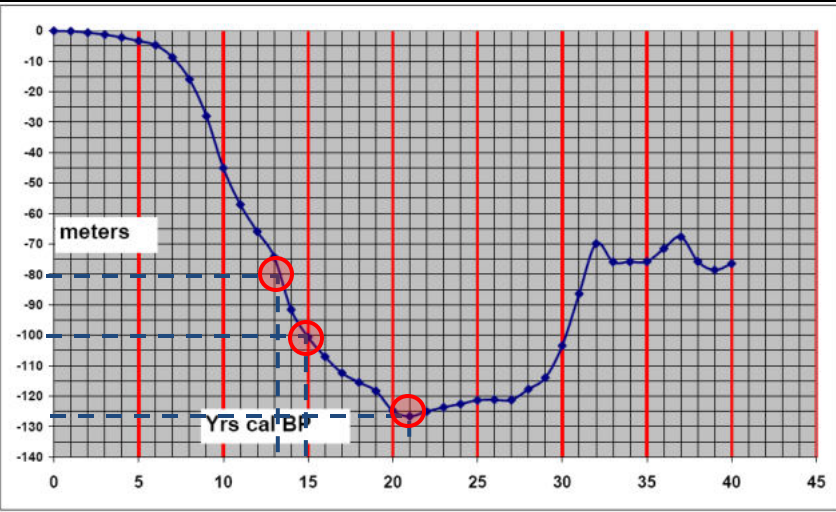


Conclusioni



La litologia dei “pilastri” sommersi che compongono parte della Sella, corrisponde alla Formazione delle Ghiaie di Messina.

Il nostro schema di emersione del Ponte Continentale viene validato dagli studi antropologici per le presenze dell’*Equus Hydruntinus* e dell’ *Homo sapiens* in Sicilia, rispettivamente datate a 18.8 e 17.5 ka cal BP.

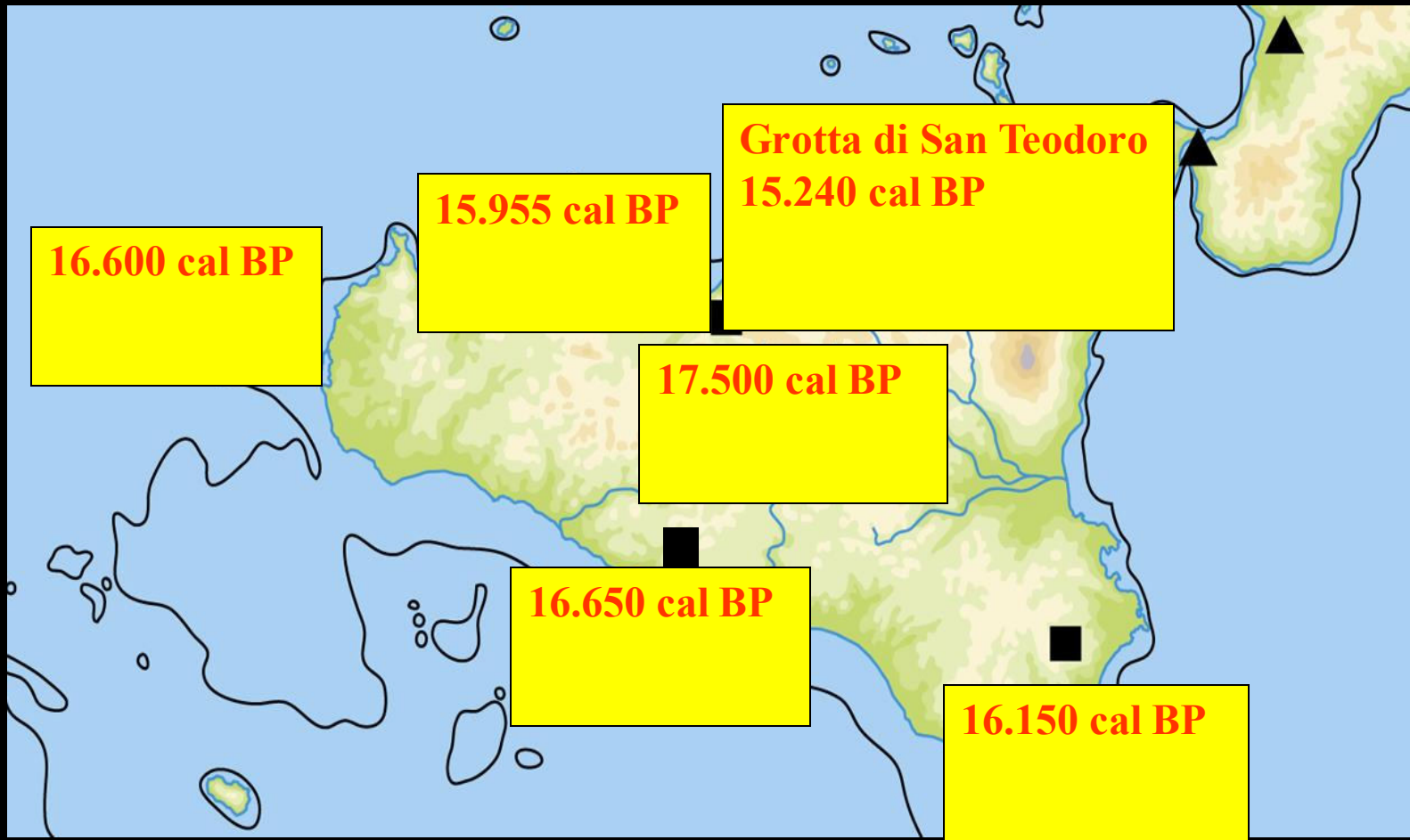


Esistono conferme a questo schema di evoluzione?

Homo sapiens in Italy about 30/43 ka

Homo sapiens in Sicily, about 17 ka





Intervallo di tempo minimo: 21.5-20 ka;
Intervallo di tempo massimo: 18-29 ka;



Mondeval de Sora, VF1, Sauveterriano
Industria litica
1 nuclei e lamelle (Scaglia Variiegata)
2 nuclei e lamelle (Scaglia Rossa)
3 nuclei e lamelle (Biancone)



Mondeval de Sora, VF1, Sauveterriano
Noduli di ocre (ossidi di ferro)



Mondeval de Sora, VF1, Sauveterriano
Industria litica
1 nuclei
2 bulini
3 grattatoi



Mondeval de Sora, VF1, Sauveterriano
Conchiglia con foro per sospensione
(Columbella rustica)



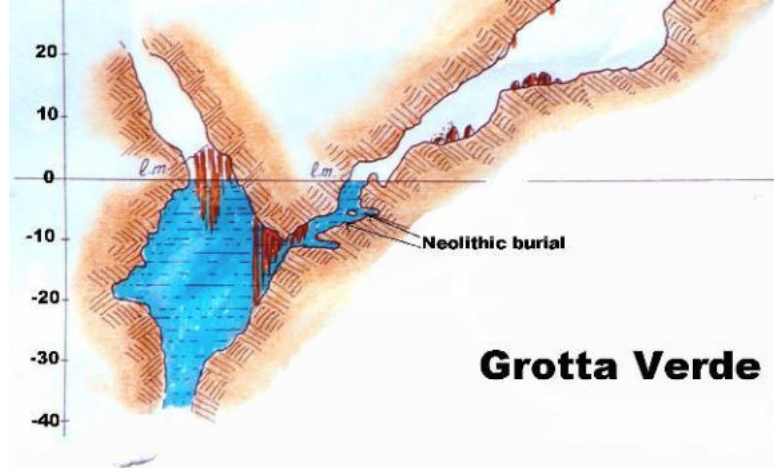
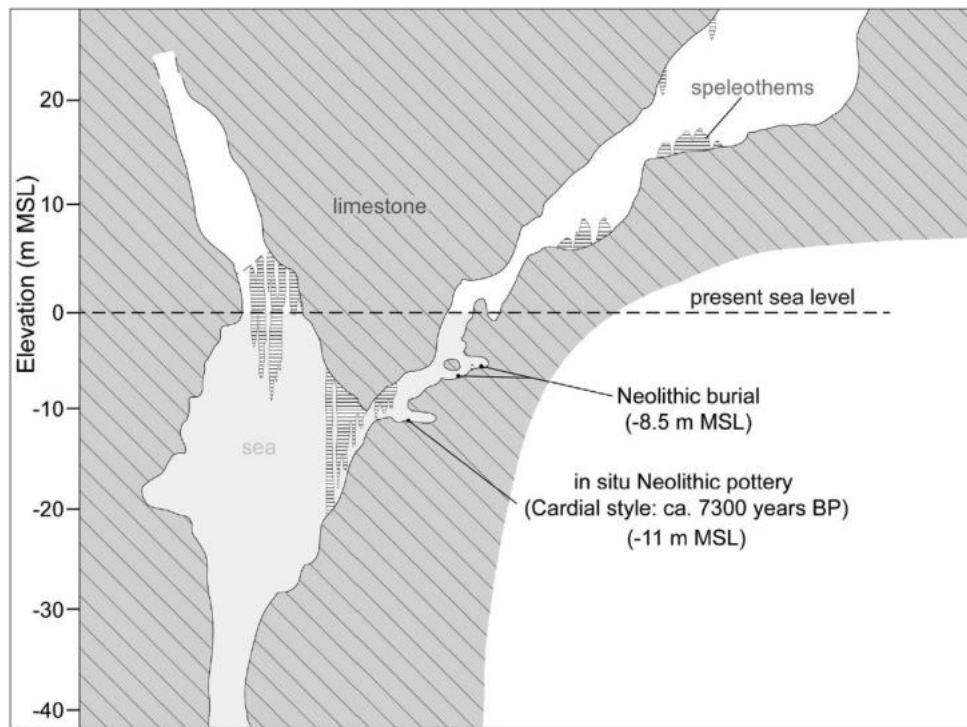
Mondeval de Sora, VF1, Sauveterriano
Industria litica
troncature
schegge a ritocco erto
lamelle a dorso



Mondeval de Sora, VF1, Sauveterriano
Canino atrofico di cervo con foro
per sospensione



Mondeval de Sora, VF1, Sauveterriano
Industria litica



Grotta Verde

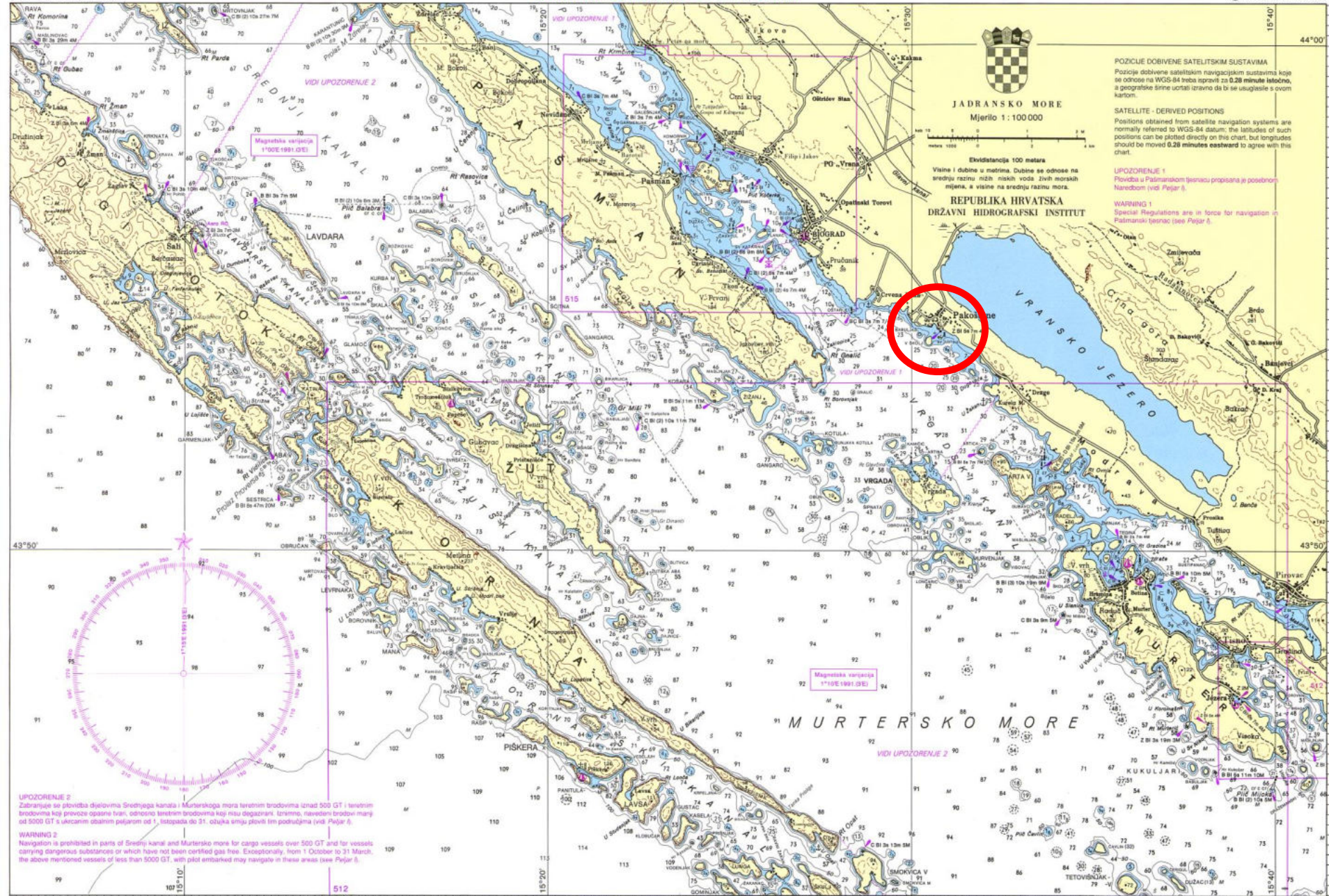
Capo Caccia, Sassari, Sardegna, si tratta dell'affioramento più antico e più interessante mai studiato in Italia, scoperto negli anni '70, fu studiato dalla locale Soprintendenza con notevoli difficoltà tecniche, buio, sospensioni fangose, ecc. Ma vennero ritrovate inumazioni in situ, nelle tombe scavate nella roccia a circa 8 metri di profondità, le ceramiche cardiali tipiche del Neolitico Antico Sardo dettero la possibilità di una datazione precisa: 7300 anni fa (Antonioli et al., 1996, Benjamin et al 2017). Alcuni vasi trovati in situ, sull'orlo di una profonda cavità sulla quale, ai tempi delle inumazioni l'acqua dolce galleggiava su quella salata, dettero la possibilità di attribuire al livello del mare di 7.3 mila anni fa una quota inferiore a -11 metri. Tutto ciò in perfetto accordo con quanto predetto dal Modello di Lambeck et al., 2011.



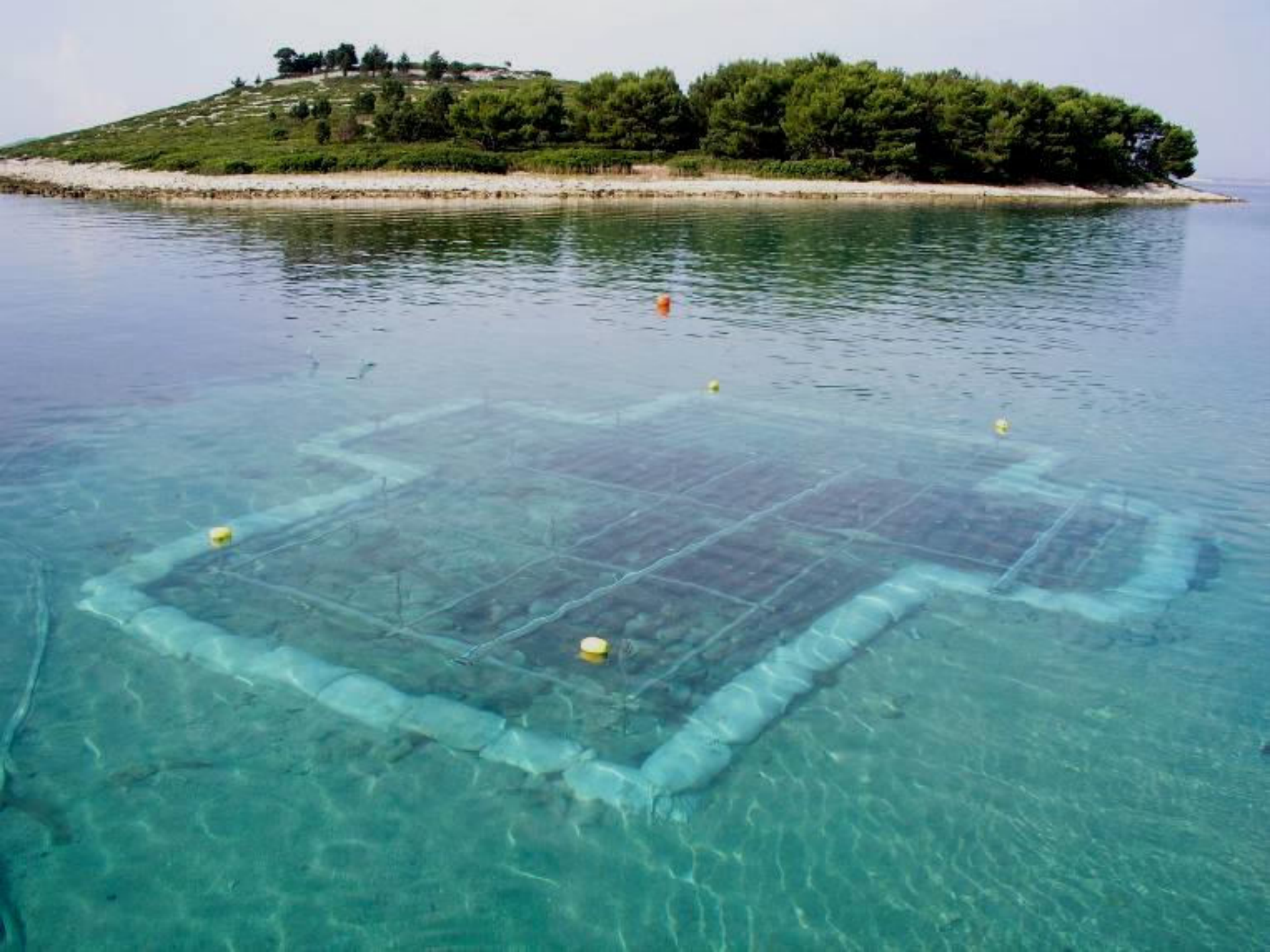
Data SIO, NOAA, U.S. Navy, NGA, GEBCO
Image © 2009 TerraMetrics
Image © 2009 DigitalGlobe
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1125 mi

1°38'57.05" E











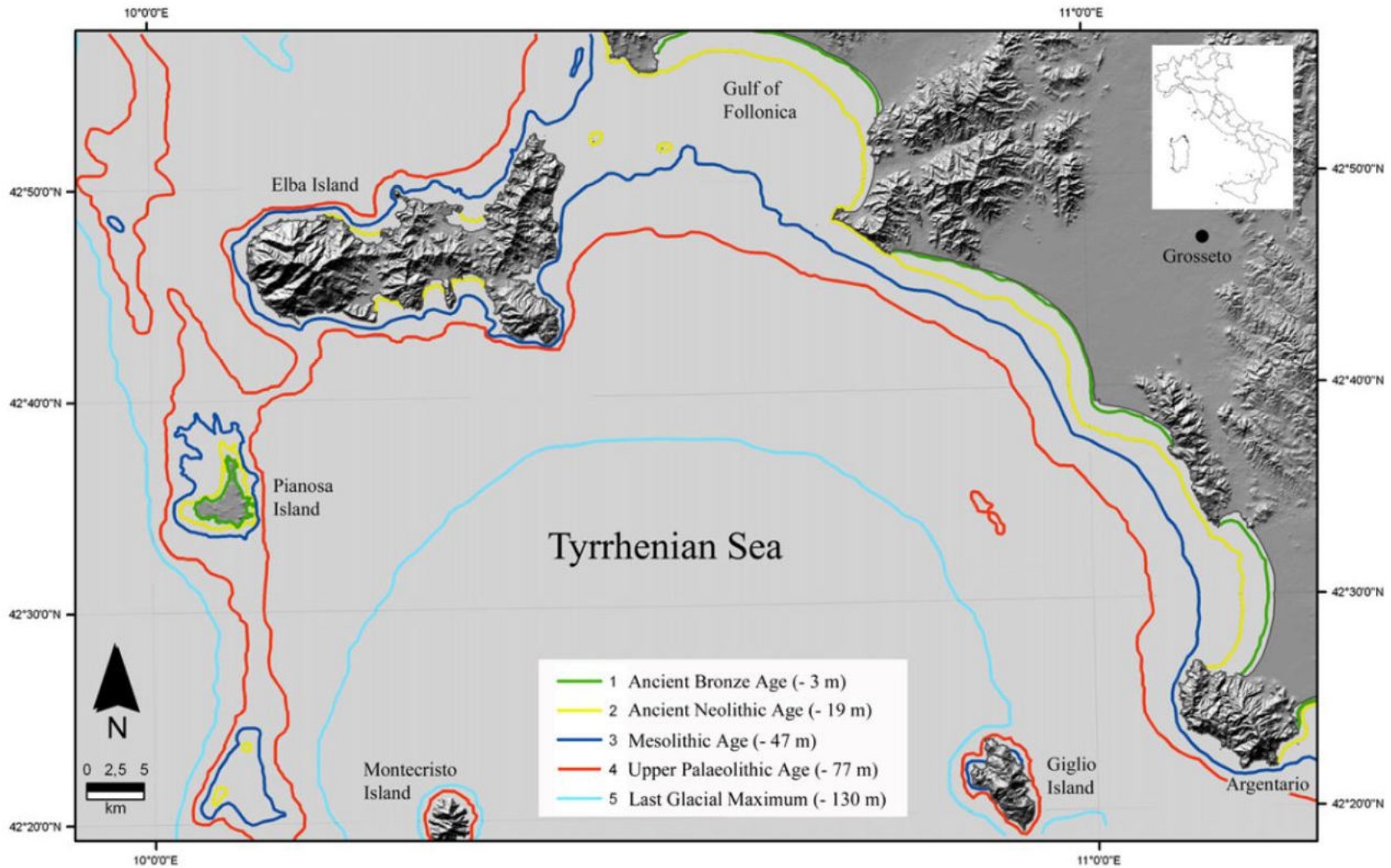


Fig. 11. Palaeocoastline variations since 20 ka cal BP on central Italy coast. See Fig. 10 for timing and sea level change curves.

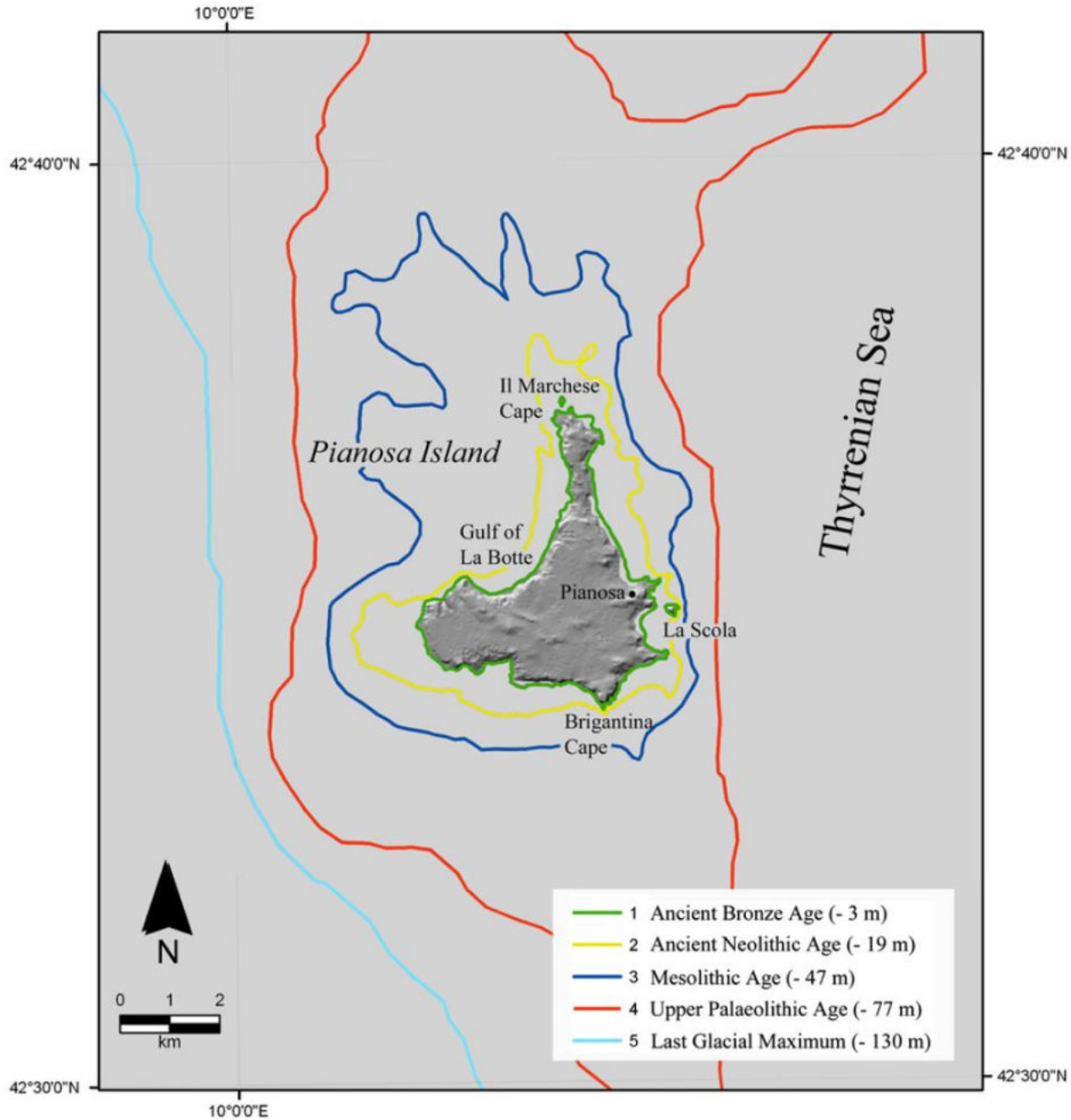
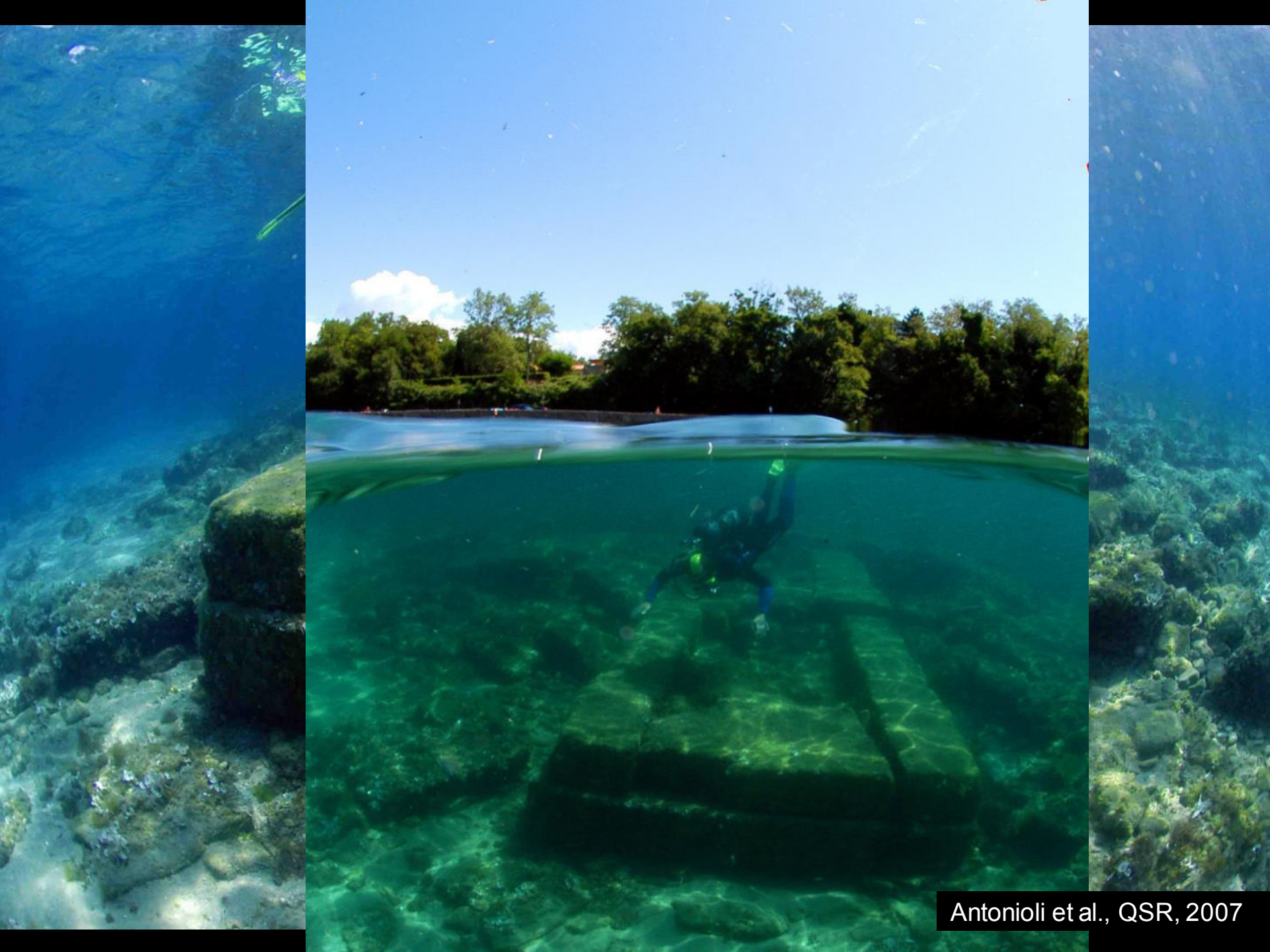


fig. 12. Palaeocoastline variations since 20 ka calBP at Pianosa. See Figs. 10 and 11 for timing and sea level change curves.



**Prof. Rita
Auriemma molo
romano di Punta
Sottile, Trieste**







Ponza fishtanks

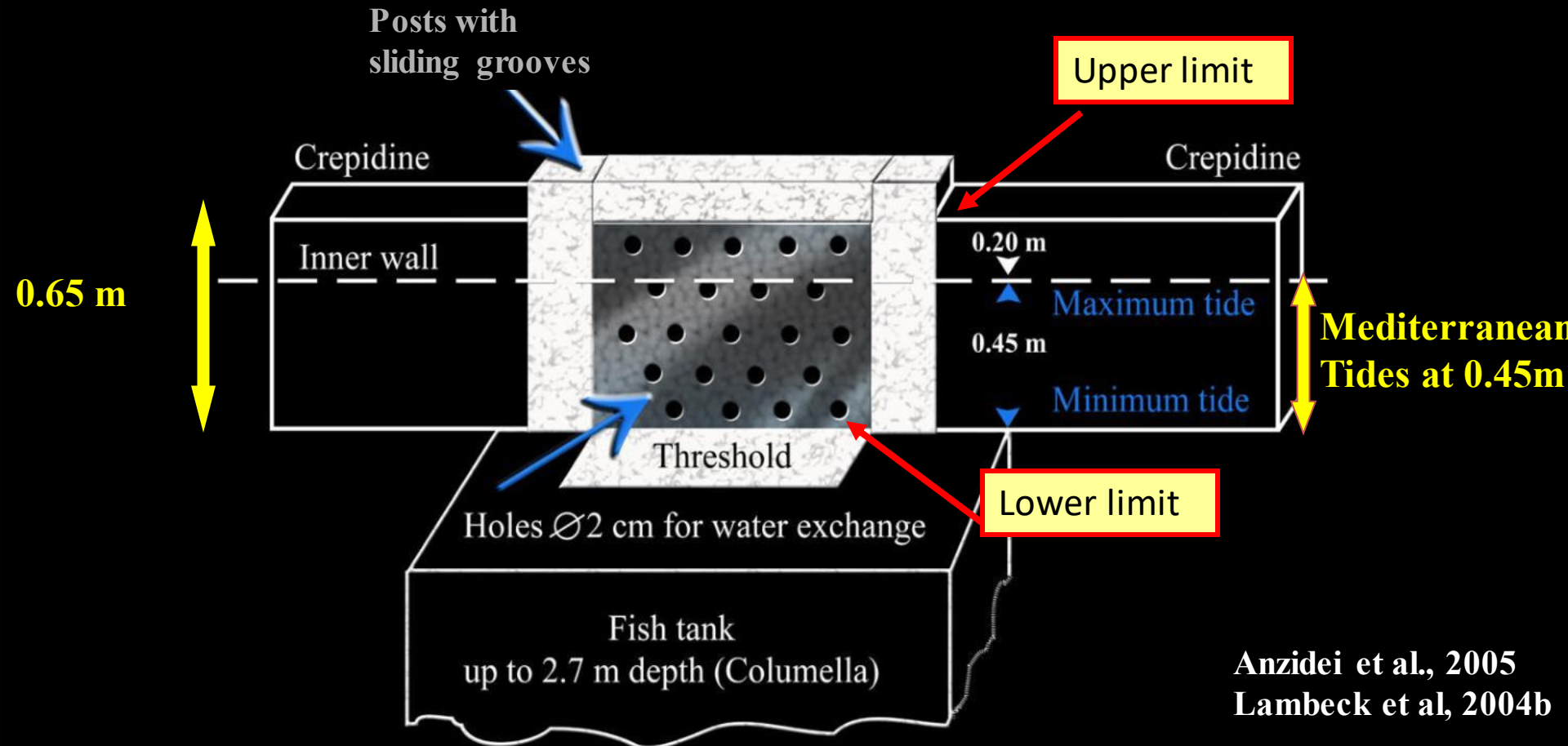


La piscina di allevamento per pesci di Punta della Vipera, Civitavecchia, uno dei migliori esempi di realizzazione ingegneristica di 2000 anni fa, in questo caso i canali di entrata ed uscita dell'acqua si trovano a -1,28 metri. Nel passato, misure errate effettuate sui muretti e non dei canali avevano fornito la quota di -50 cm.



Sluice gates: the precise ~2ka benchmarks

Sketch of a sluice gate for the water exchange in a Roman Fish tank



The top of the sluice gate coincides with the elevation of the lowest level foot-walk (crepidine), to a position above the highest tide level.



Millstone coastal quarries of the Mediterranean: A new class of sea level indicator

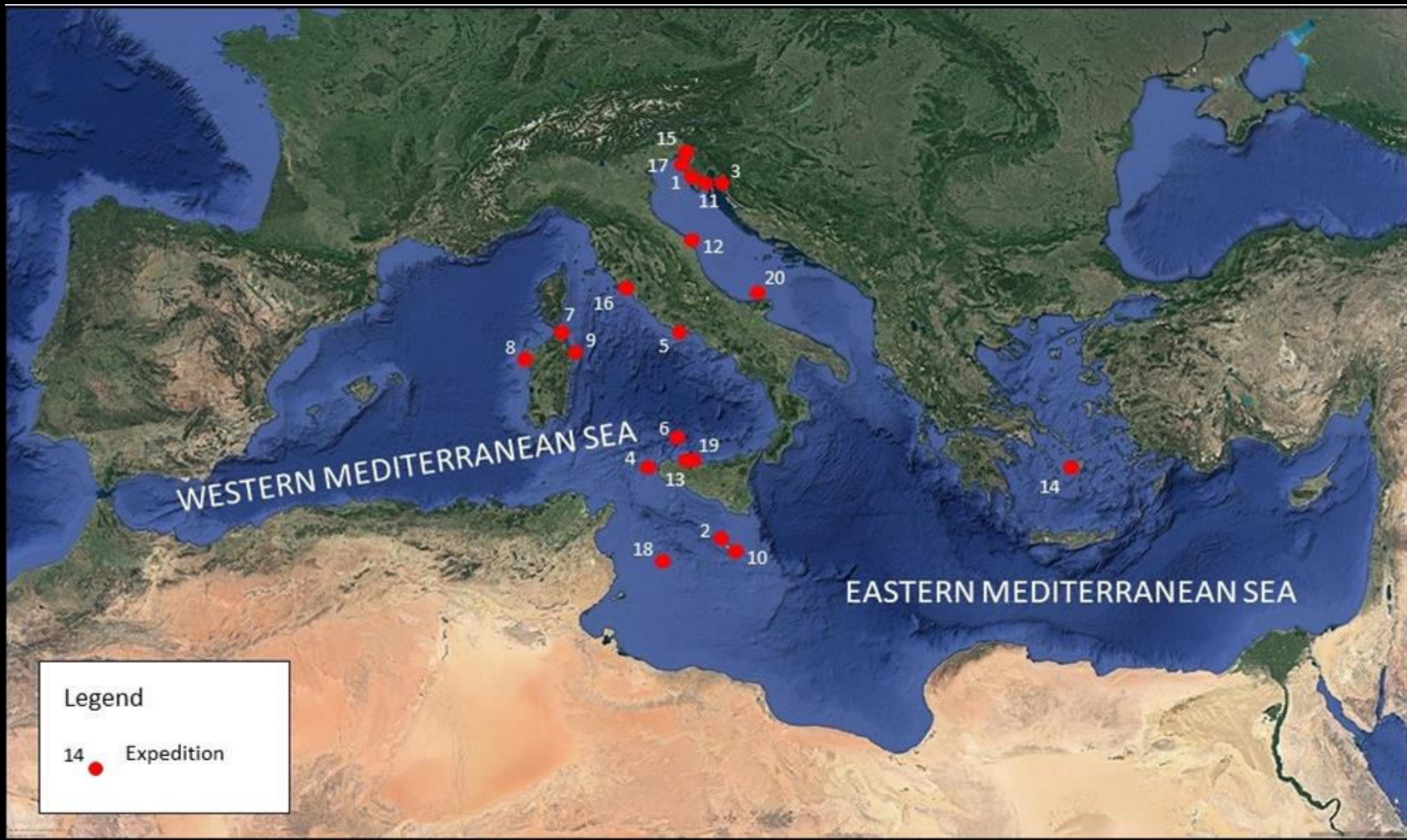
V. Lo Presti ^{a,*}, F. Antonioli ^b, R. Auriemma ^c, A. Ronchitelli ^d, G. Scicchitano ^{e,f},
C.R. Spampinato ^{e,g}, M. Anzidei ^h, S. Agizza ⁱ, A. Benini ^j, L. Ferranti ^k,
M. Gasparo Morticelli ^a, C. Giarrusso ^l, G. Mastronuzzi ^m, C. Monaco ^e, A. Porqueddu ⁿ

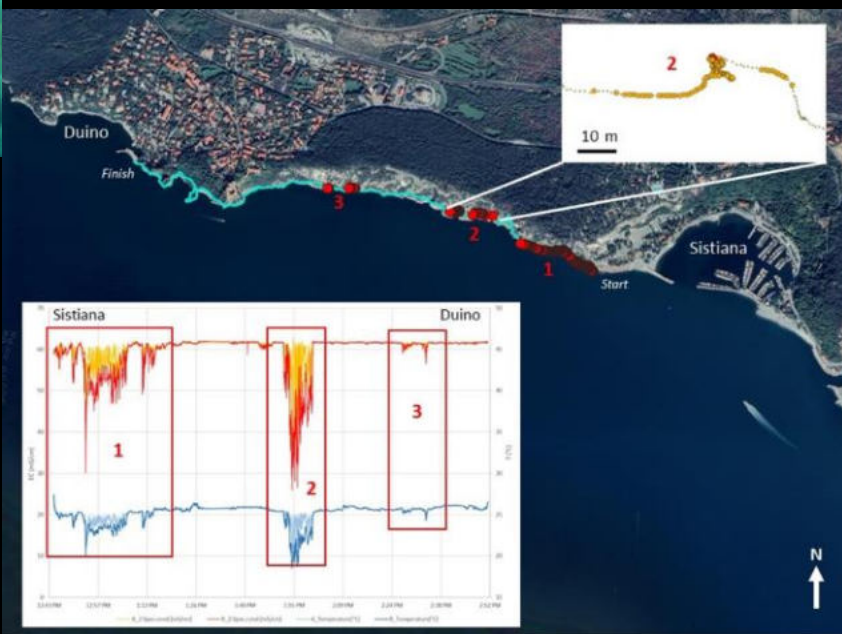
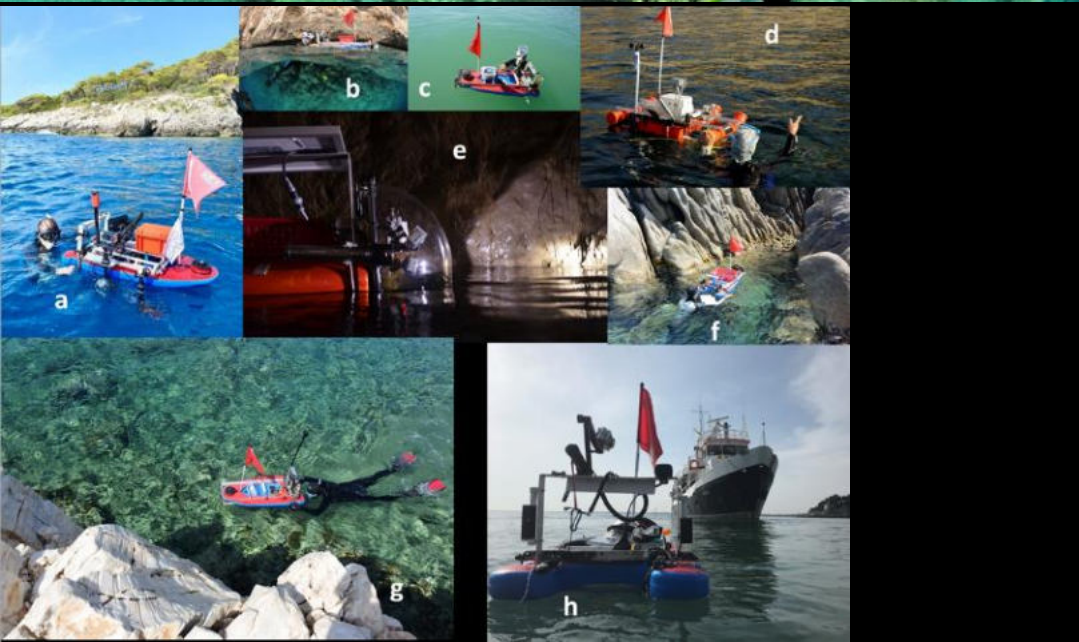
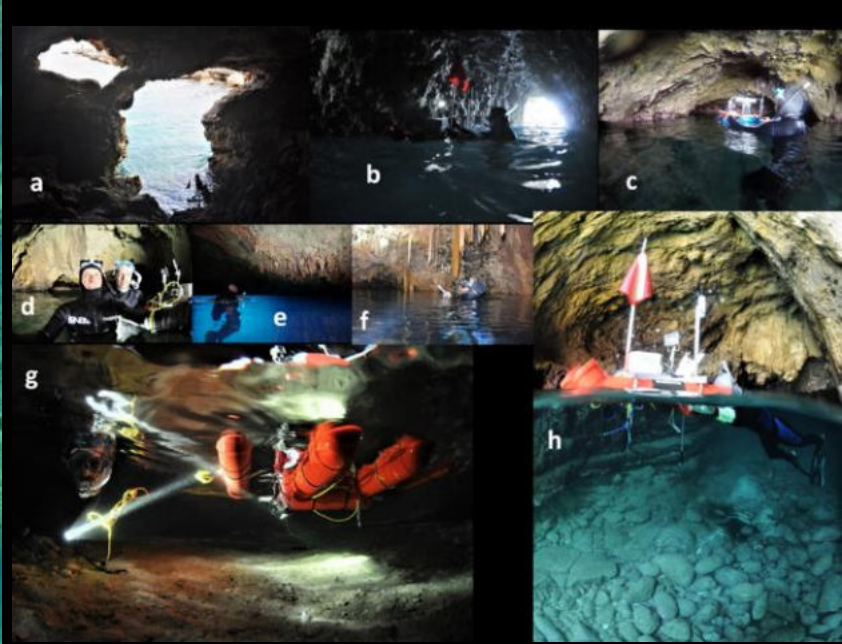
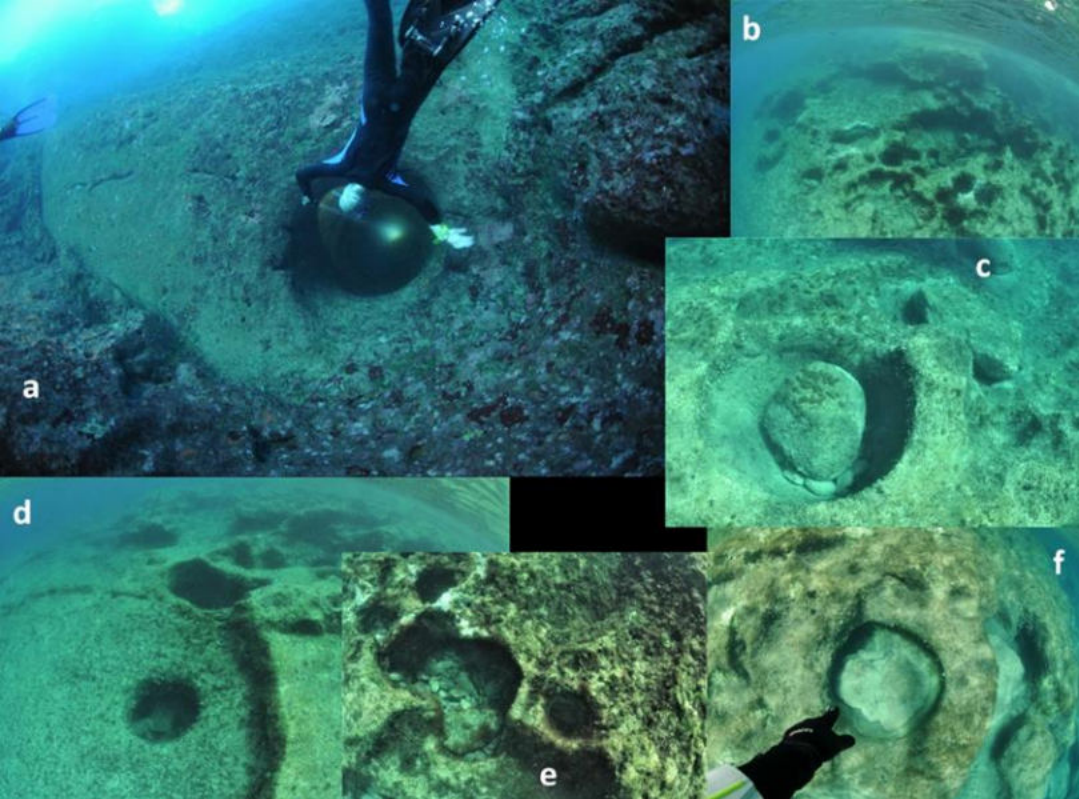




Geoswim a
Paros,
Grecia, 2017







A ID	B Year	C Location	D Videos (Y/N, a/b/w, duration (hours))	E Time-lapse images (Y/N, a/b/w)	F Outline images (Y/N, a/b)	G Total length (km)	H Days of survey	I Literature
1	2012	W Istrian Peninsula (Croatia, Slovenia, Italy)	(Y, a/b)	(N)	(Y, a/b)	253,2	27	Furlani, 2012; Furlani et al. 2014a
2	2013	Gozo and Comino (Malta)	(Y, a/b)	(N)	(Y, a/b)	57	7	Furlani et al., 2017a
3	2013	Stara Baska (Krak, Croatia)	(Y, a/b)	(N)	(Y, a/b)	2,2	1	/
4	2014	Egadi Islands (Italy)	(Y, w)	(N)	(Y, a/b)	67	7	Busetto et al., 2015 Furlani et al., 2021 Antonioli et al 2021
5	2015	Gaeta Promontory (Latium, Italy)	(Y, w)	(N)	(Y, a/b)	2,5	1	Furlani et al., 2021a
6	2015	Ustica (Sicily, Italy)		(N)	(Y, a/b)	14	2	Furlani et al., 2017b
7	2015	Razzoli, Budelli, Santa Maria (Sardinia, Italy)	(N)	(Y, w)	(Y, a/b)	22,5	3	Furlani et al., 2021a
8	2015	Capo Caccia (Sardinia, Italy)	(N)	(Y, w)	(Y, a/b)	26	2	Furlani et al., 2021a
9	2015	Tavolara (Sardinia, Italy)	(N)	(Y, w)	(Y, a/b)	14,9	2	Furlani et al., 2021a
10	2015	Malta (Malta)	(N)	(Y, w)	(Y, a/b)	19,2	3	Furlani et al., 2021a
11	2015	SE Istria (Croatia)	(N)	(Y, w)	(N)	7	1	Vaccher, unpublished thesis
12	2016	Monte Conero (W Adriatic Sea, Italy)	(N)	(Y, w)	(Y, a/b)	2,9	1	Furlani et al., 2018
13	2016	Addaura (Palermo, Sicily, Italy)	(N)	(N)	(Y, a/b)	7	1	Caldareri et al., 2018
14	2017	Paros (Greece)	(Y, w)	(Y, w)	(Y, a/b)	24	8	Furlani et al., 2021b
15	2017	Sistiana-Duino (Gulf of Trieste, Italy)	(N)	(Y, a/b)	(N)	2	1	Furlani and Biolchi, 2018
16	2018	Ansedonia and Argentario (Tuscany, Italy)	(Y, w)	(Y, a/b)	(Y, a/b)	10	2	Furlani et al., 2021
17	2019	Savudrija (Croatia)	(N)	(Y, a/b)	(N)	1.1	1	Furlani et al., 2021
18	2020	Isole Pelagie (Sicily, Italy)	(Y, a)	(Y, a/b)	(Y, a/b)	43.9	9	/
19	2020	Arenella (Sicily Italy)	(N)	(n, a/b)	(Y, a/b)	0.6	1	Furlani et al., 2021 Antonioli et al. 2021
20	2021	Isole Tremiti (Puglia, Italy)	(Y, a/b)	(Y, a/b)	(Y, a)	5.7	1	/
	TOTAL					532	81	



Article
**Preserv
Biologi**


Stefano Furl
Alice Busetti
Elisa Dal Bo
Thalassia Gi
Alessandro I
Matteo Vaccl




Citation: Furlani,
Antonioli, F.; Agab
Boccali, C.; Busetti
Canziani, F.; Chem
Preservation of Me
Erosional Landform
Structures as Sea L
Matter of Luck? W
<https://doi.org/10>



40 coastal area that could be flooded by the sea for 2100

 completed maps

 maps to finish



FLOODING SCENARIO AT FOUR ITALIAN COASTAL PLAINS USING THREE RELATIVE SEA LEVEL RISE MODELS: THE NORTH ADRIATIC AREA



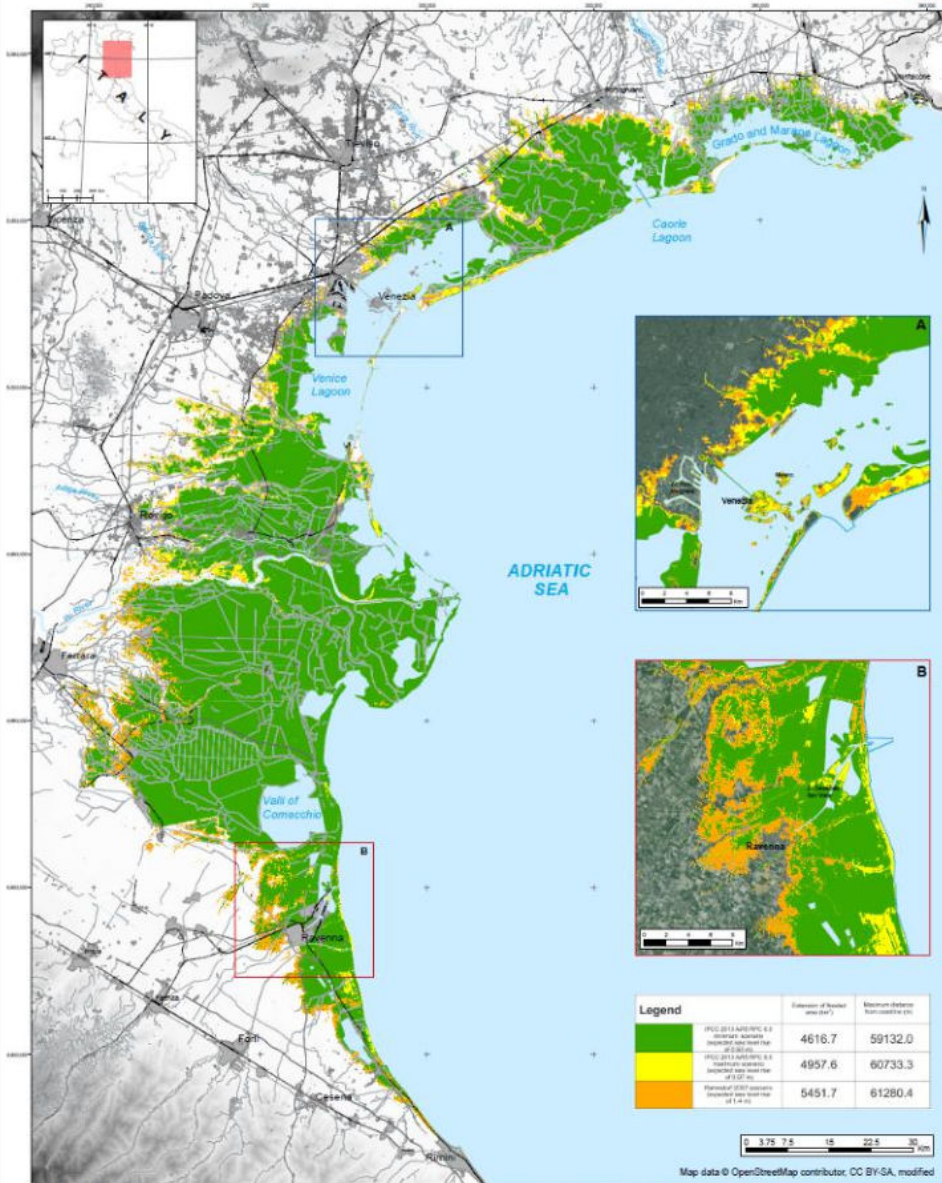
A. Marsico¹, S. Lisco¹, V. Lo Presti², F. Antonjoli², A. Amorosi³, M. Anzidei⁴, G. Delana⁵, G. De Falco⁶,
A. Fontana⁷, G. Fontolan⁸, M. Moretti¹, P. Orru⁹, G. Sannino⁹, E. Serpelloni⁴, A. Vecchio¹⁰, G. Mastroruzzi¹

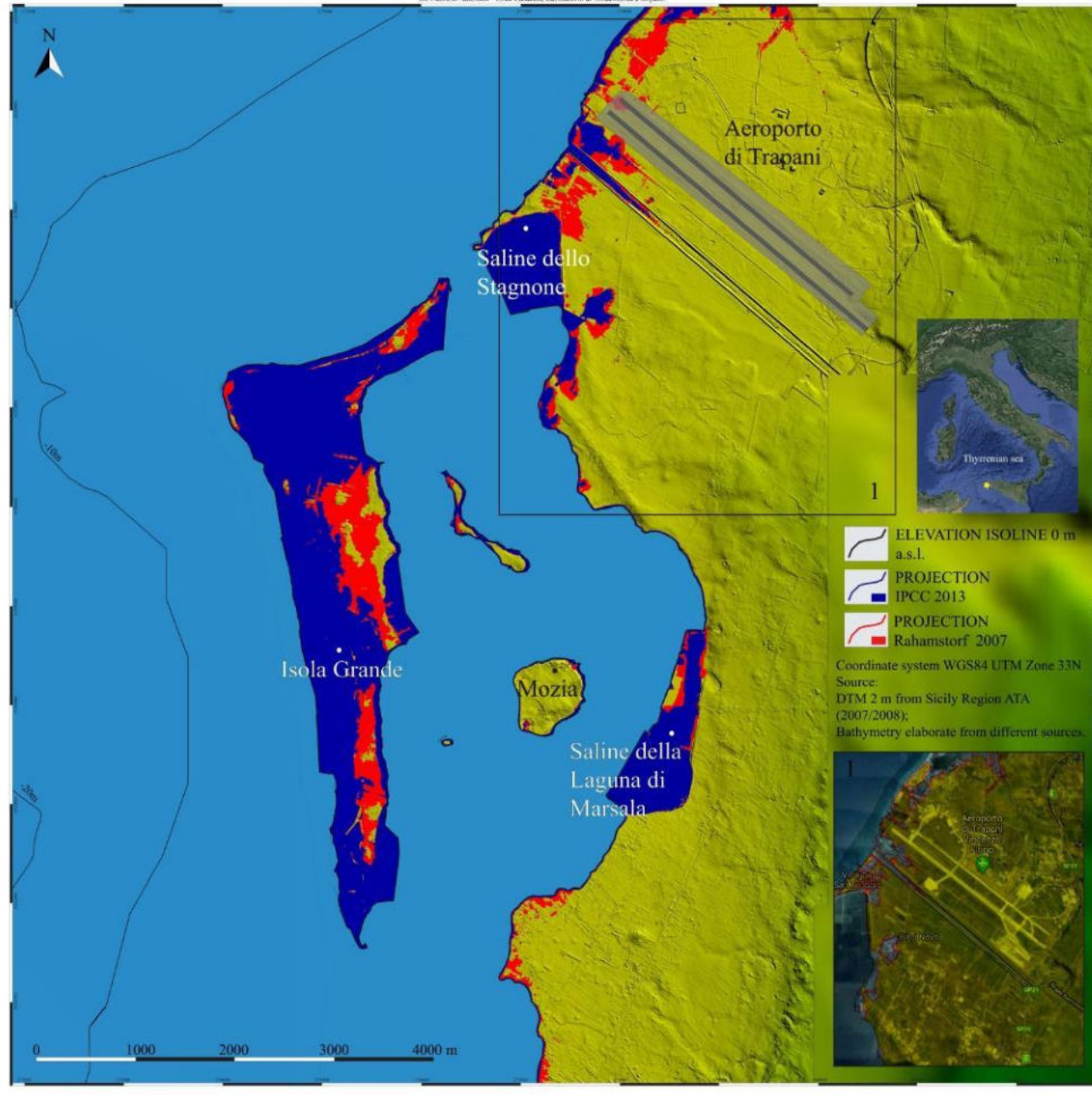
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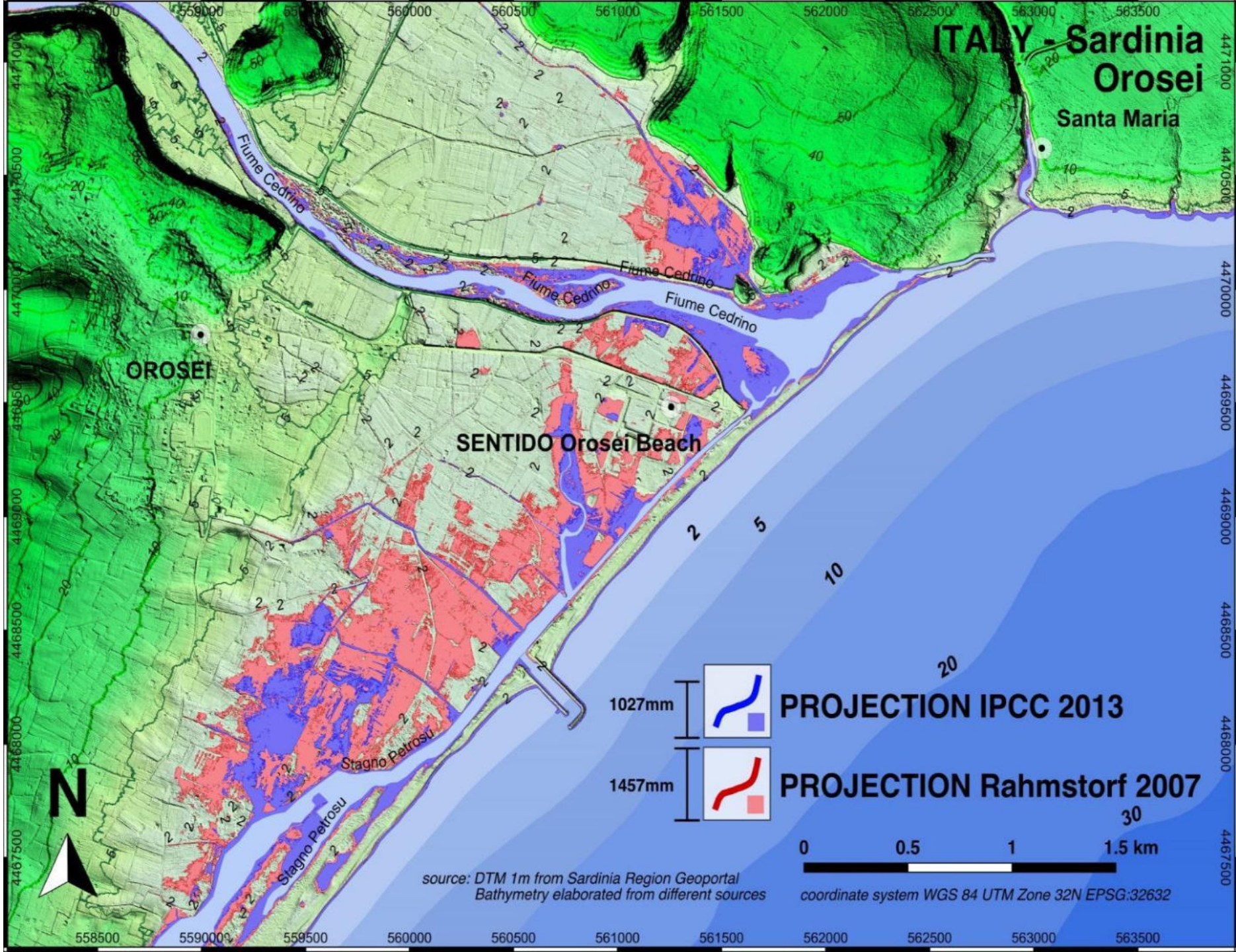
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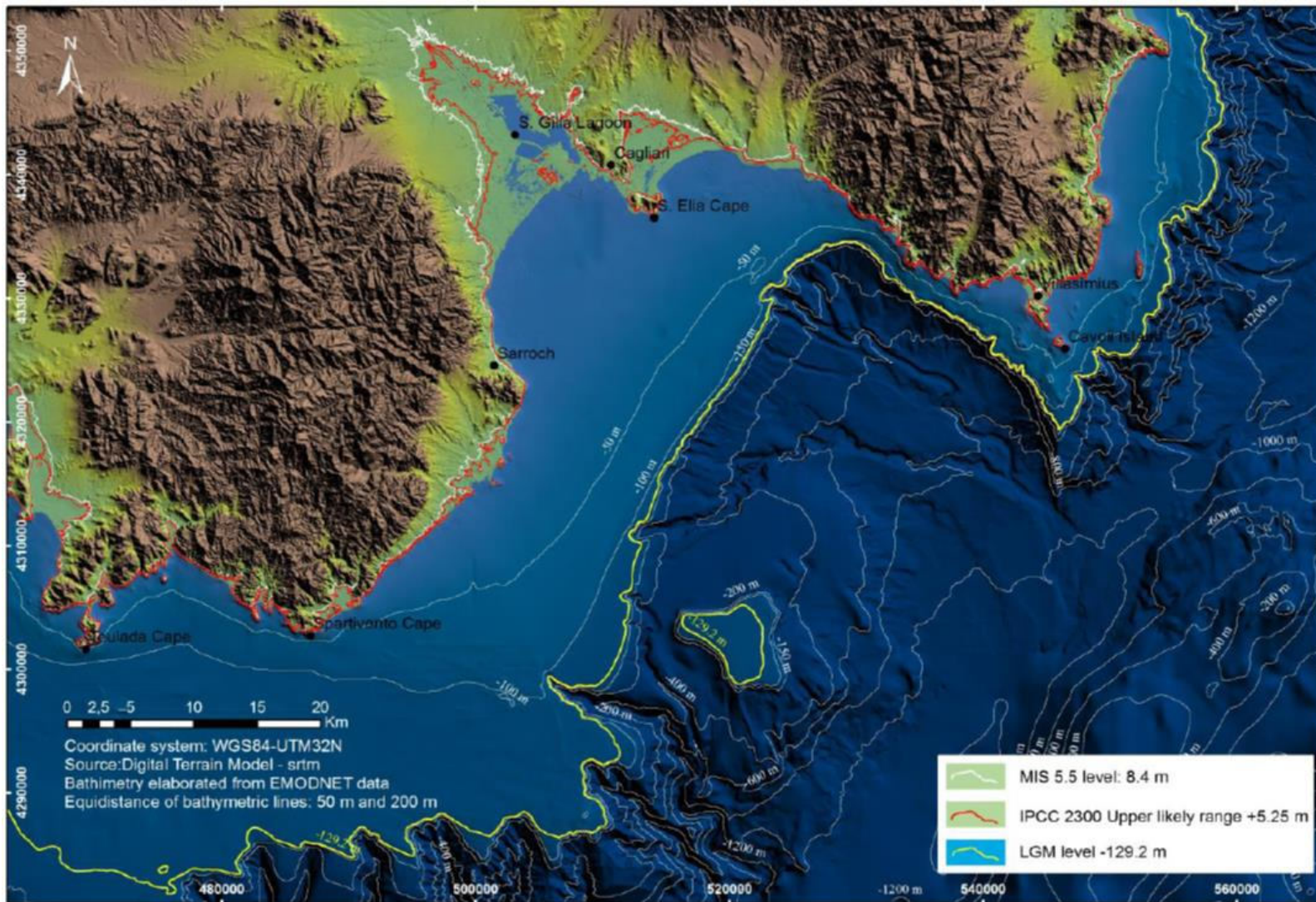


Figure 16 The Cagliari plane map showing the potential submersion area using IPCC AR5 RCP 8.5 for 2100 and 2300: The MIS 5.5 extension occurred 119 ka BP.

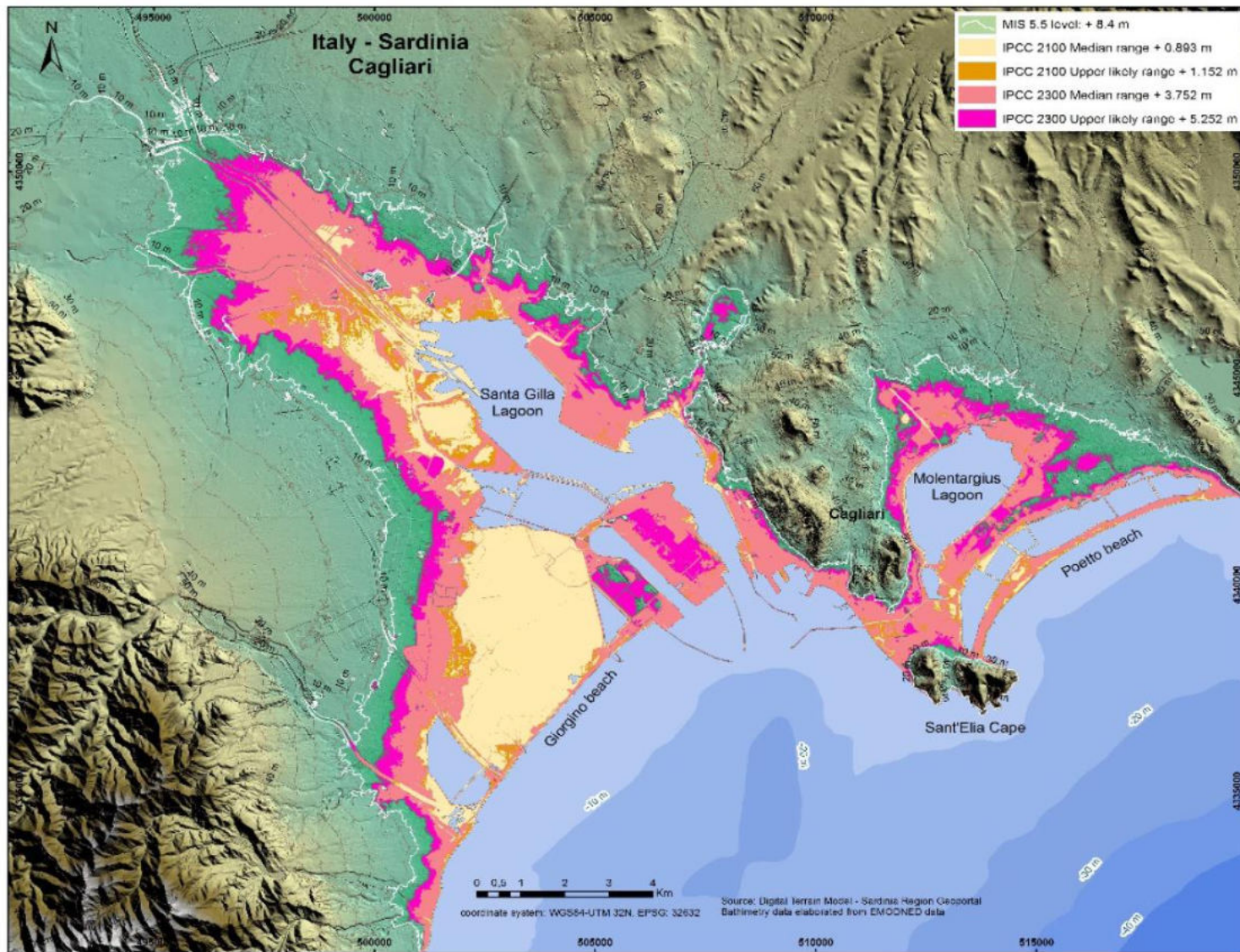


Figure 17. Map of Cagliari plane, (see also Figure 2 for location). The potential submersion area, using IPCC AR5 RCP 8.5 projections at 2100 and 2300.

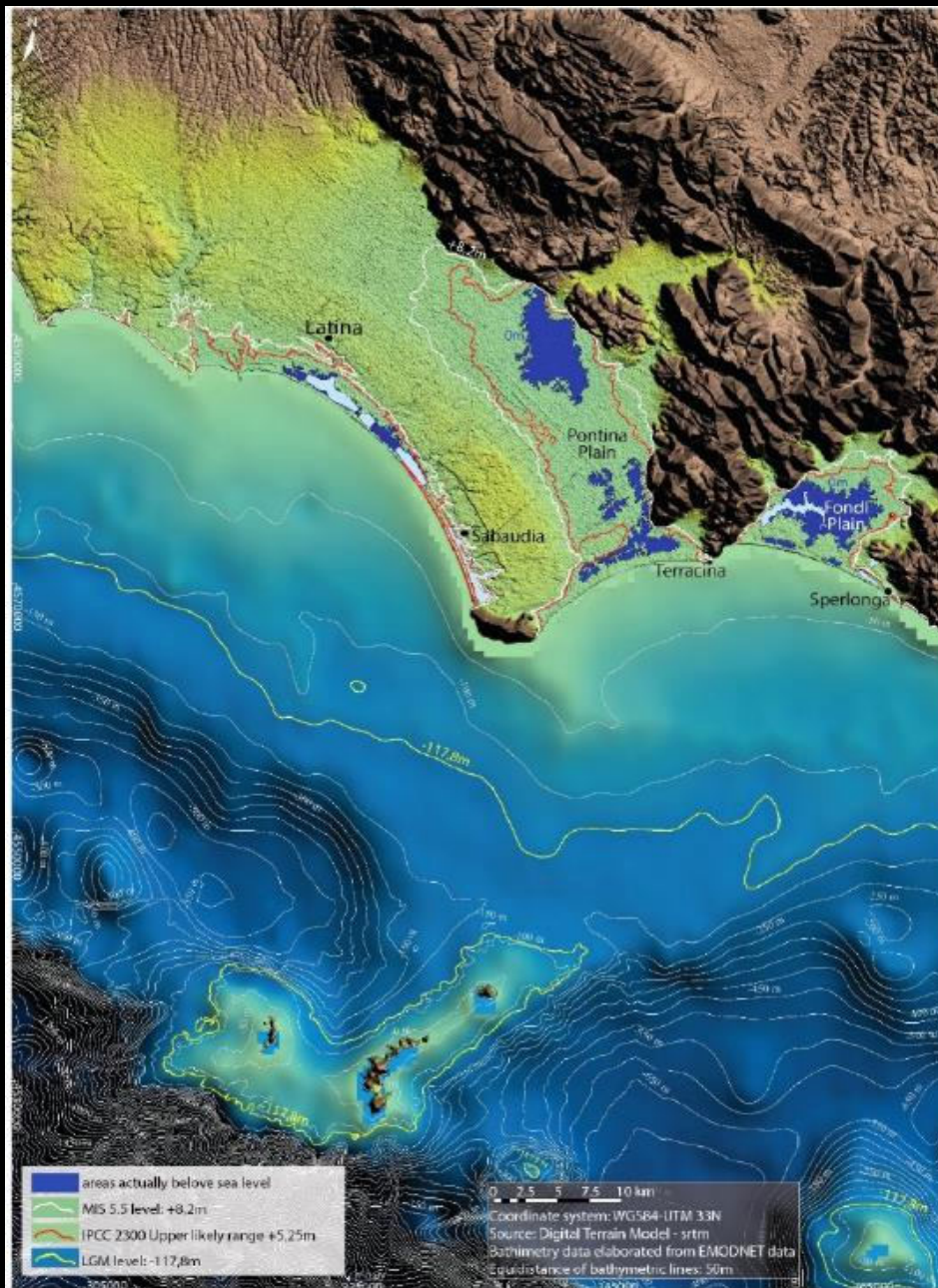


Figure 15. The Pontina and Fondi Plains maps showing the potential submersion area using IPCC AR5 RCP 8.5 for 2100 and 2300: the MIS 5.5 extension occurred 119 ka BP.



Chloe Sladden



Farrell e Mortensen interpretano
due dei **soccorritori**
intervenuti per salvarli.

00:51 / 01:25

"TREDICI VITE", IL FILM CHE HA SCATENATO IL PANICO DI COLIN FARRELL







Thalassia Giaccone

A film by Fabrizio Antonioli and Thalassia Giaccone



2100

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project management FLAVIO LENOCI
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